

**Introdução a Física  
Atmosférica - PGF-3521  
Aula 3 – Water in the  
climate system**

*Henrique Barbosa*  
*Departamento de Física Aplicada*  
*Instituto de Física da USP*  
IF-USP, Ed. Milênio, Sala 4  
R. 91 8985  
[\*hbarbosa@if.usp.br\*](mailto:hbarbosa@if.usp.br)

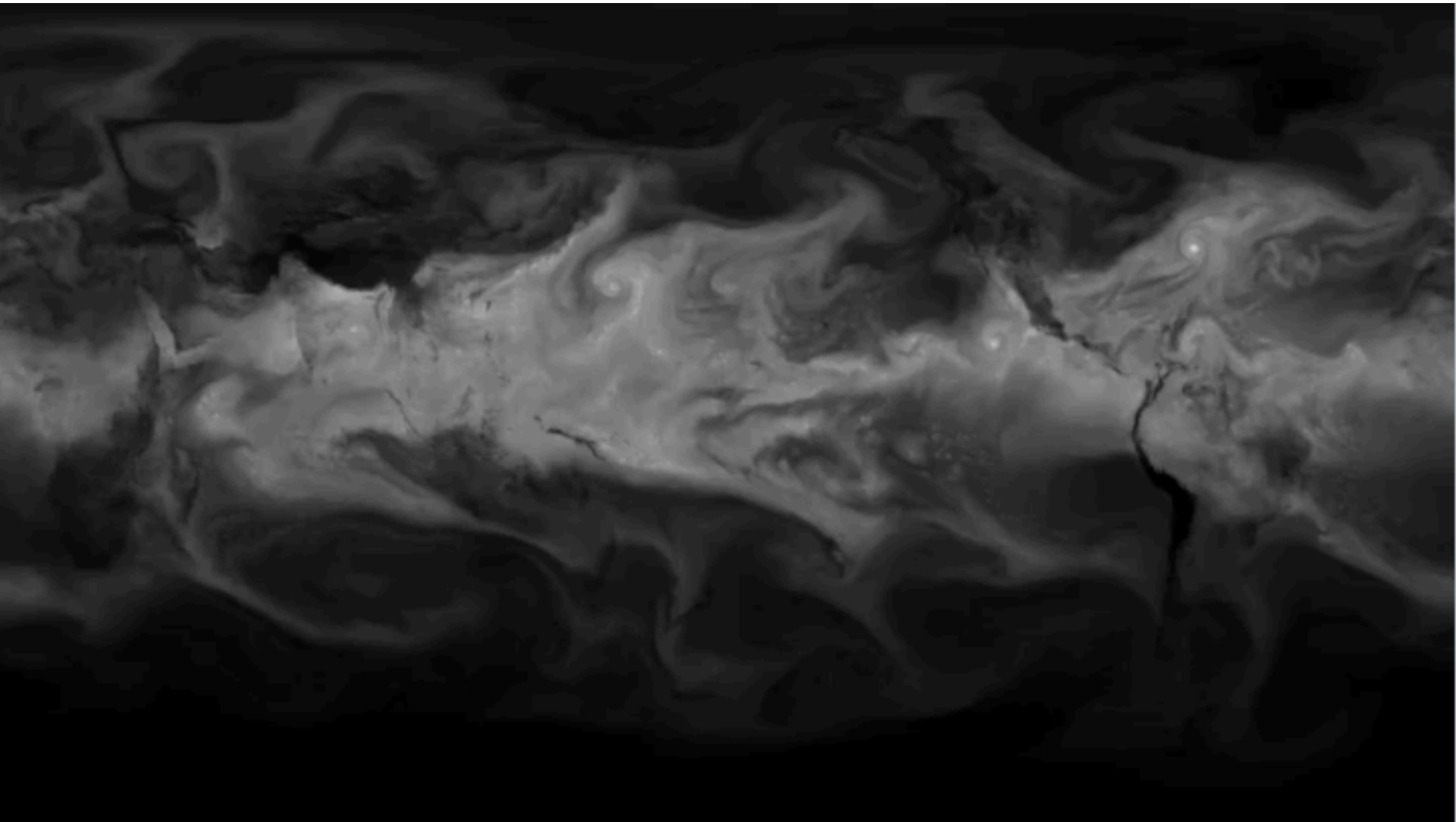


# Water in the climate system

Its physical properties determine

- How strong greenhouse effect is;
- Planetary albedo;
- Thermodynamic structure of the troposphere;
- Large scale circulation;
- Hydrological cycle;

# Total column water vapor



# Water is rare

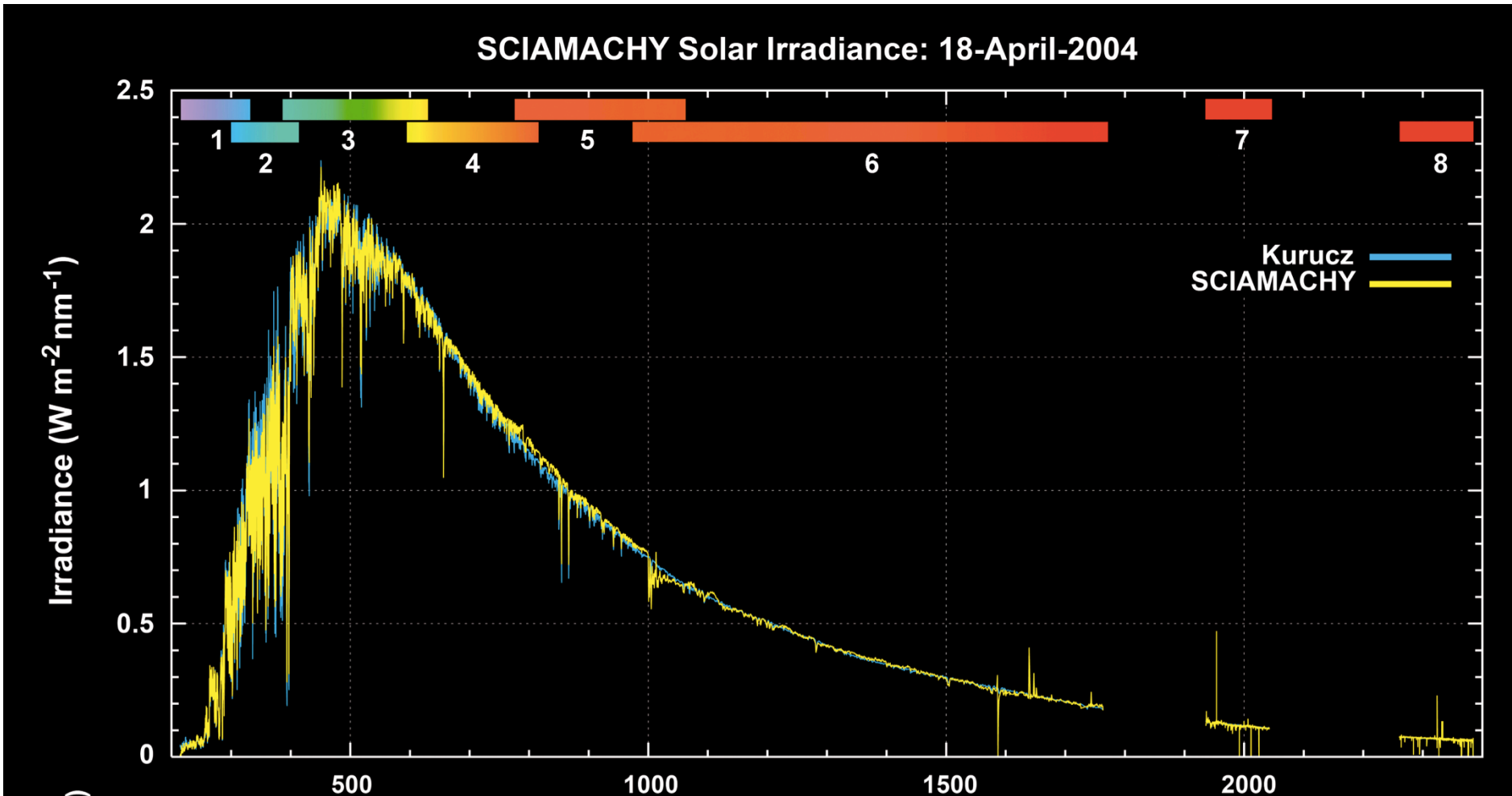
Water in the atmosphere, be it vapor, liquid or solid, is the least likely place to find a water molecule in the climate system.

<b>H<sub>2</sub>O</b>	<b>Equivalent in meters</b>
Liquid or solid water in the atmosphere	0.0001 m
Vapor in the atmosphere	0.025 m
Water in soil, lakes, rivers and glaciers	50 to 75 m
Oceans	2800 m

Its important role steams from its radiative properties



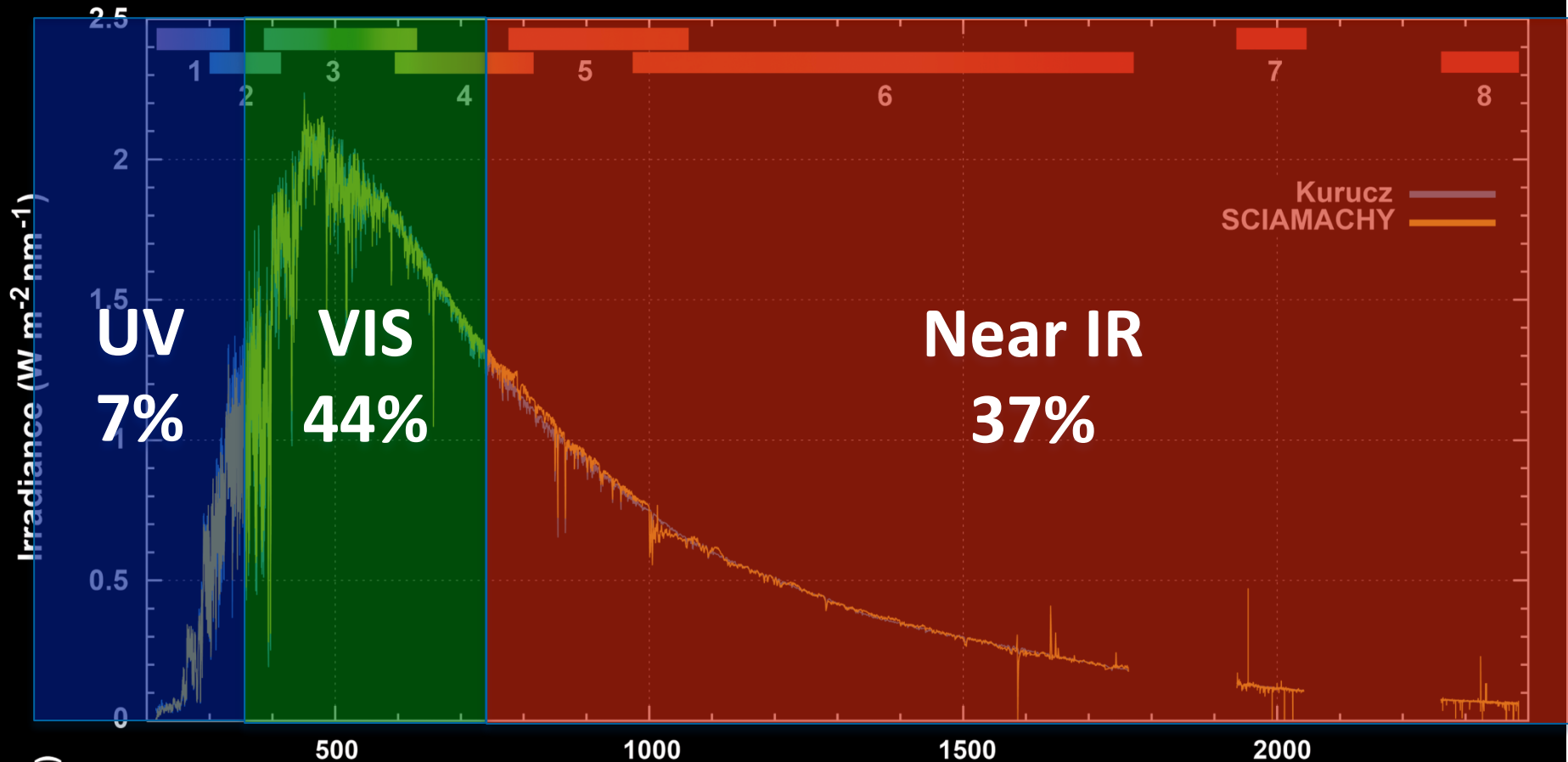
# Observed Solar Spectrum



BOOK: SCIAMACHY - Exploring the Changing Earth's Atmosphere  
<http://atmos.caf.dlr.de/projects/scops/>

$$S_0 \sim 1365 \text{ W/m}^2$$

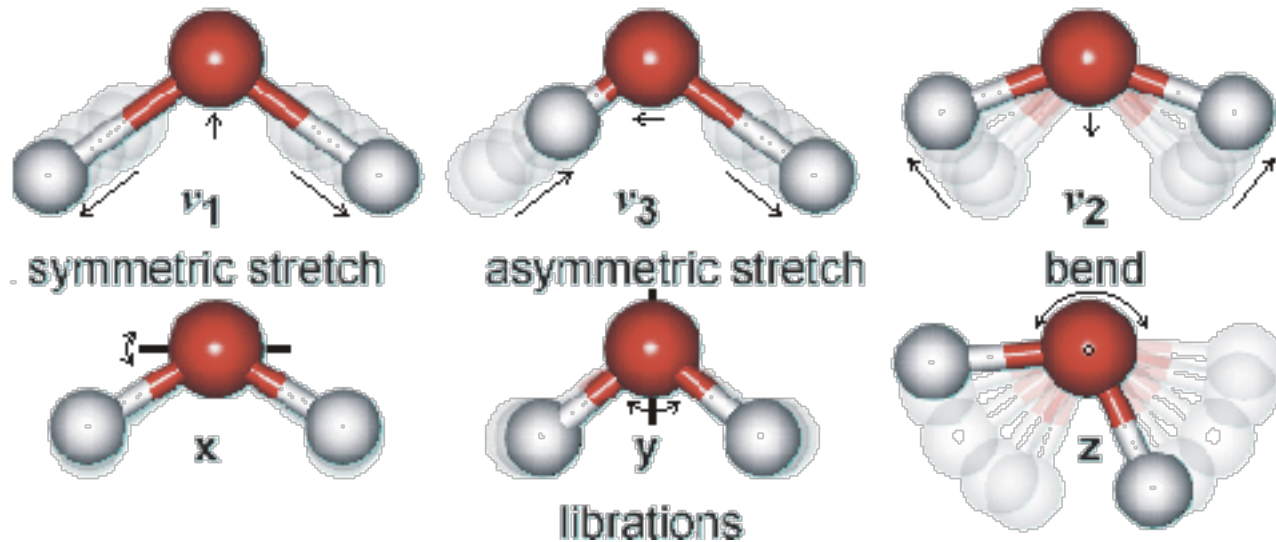
SCIAMACHY Solar Irradiance: 18-April-2004



BOOK: SCIAMACHY - Exploring the Changing Earth's Atmosphere  
<http://atmos.caf.dlr.de/projects/scops/>

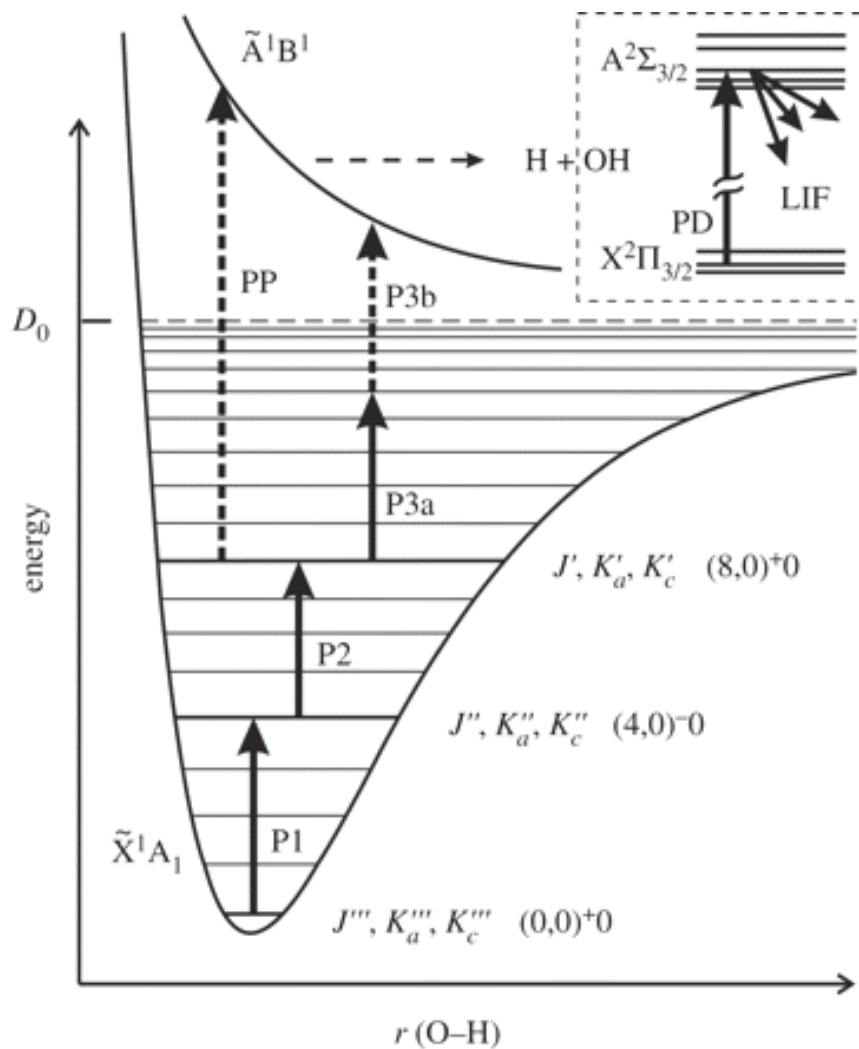
# H<sub>2</sub>O vibration and rotation

- The water molecule has a very small moment of inertia on rotation which gives rise to rich combined vibrational-rotational spectra in the vapor containing tens of thousands to millions of absorption lines.

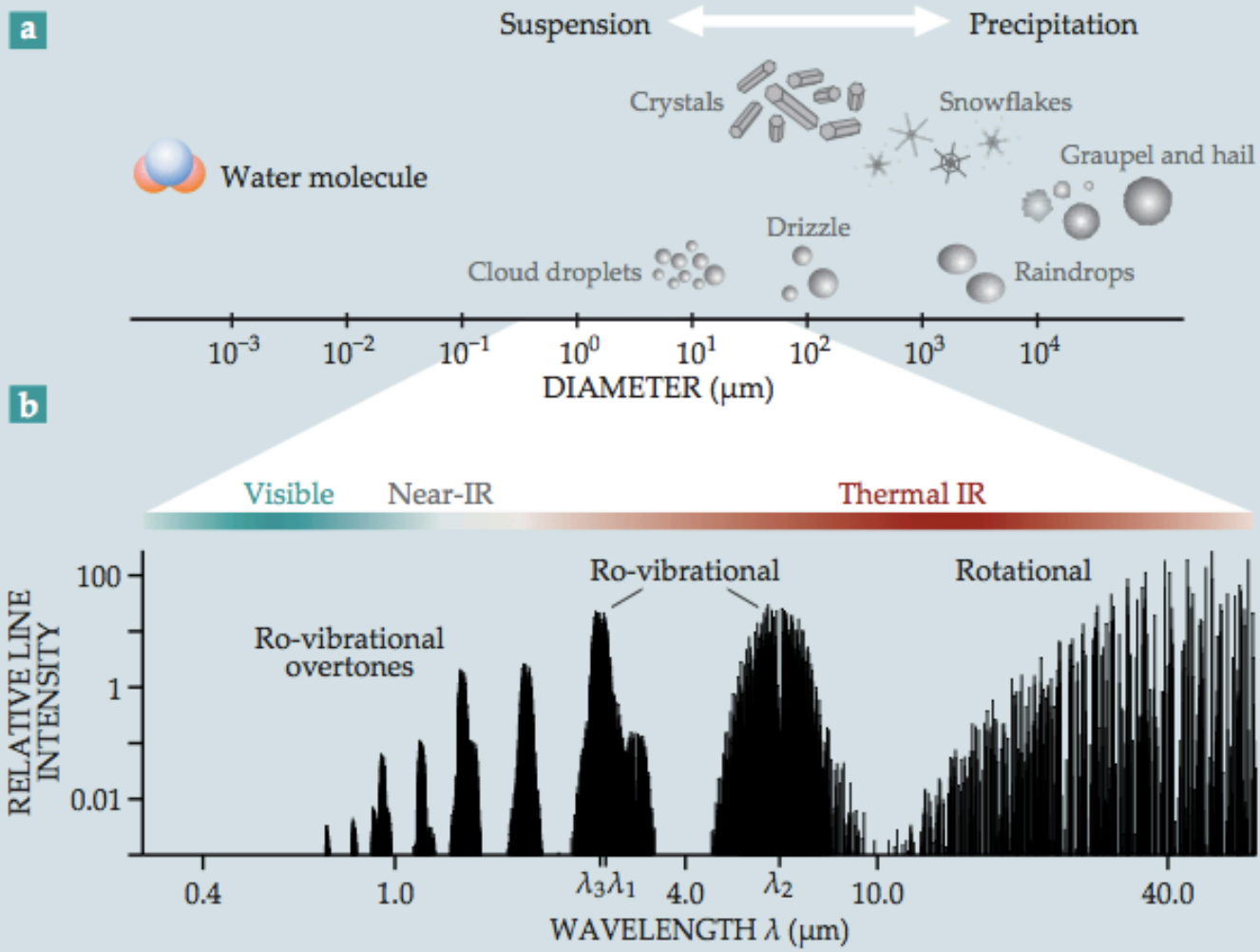




# Energy levels



- Energy-level diagram of double- and triple-resonance vibrational overtone excitation (photons P<sub>1</sub>–P<sub>3a</sub>) followed by photodissociation (dashed arrows, photons P<sub>3b</sub> or PP) and OH fragment



**Figure 1. Hydrometeors (a)**, the condensed forms of water in the atmosphere, come in several sizes. They mostly scatter visible light but absorb over a broad range of the IR. **(b)** The near- and thermal-IR regions of the spectrum excite the molecule and produce its rotational-vibrational (or ro-vibrational) and rotational bands. Specific lines  $\lambda_1$ ,  $\lambda_2$ , and  $\lambda_3$  mark the symmetric stretching mode, bending mode, and asymmetric stretching mode, respectively.

# Radiation Transmitted by the Atmosphere

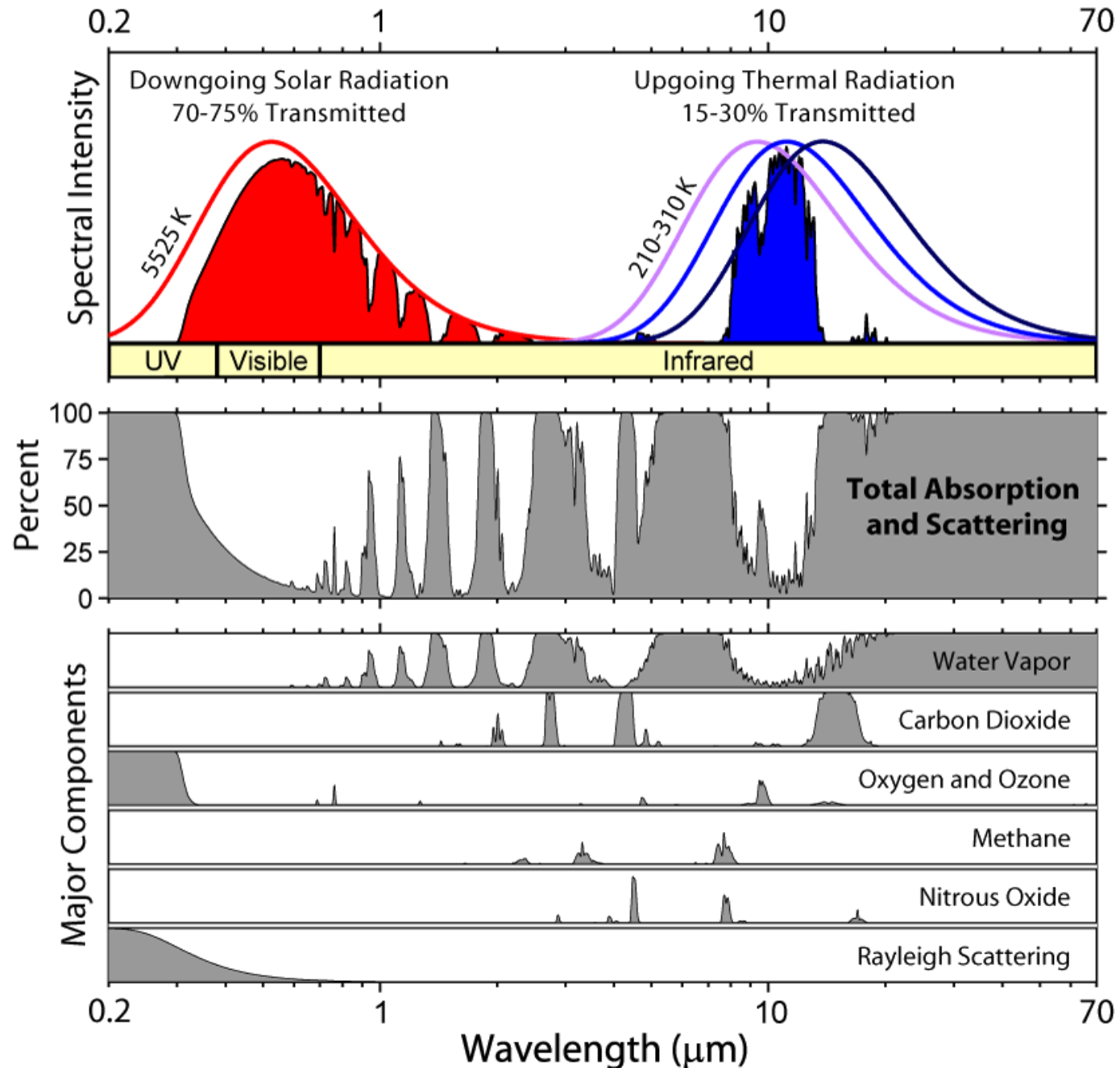
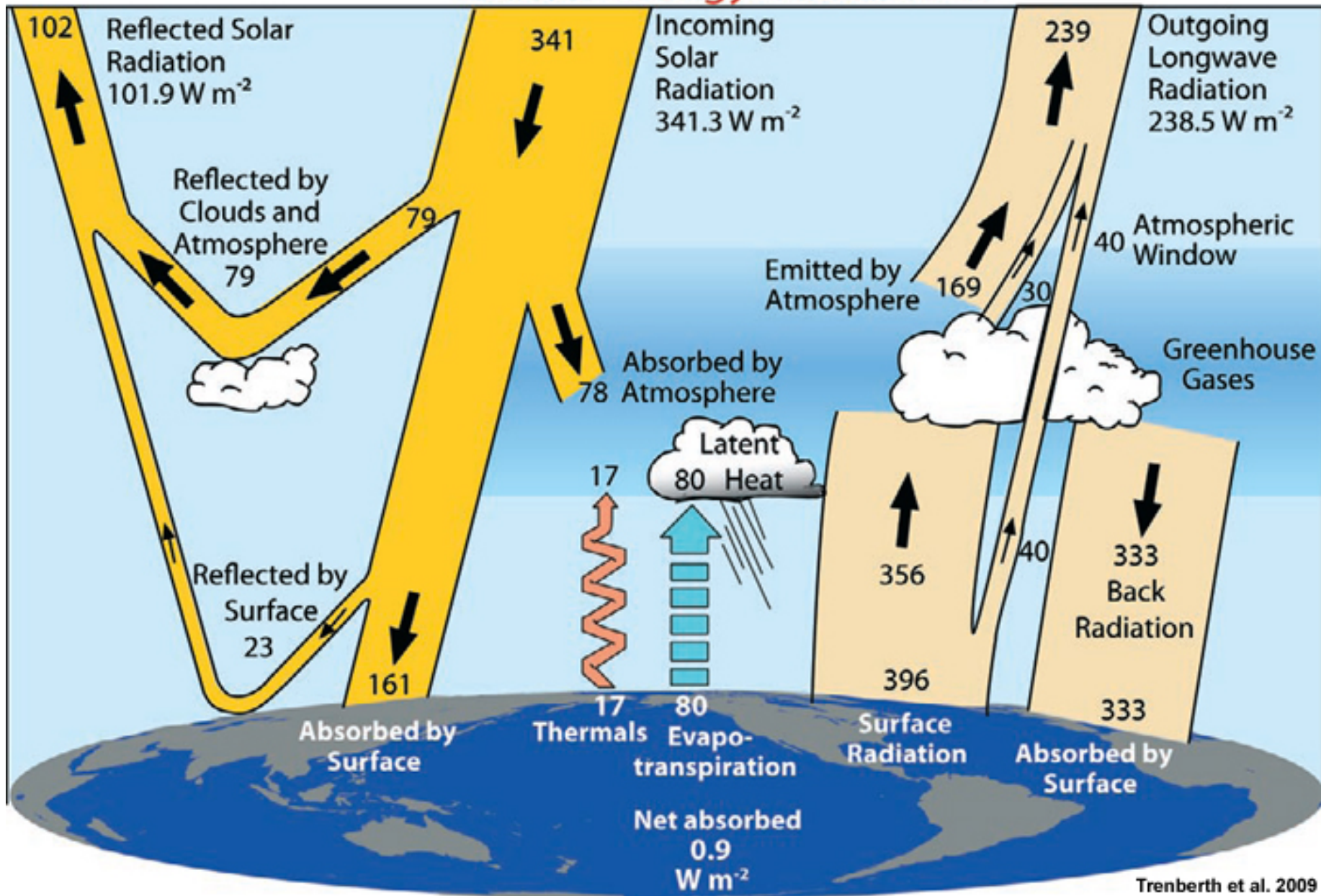


Image by:  
Robert A. Rohde  
Global Warming Art.



# Global Energy Flows $W m^{-2}$



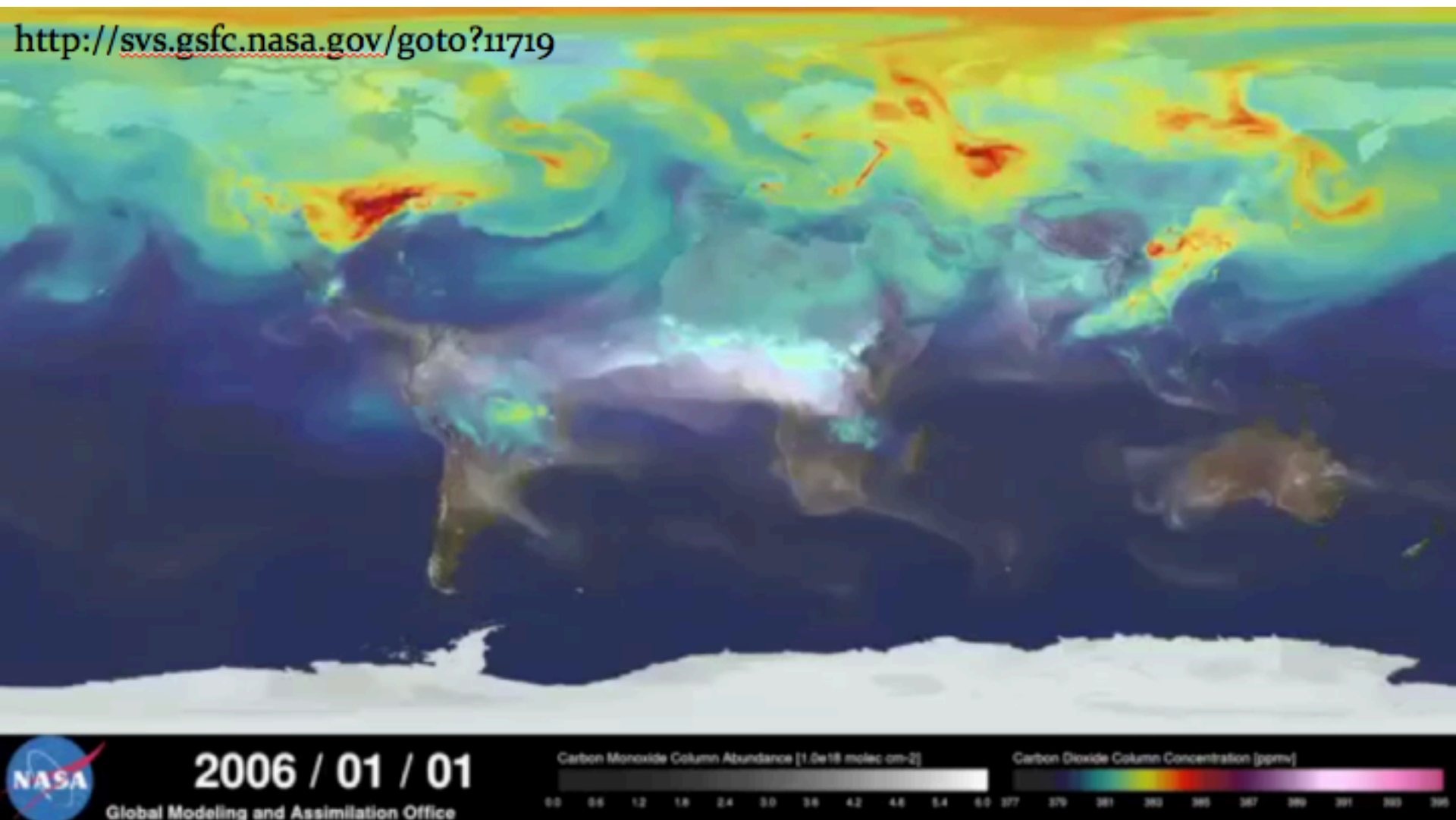
# Natural Greenhouse effect

	Radiative efficiency (W m <sup>-2</sup> /ppb)	Pre industrial conc.	Natural Greenhouse effect (W m <sup>-2</sup> )		Concentration in 2011	Antrop. Forcing (W m <sup>-2</sup> )
H <sub>2</sub> O			75	51		
CO <sub>2</sub>	1.37 10 <sup>-5</sup>	278±2 ppm	32	24	390.4±0.2 ppm	1.82
O <sub>3</sub>			10	7		0.35
CH <sub>4</sub>	3.63 10 <sup>-4</sup>	722±25 ppb	8	4	1803.2±1.2 ppb	0.48
N <sub>2</sub> O	3.03 10 <sup>-3</sup>	270±7 ppb			324.3±0.1 pbb	0.17
CF <sub>4</sub>	0.1	34.7±0.2 ppt			79.0±0.1 ppt	0.0041
Outros						0.01
Total			125	86		2.83

Hartmann et al, IPCC (WG-I) 2013  
Kiehl and Trenberth, BAMS, 1997

H<sub>2</sub>O is not a forcing because it is not directly emitted as CO<sub>2</sub>

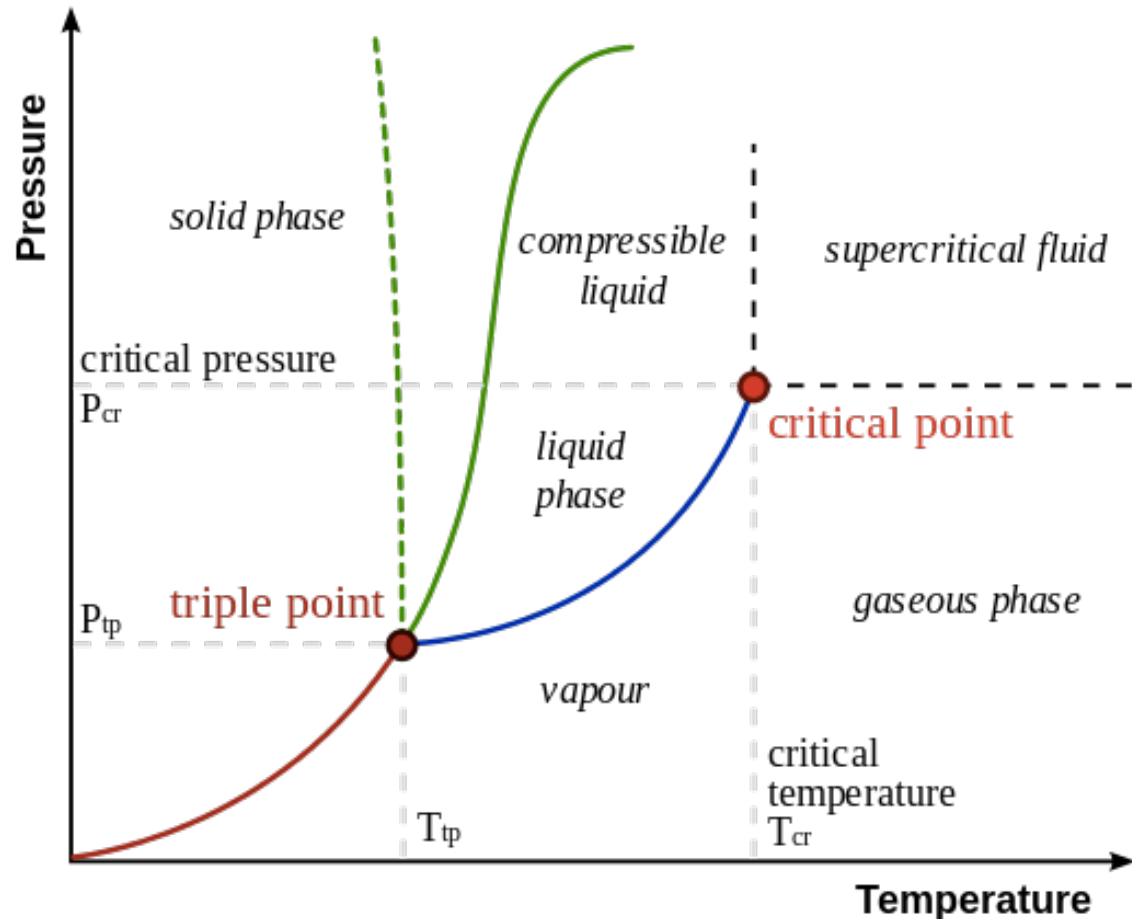
<http://svs.gsfc.nasa.gov/goto?11719>





# Vapor phase diagram

- Vapor refers to a gas phase at a temperature where the same substance can also exist in the liquid or solid state, below the critical temperature of the substance.
- $T_{cr}(H_2O) = 374\text{ }^\circ\text{C}$
- $T_{tp}(H_2O) = 0.01\text{ }^\circ\text{C}$



# Water vapor partial pressure

Clausius-Clapeyron:

$$\frac{de_s}{dT} = \frac{L_v(T)e_s}{R_v T^2}$$

Saturation vapor pressure:

$$e_s(T) = 6.1094 \exp\left(\frac{17.625T}{T + 243.04}\right)$$



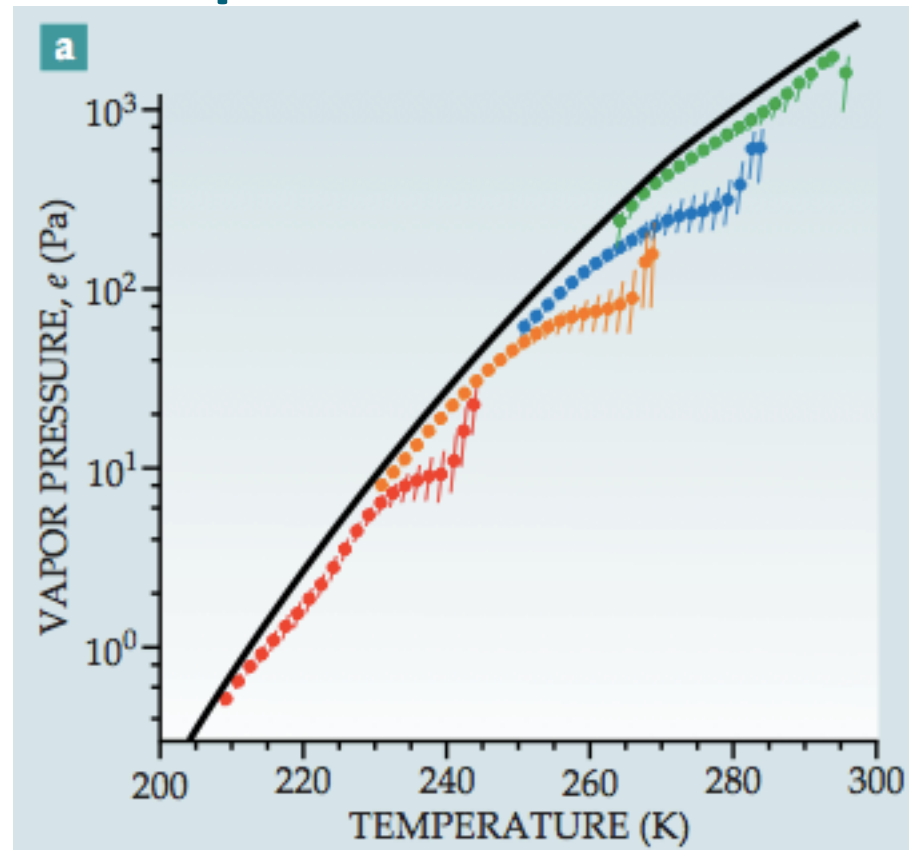
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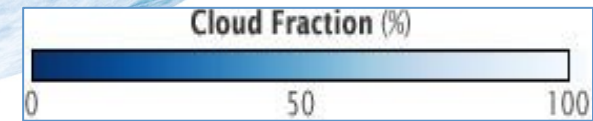
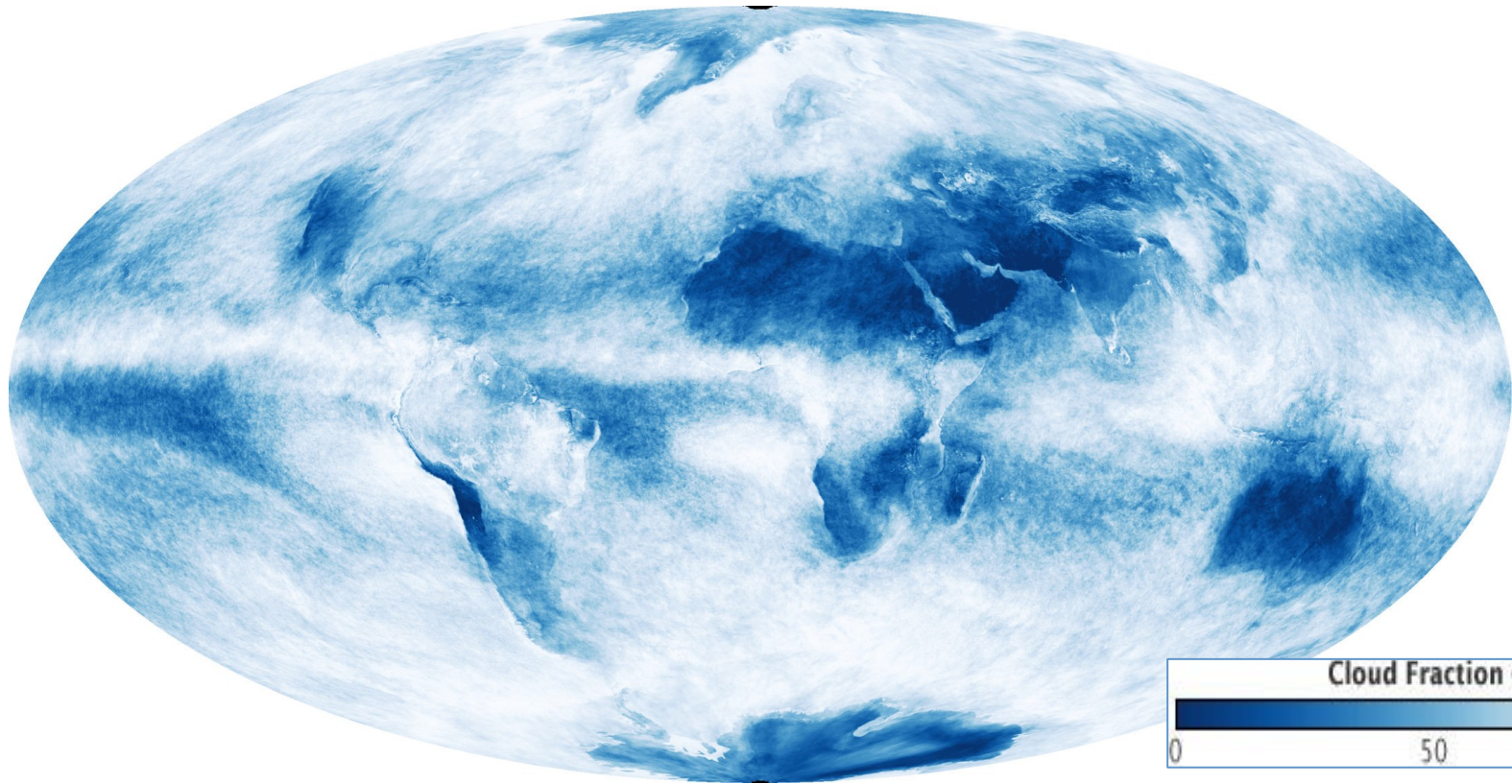
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**Figure 2.** The atmospheric vapor pressure as a function of temperature (a) is bounded by the atmosphere's vapor pressure when saturated with water (solid line). The data are shown as the median (dot) and range (error bar) of 228 monthly values at different isobaric levels—900 hPa (green), 700 hPa (blue), 500 hPa (orange), and 300 hPa (red). (Data are provided by the European Centre for Medium-Range Weather Forecasts.) (b) Based on satellite measure-

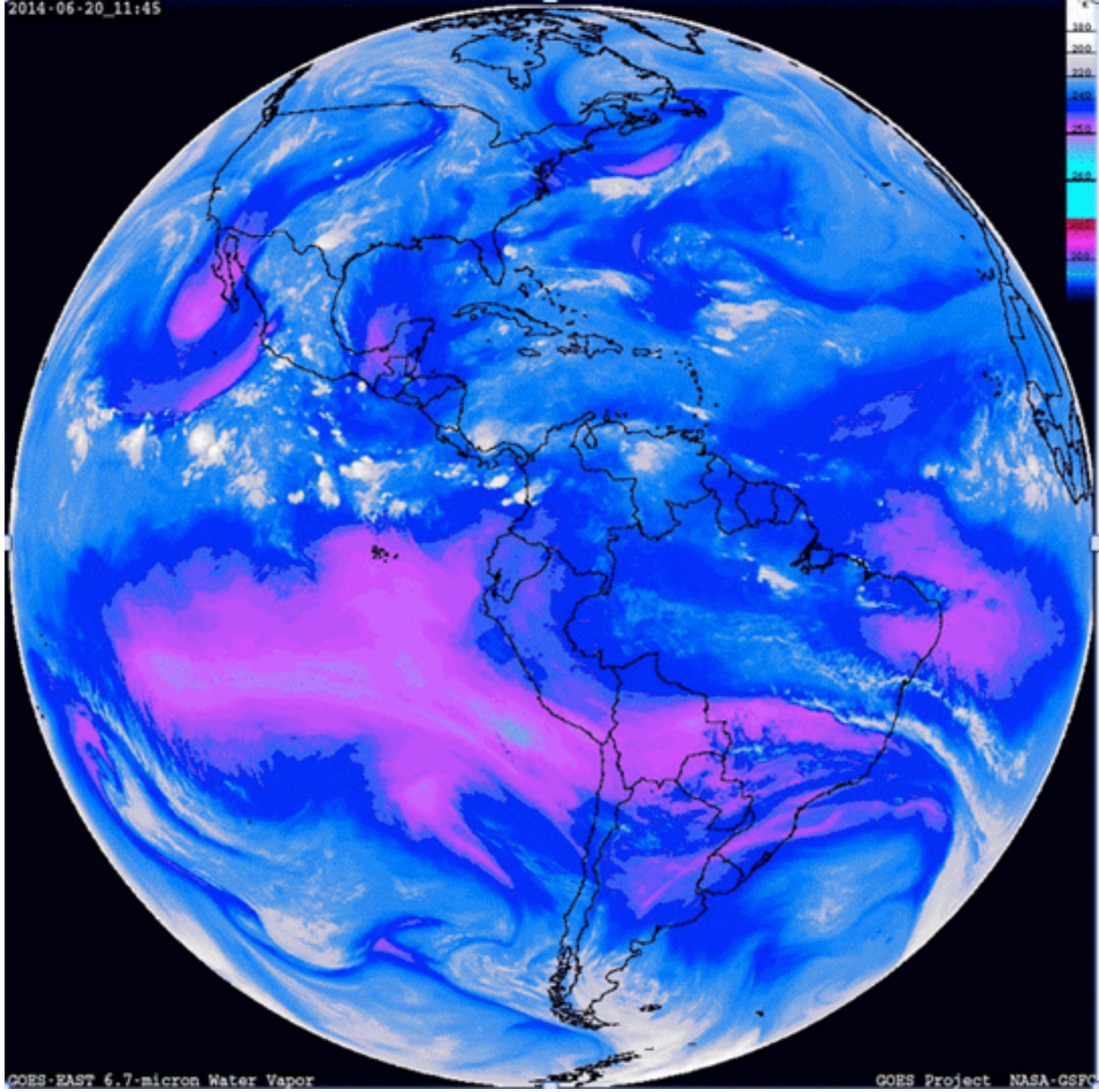




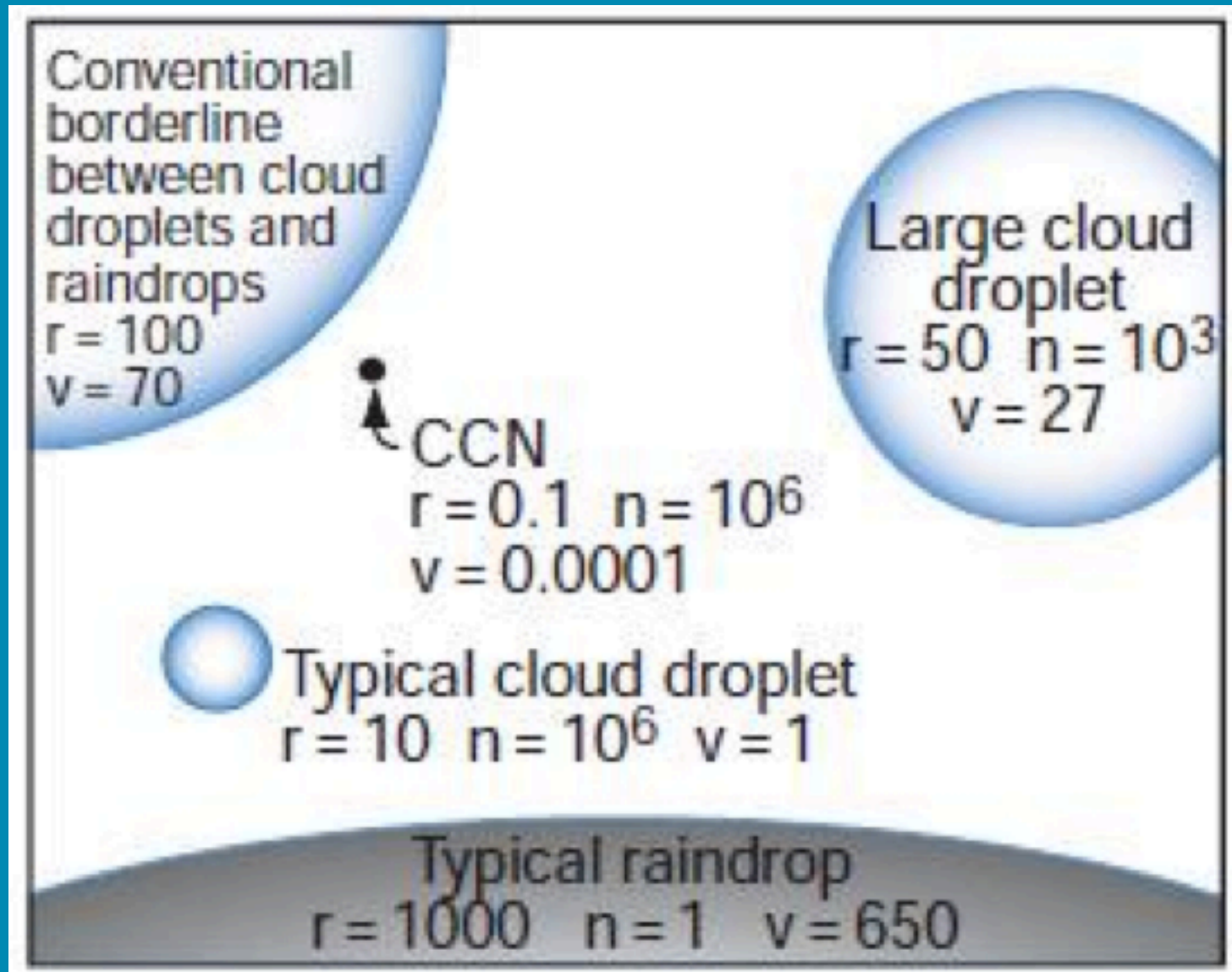
	Cloud density			All densities
	Thin	Thick	Opaque	
Cloud level	$N_e < 0.5$ $\sigma_{vis} < 1.4$	$0.5 < N_e < 0.95$ $1.4 < \sigma_{vis} < 6$	$N_e > 0.95$ $\sigma_{vis} > 6$	
High (<440 mb)	15%	15%	3%	33%
Middle (440–700 hPa)	7%	10%	9%	26%
Low (>700 hPa)		2%	47%	49%
Total	20%	23%	32%	75%



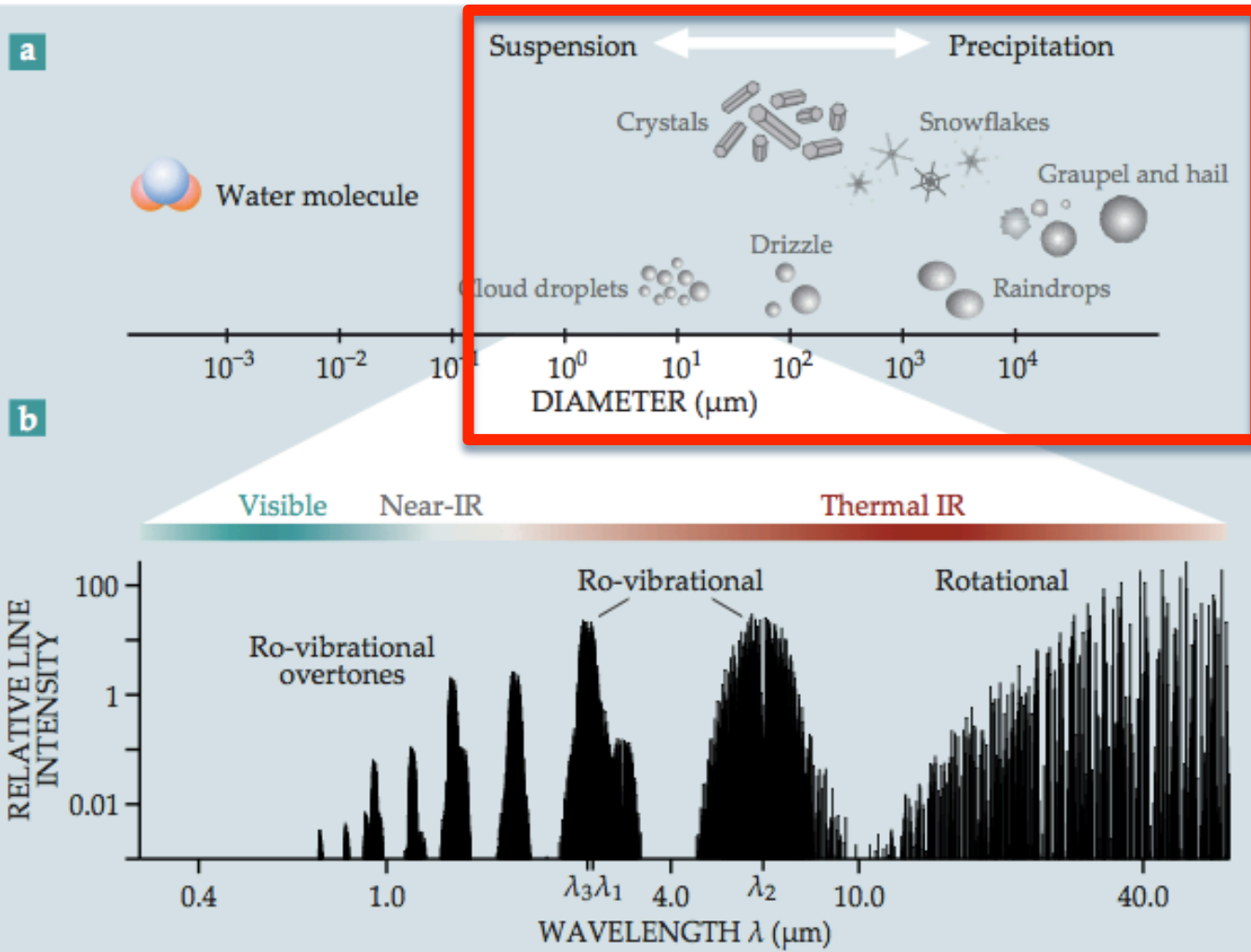
<http://www.nasa.gov/content/goddard/a-new-view-nasanoaa-water-vapor-animations-over-oceans>



# 6 orders of magnitude in volume



# Clouds (liq/ice) strongly interact with SW and LW radiation

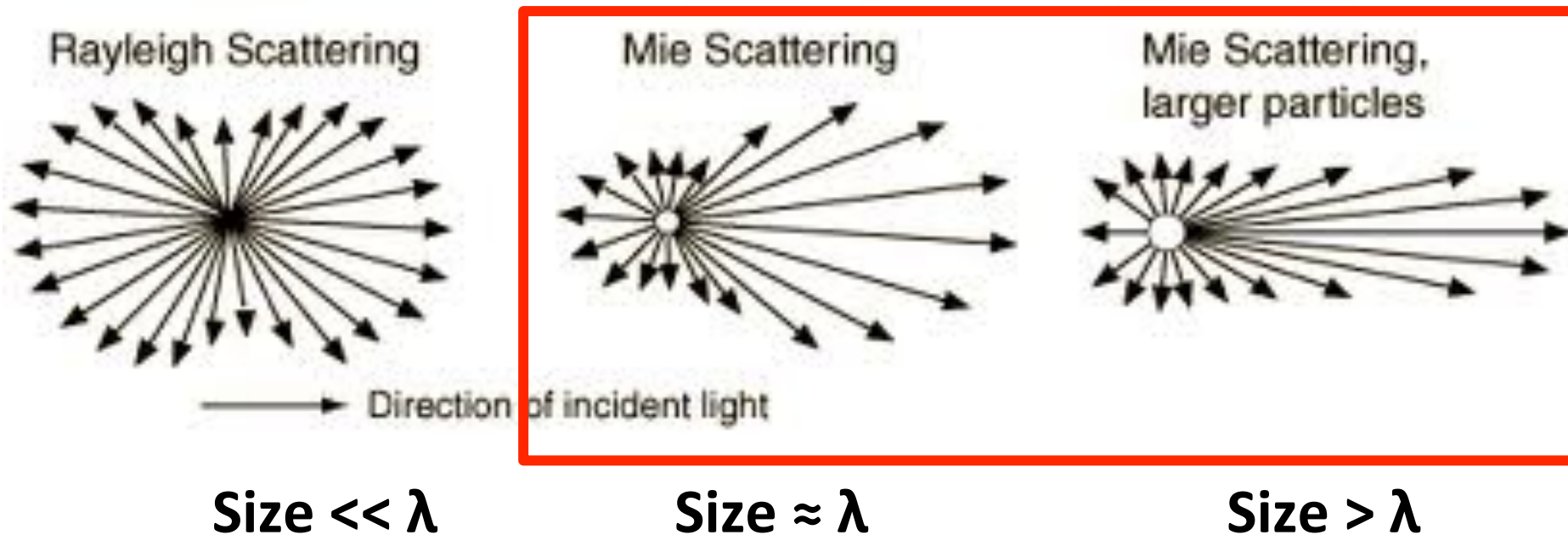


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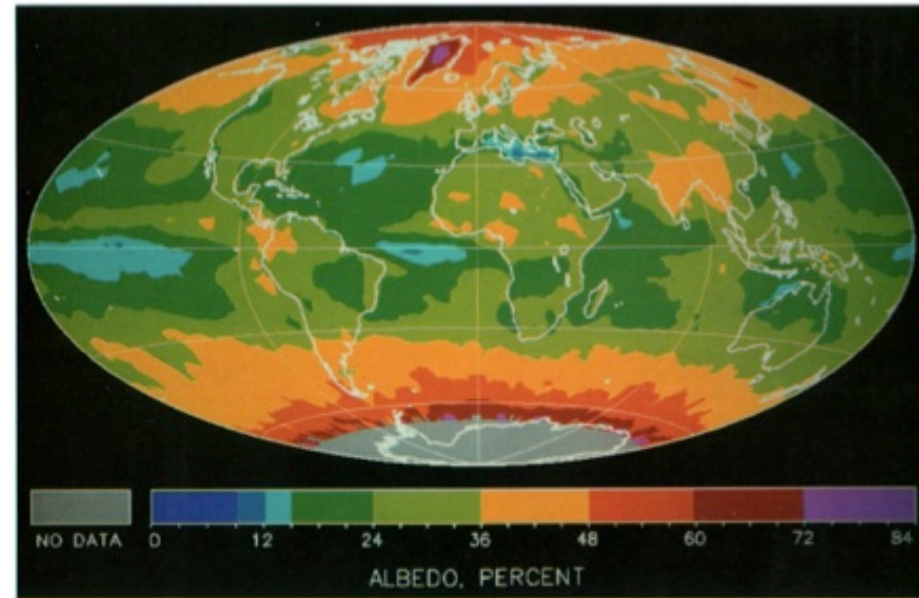
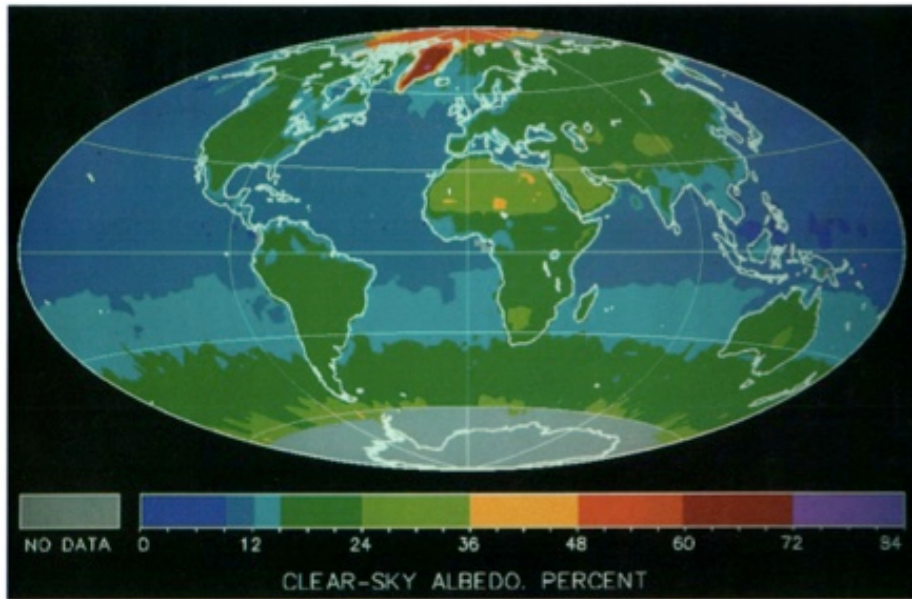


# Particles are very complicated

- Classical electromagnetism
  - Rayleigh scattering - molecules
  - Mie scattering – aerosol and droplets



# Effects on SW radiation

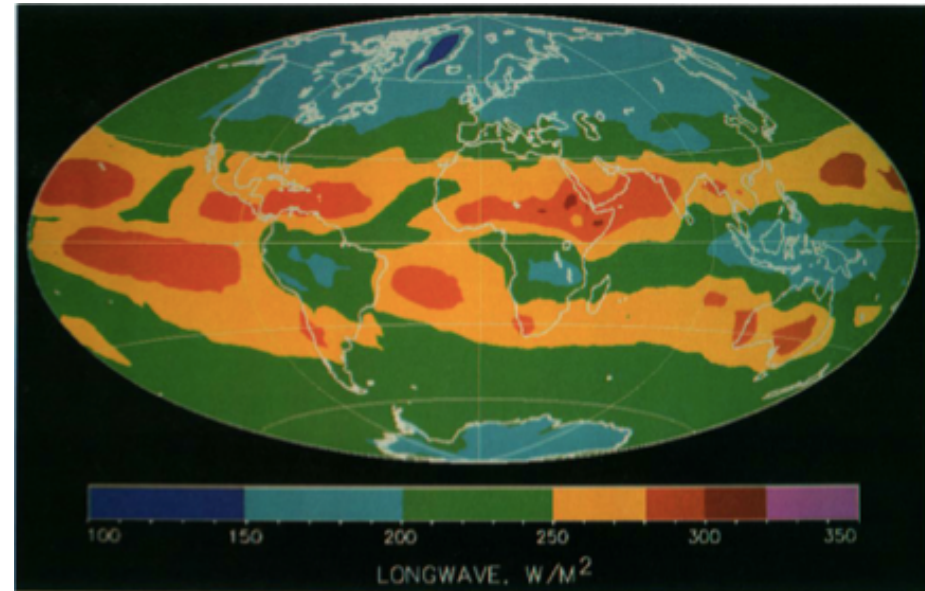
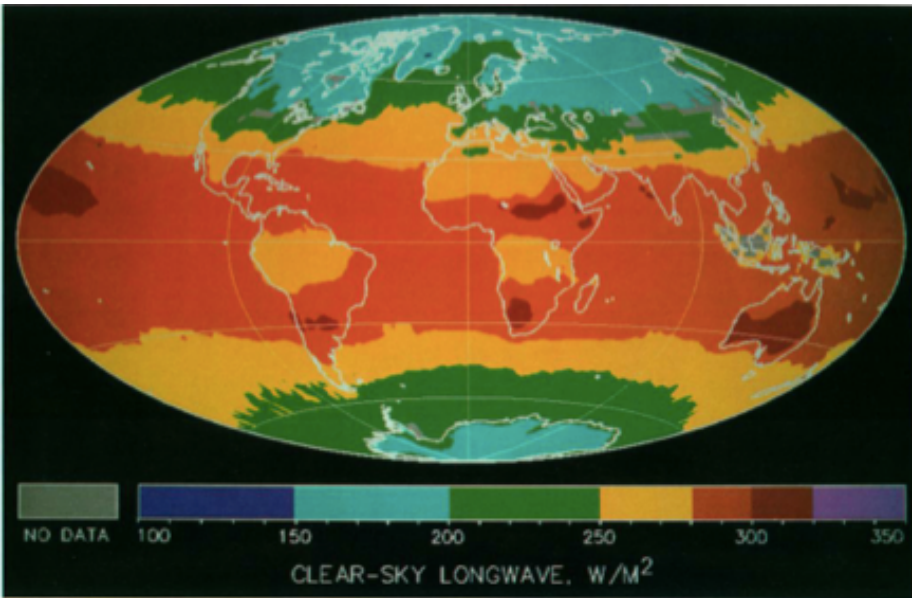


Harrison et al, JGR 1990

TABLE 1. Summary of Cloud Radiative Forcing Parameters ( $W/m^2$ )

Date	Longwave	Clear-Sky Longwave	Longwave Cloud Forcing	Shortwave Absorbed	Clear-Sky Shortwave Absorbed	Shortwave Cloud Forcing	Net Cloud Forcing
April 1985	234.5	265.8	31.3	236.5	281.6	-45.1	-13.8
July 1985	237.5	267.6	30.1	234.4	281.1	-46.7	-16.6
Oct. 1985	234.1	266.3	32.2	243.0	293.1	-50.1	-17.9
Jan. 1986	231.9	262.5	30.6	243.3	295.0	-51.7	-21.1
Annual	234.5	265.6	31.1	239.3	287.7	-48.4	-17.3

# Effects on LW radiation



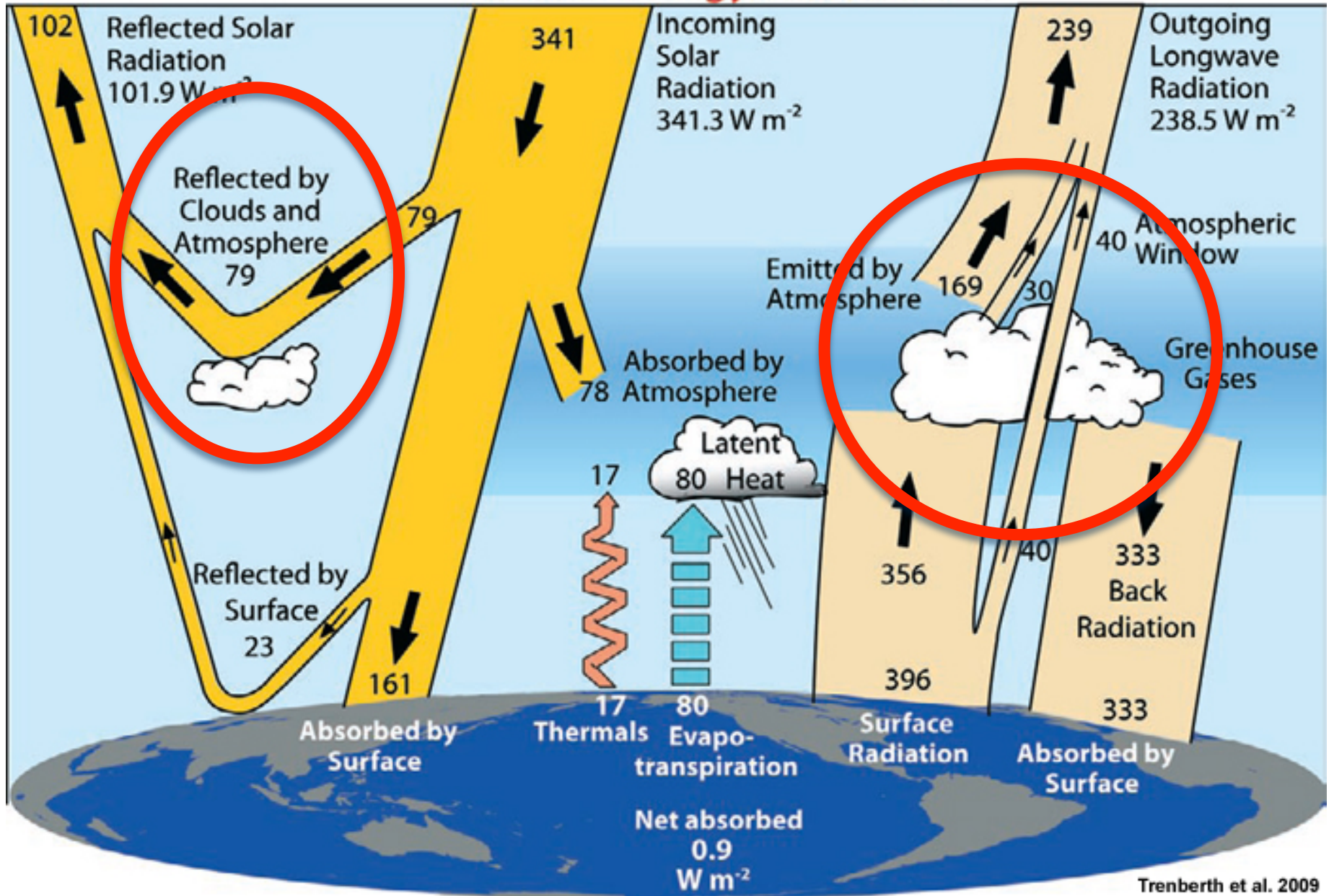
Harrison et al, JGR 1990

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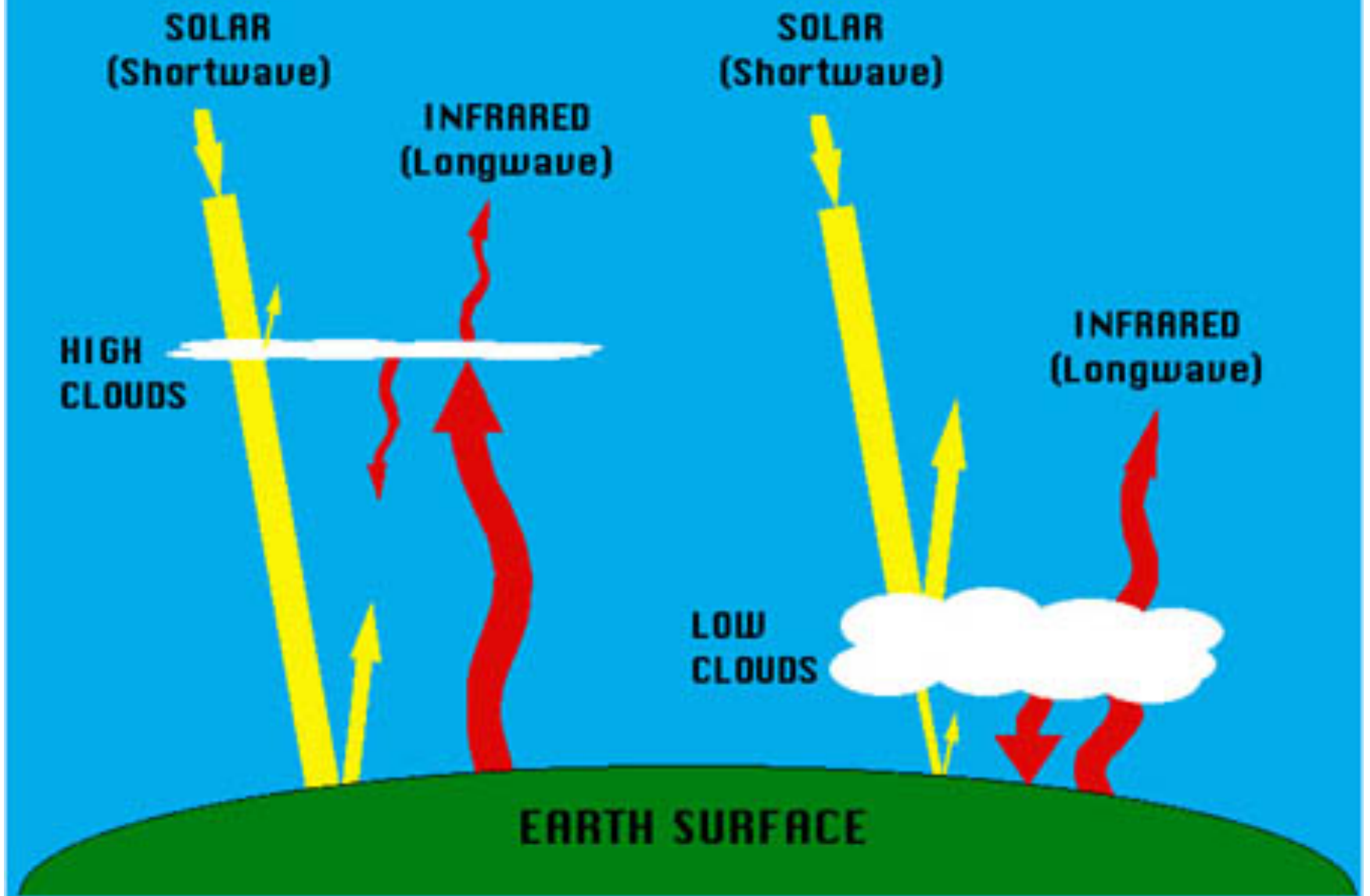
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# Global Energy Flows $W m^{-2}$

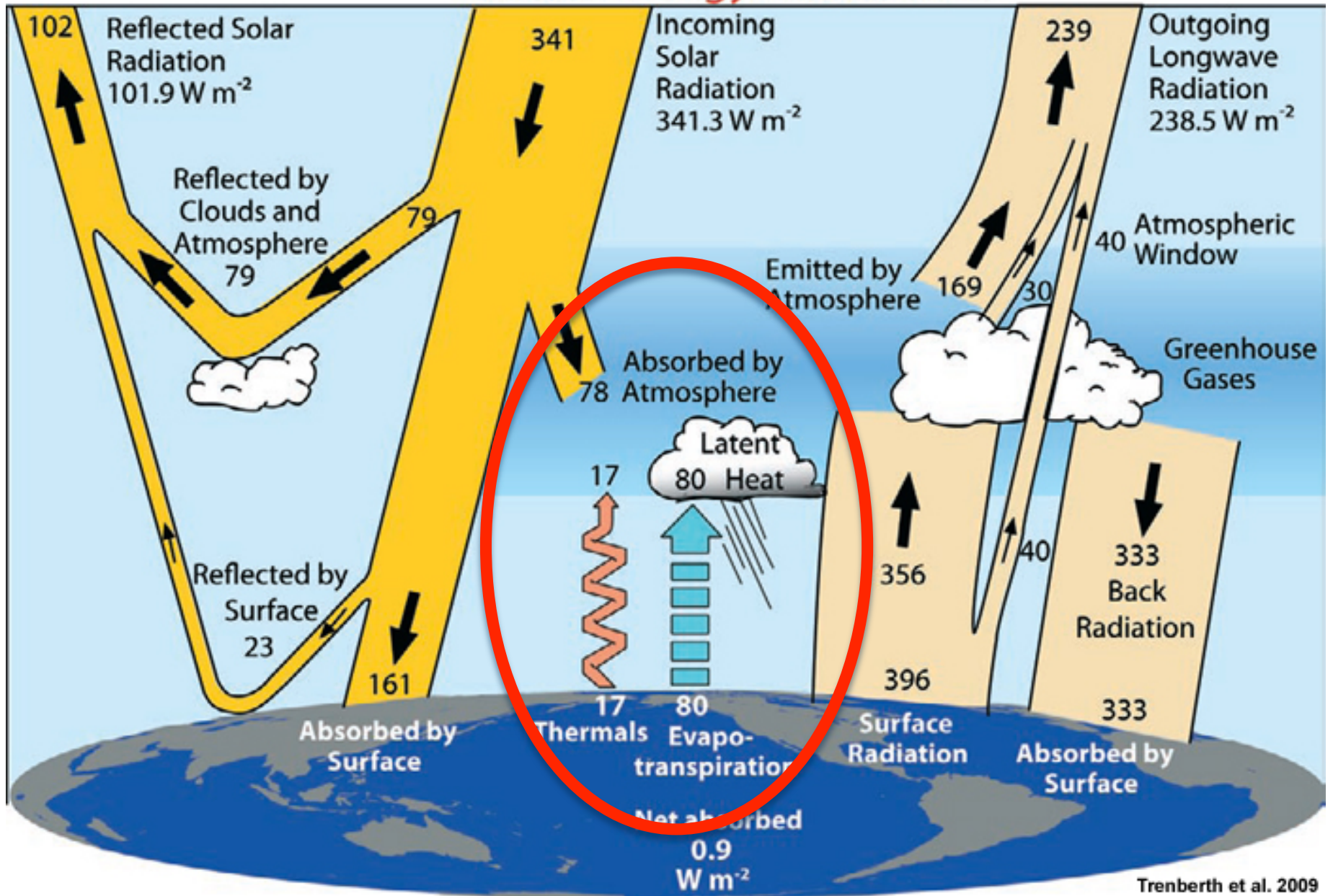


# CLOUD EFFECTS ON EARTH'S RADIATION

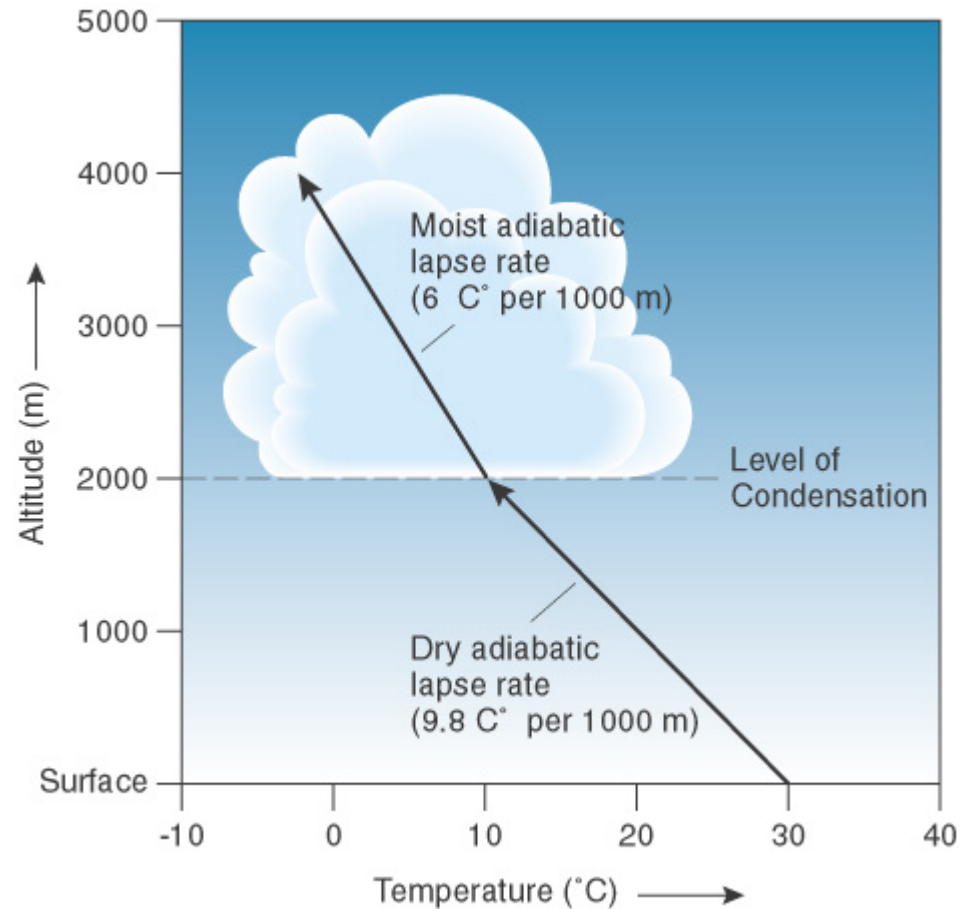
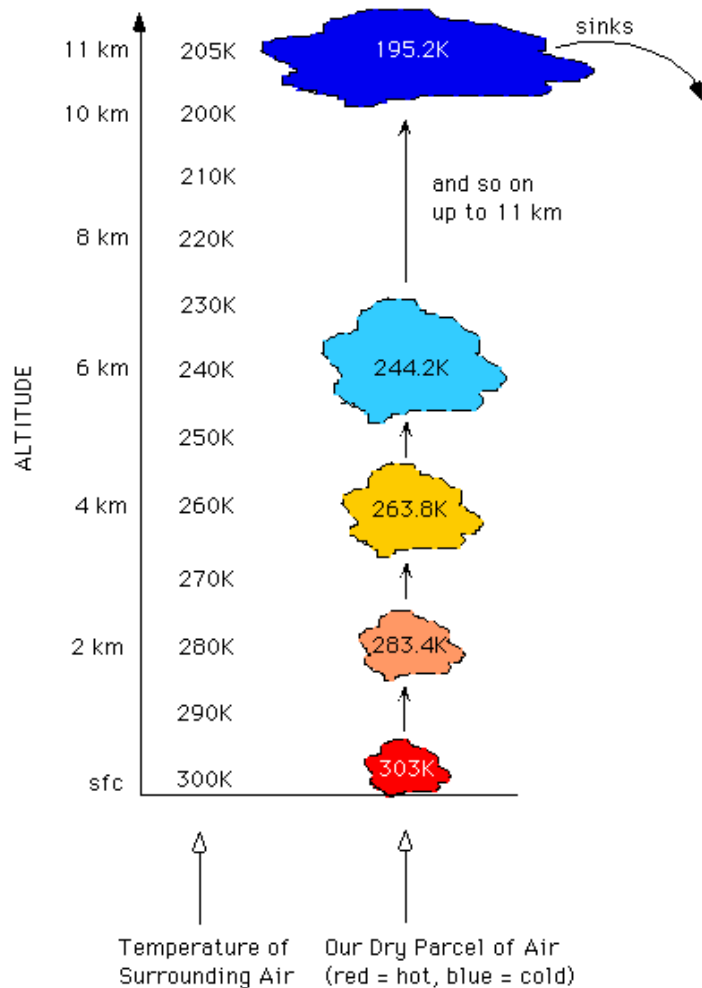




# Global Energy Flows $W m^{-2}$

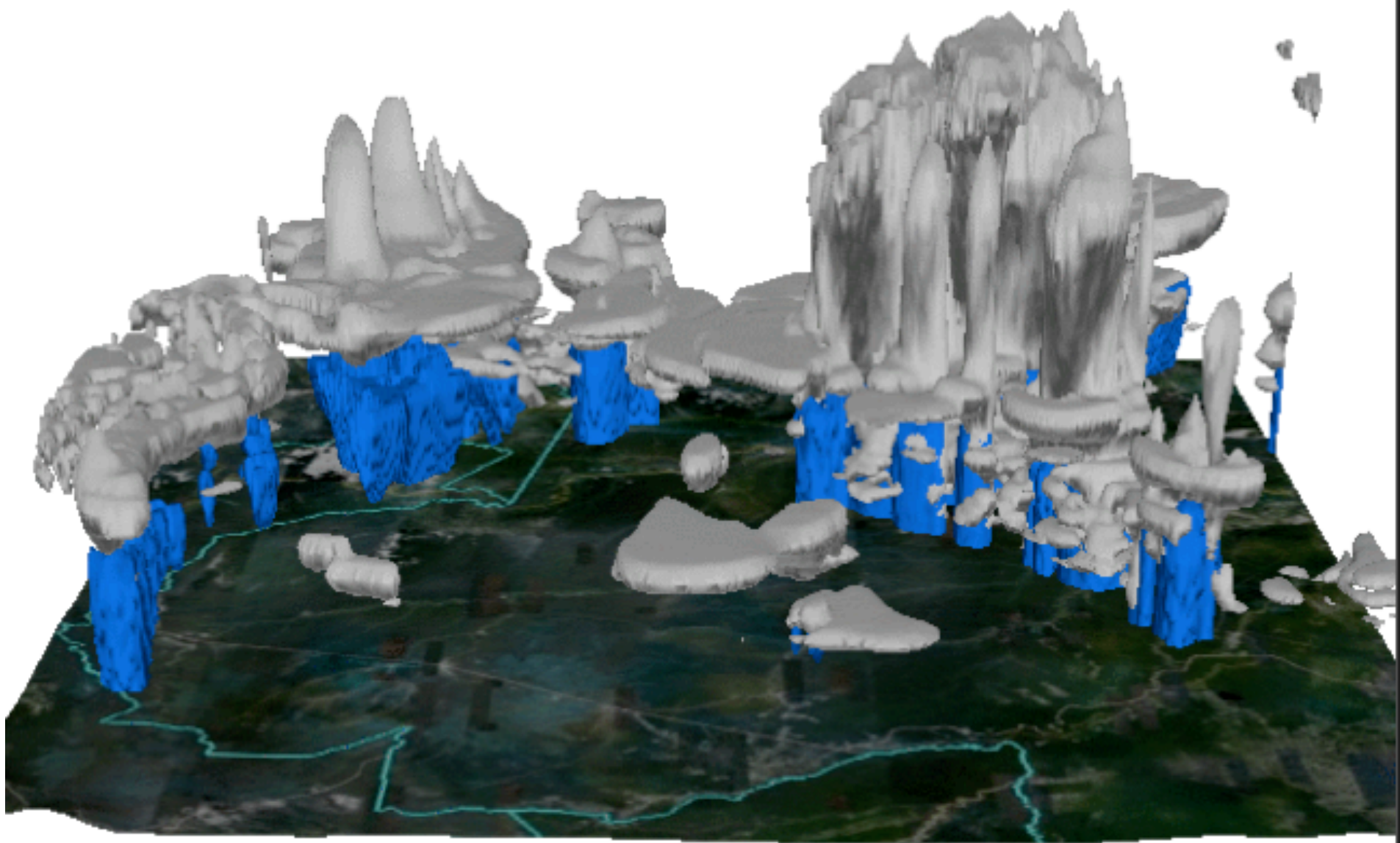


# Heat Exchange

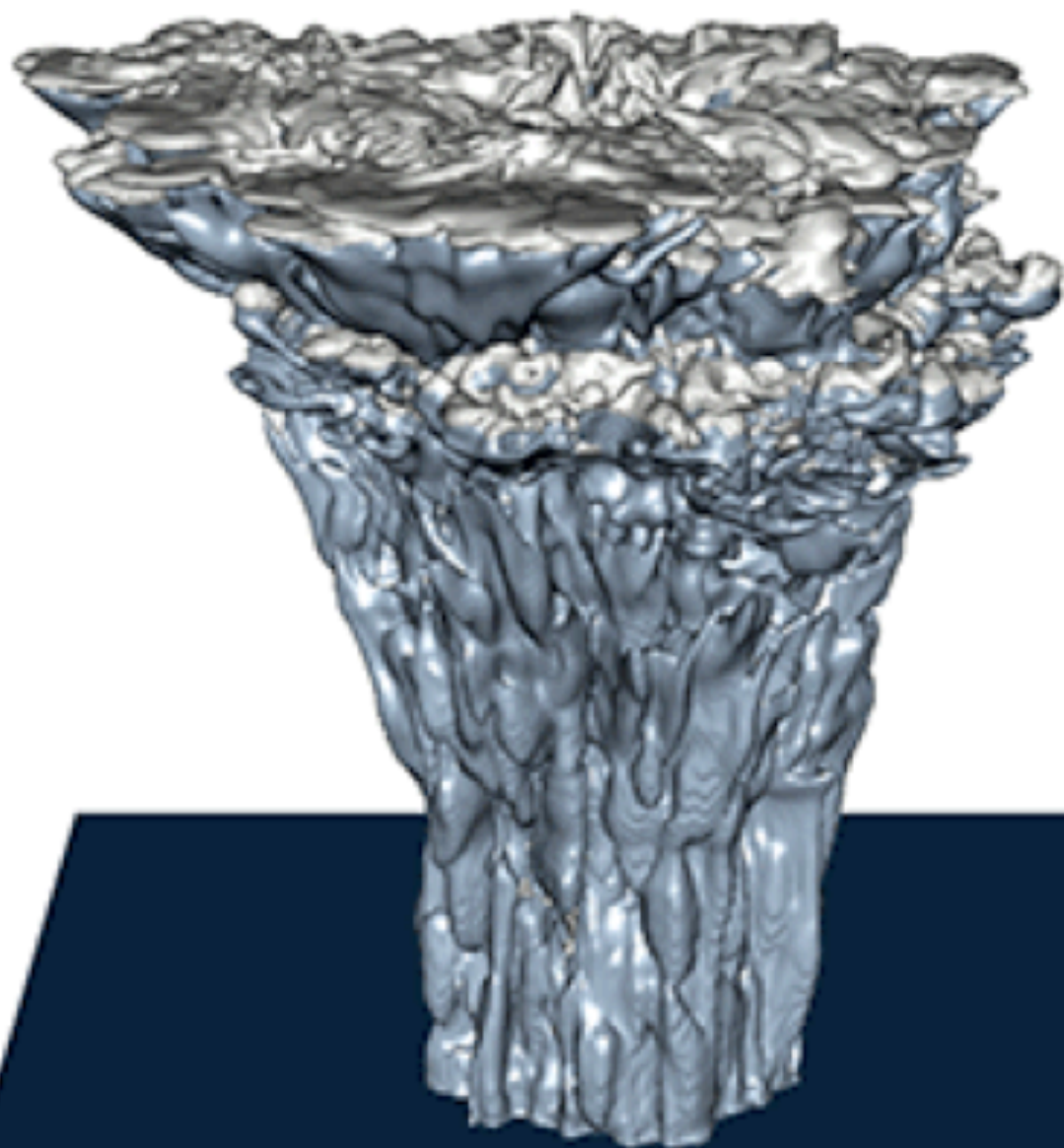


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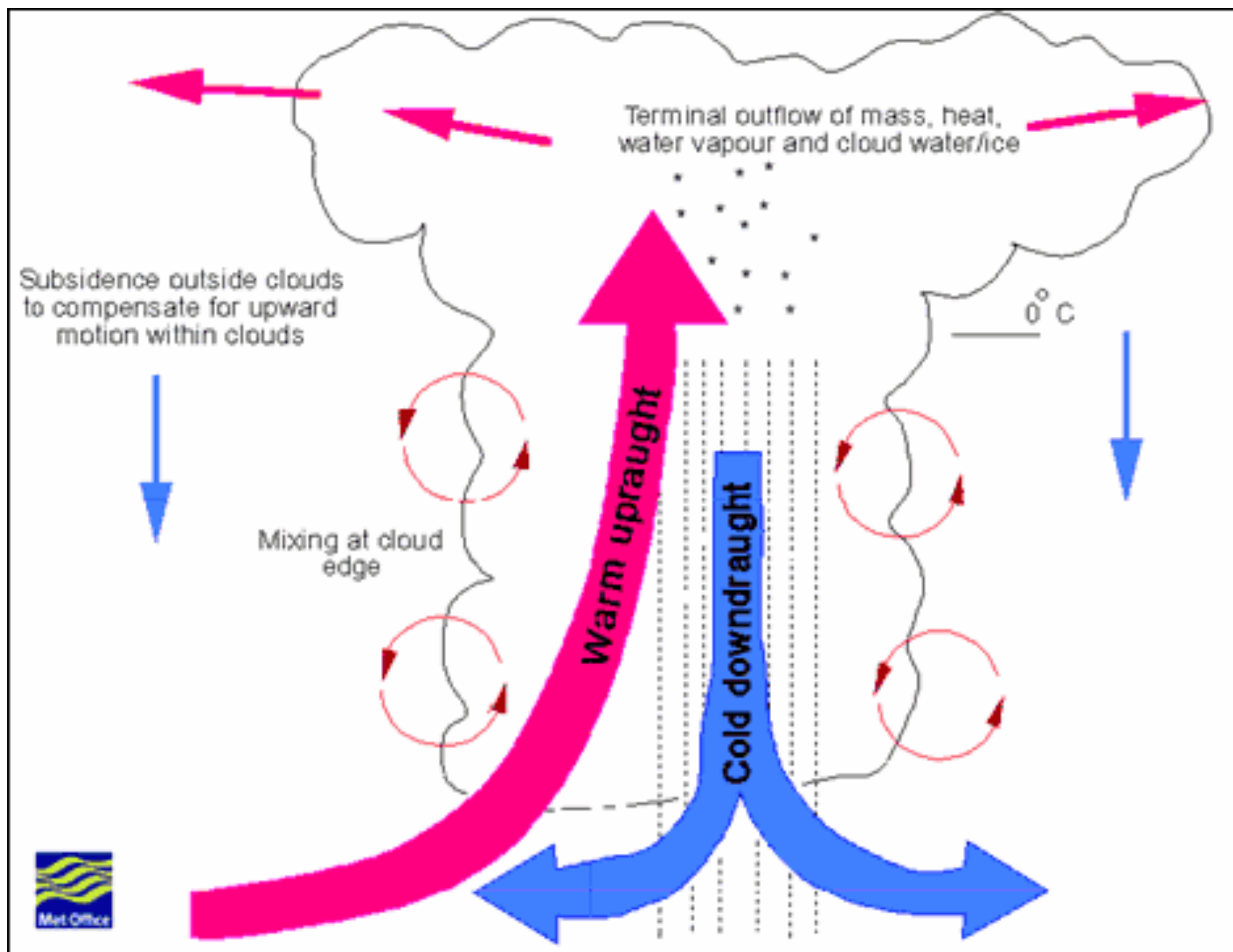
09:45:00  
26 Apr 2007  
8 of 16  
Thursday



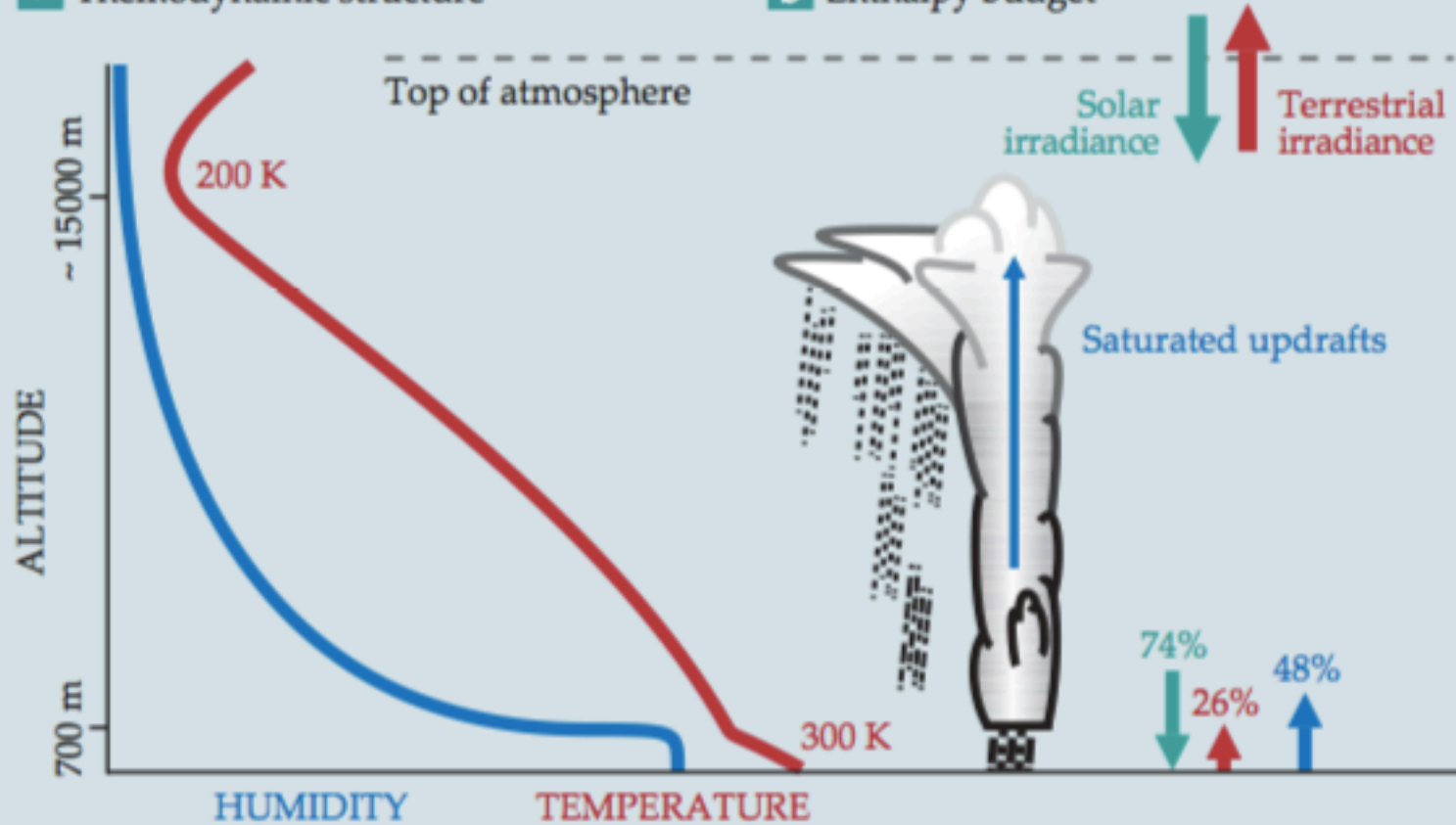




Lawrence Berkeley National Laboratory.

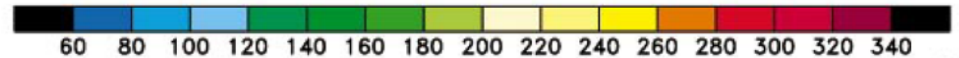
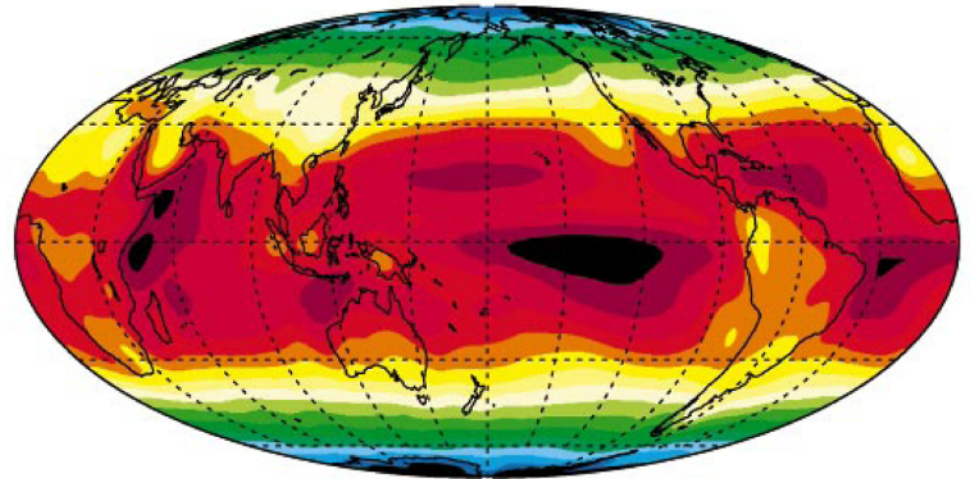




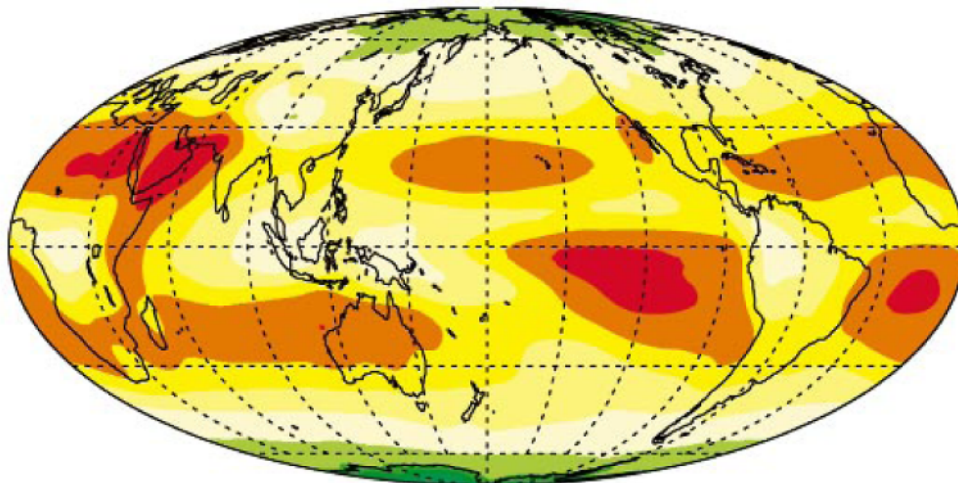
**a** Thermodynamic structure**b** Enthalpy budget

**Figure 3. The thermodynamic structure and enthalpy budget of the atmosphere.** (a) The atmosphere's temperature (red) and its absolute humidity (blue) are closely coupled. (b) At the top of the atmosphere solar and terrestrial irradiances balance one another. According to calculations, most (74%) of the incoming solar irradiance reaches the surface, but the net terrestrial irradiance at the surface is only a small fraction (26%) of its value at the top of the atmosphere. The radiative deficit (48%) is balanced by surface turbulent fluxes of enthalpy, arising mostly from evaporation, that transport warm water vapor from the surface to the troposphere, where it cools and condenses.

# Distribution on the Earth



Annual mean solar radiation budget at top (W/m<sup>2</sup>)



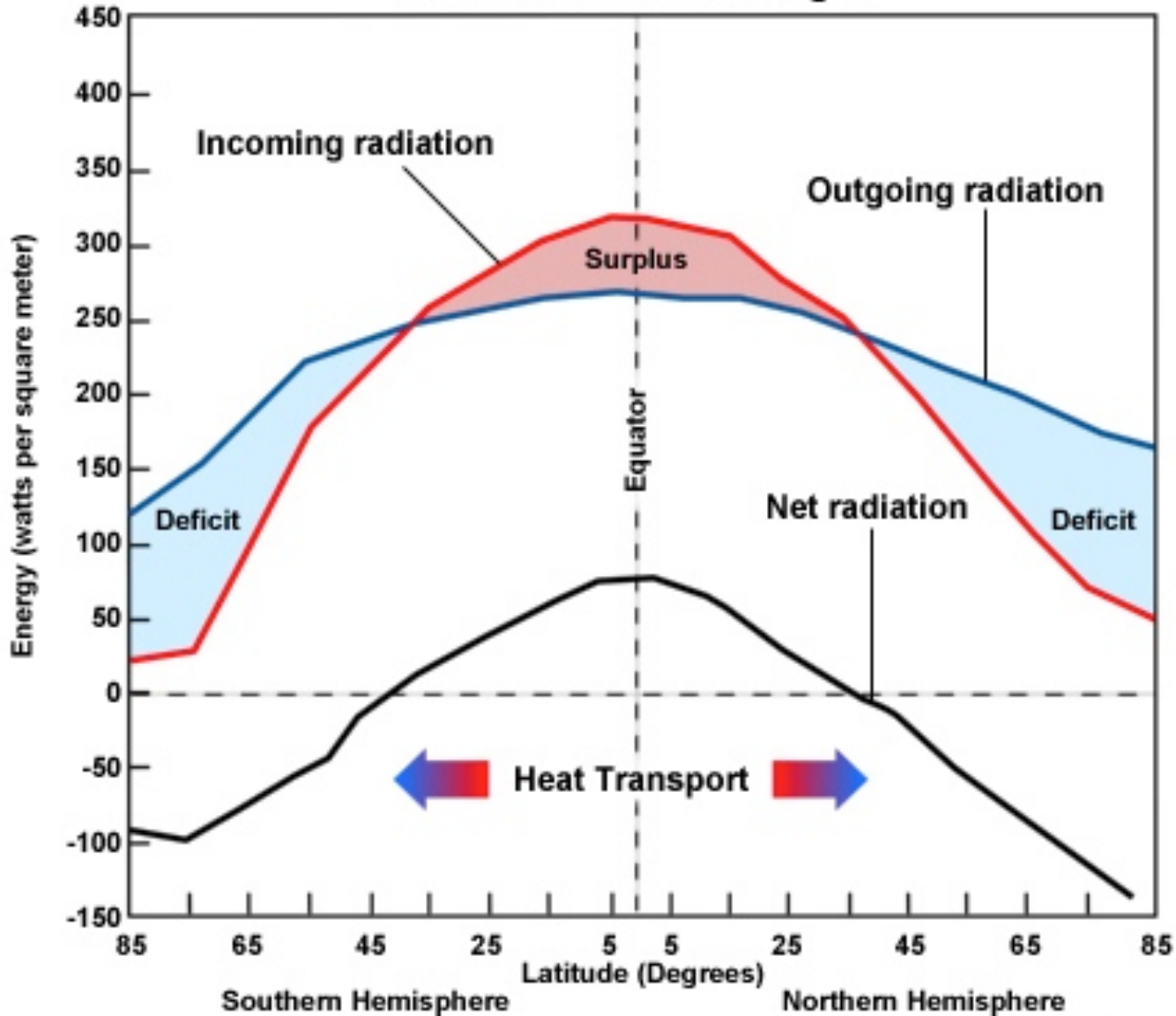
Annual mean outgoing long wave radiation at top (W/m<sup>2</sup>)



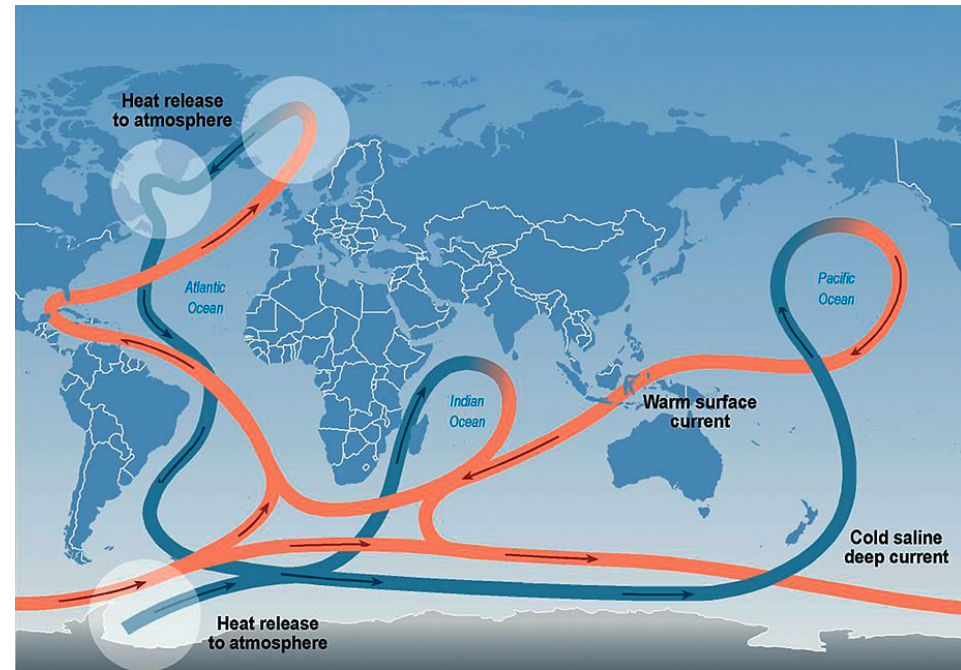
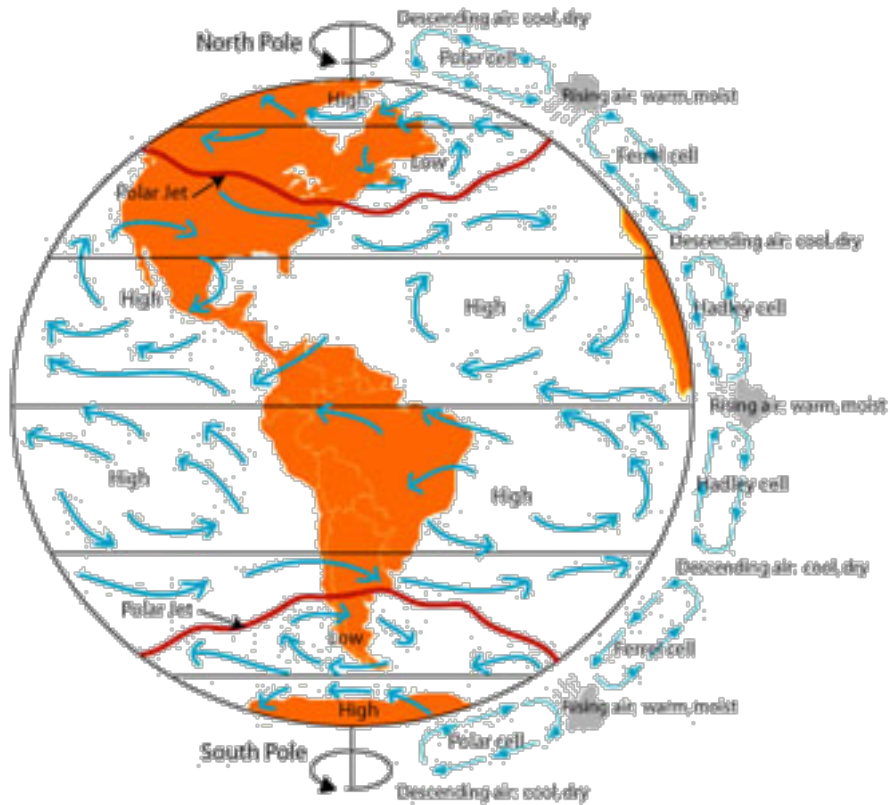
Trenberth and Stepaniak, J. Clim. (2003)

b

### Annual Radiation Budget

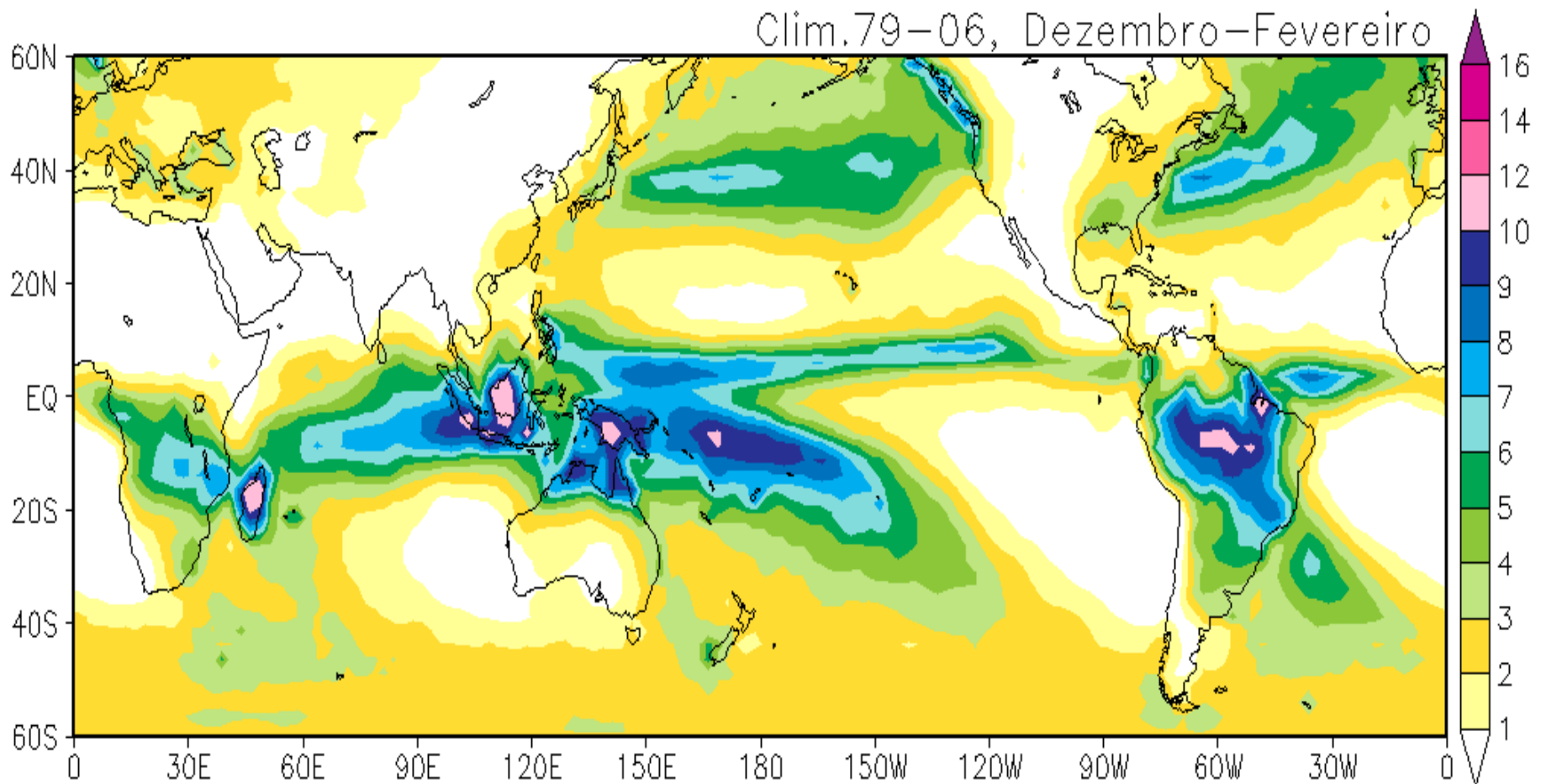


# Atmospheric and Oceanic circulation



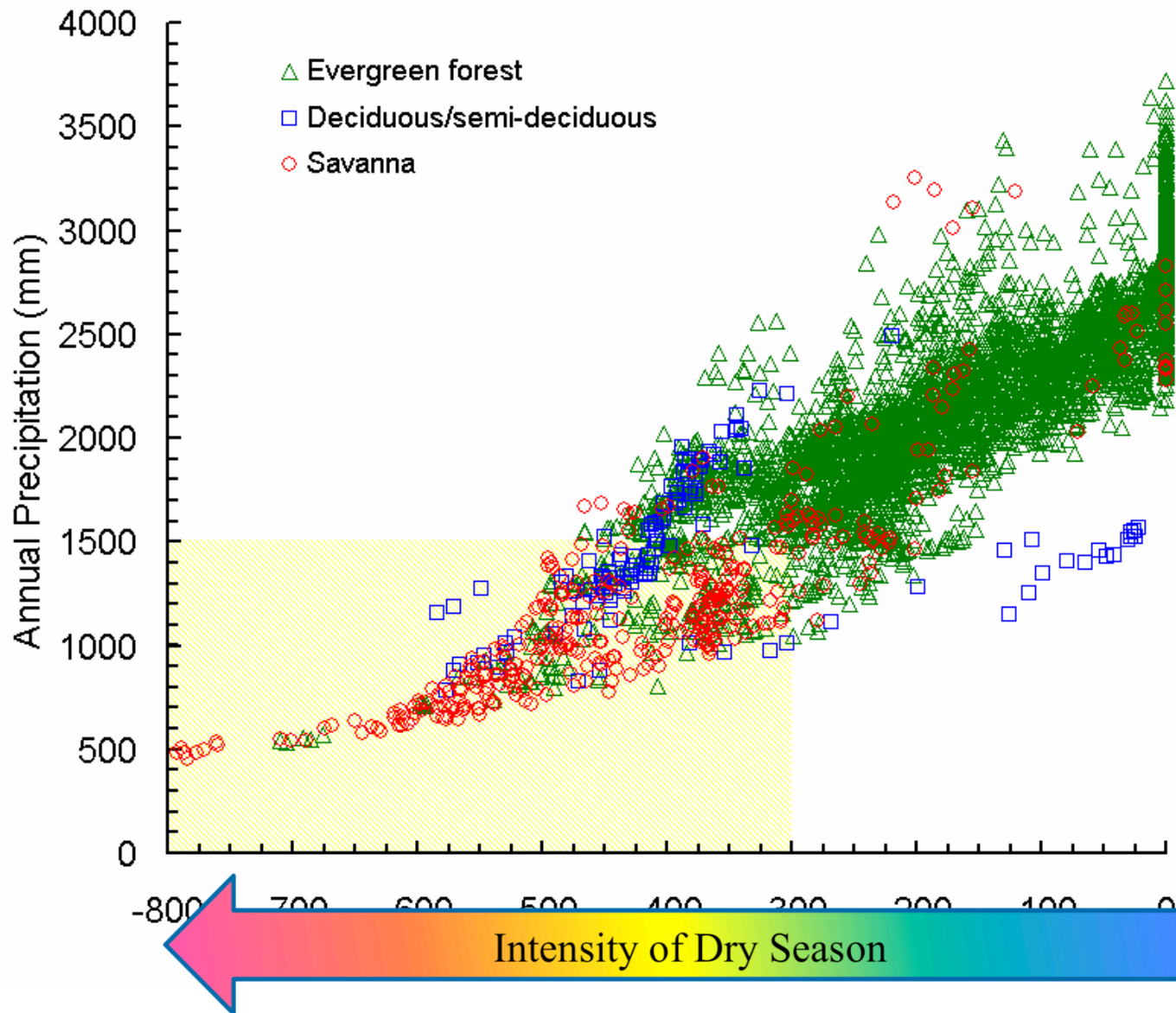


# Global Precipitation





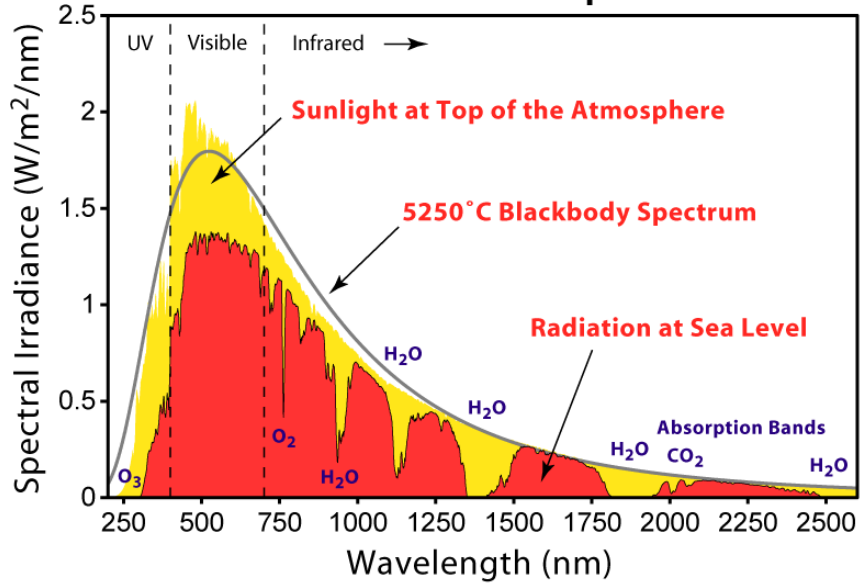
# A Rainfall Biogeography of Amazonia



Source: Malhi *et al.*, **Exploring the likelihood and mechanism of a climate-change induced dieback of the Amazon rainforest**, *Proceedings of the National Academy of Sciences*, 2010



## Solar Radiation Spectrum



**Thanks!**

[hbarbosa@if.usp.br](mailto:hbarbosa@if.usp.br)

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