

A satellite image of Earth showing the continent of South America and the surrounding Atlantic Ocean. The image is partially obscured by text on the right side.

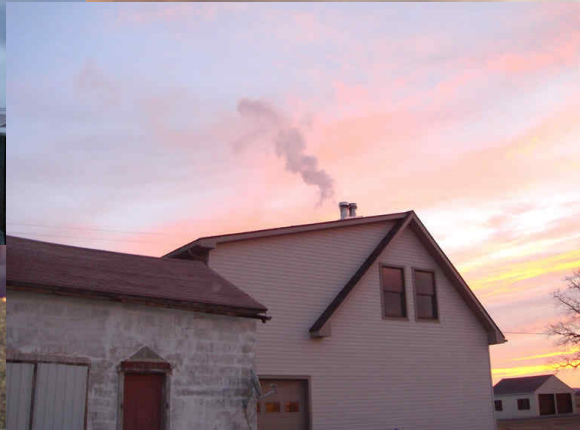
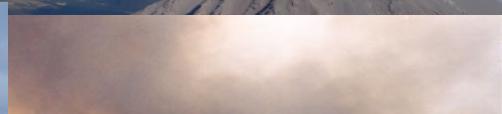
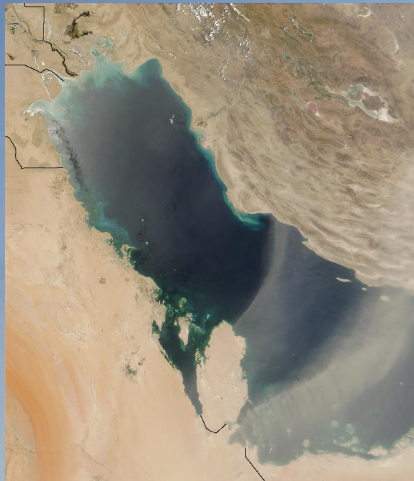
Aerossóis, aula 3

Física Atmosférica, 2019

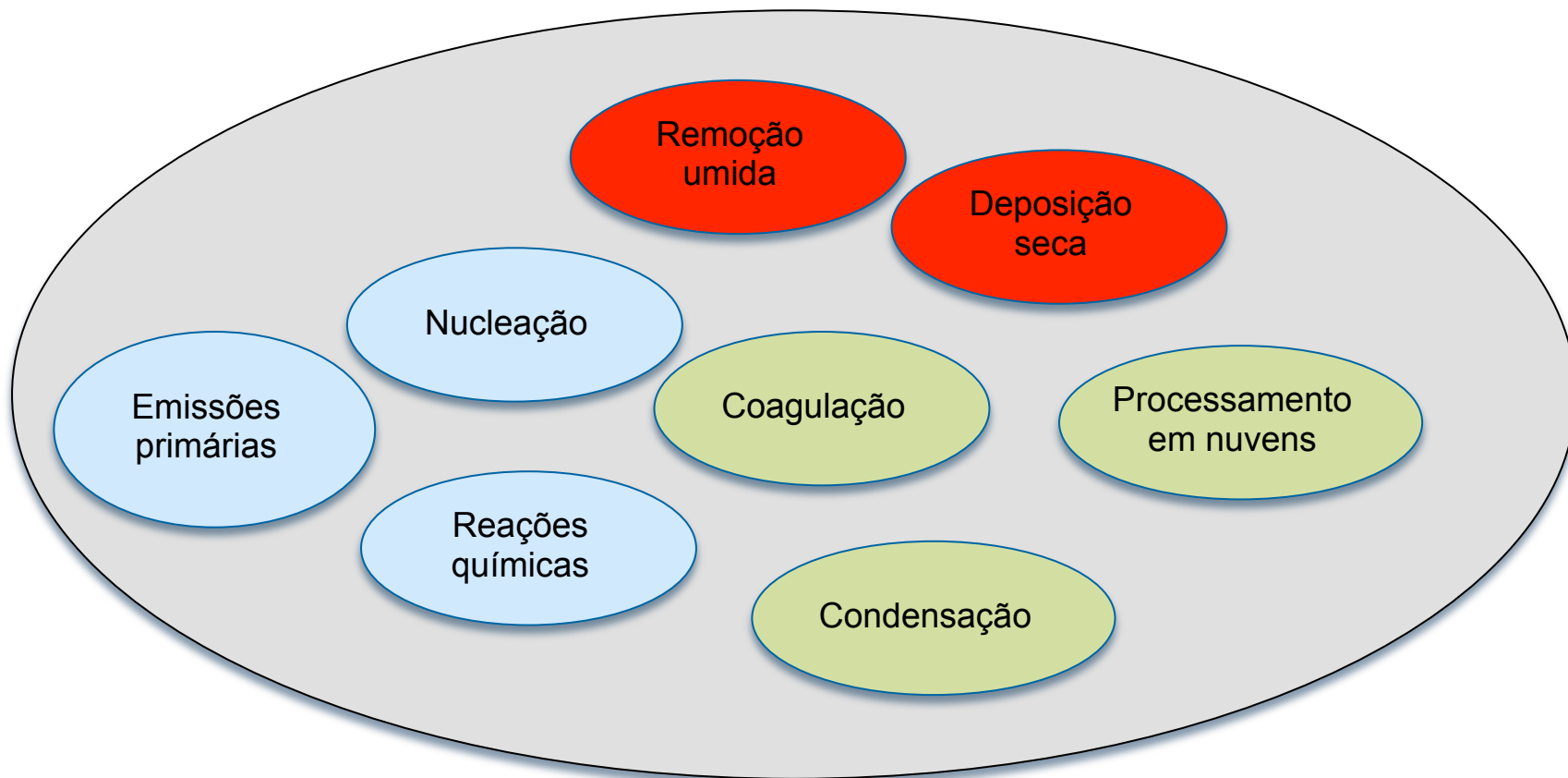
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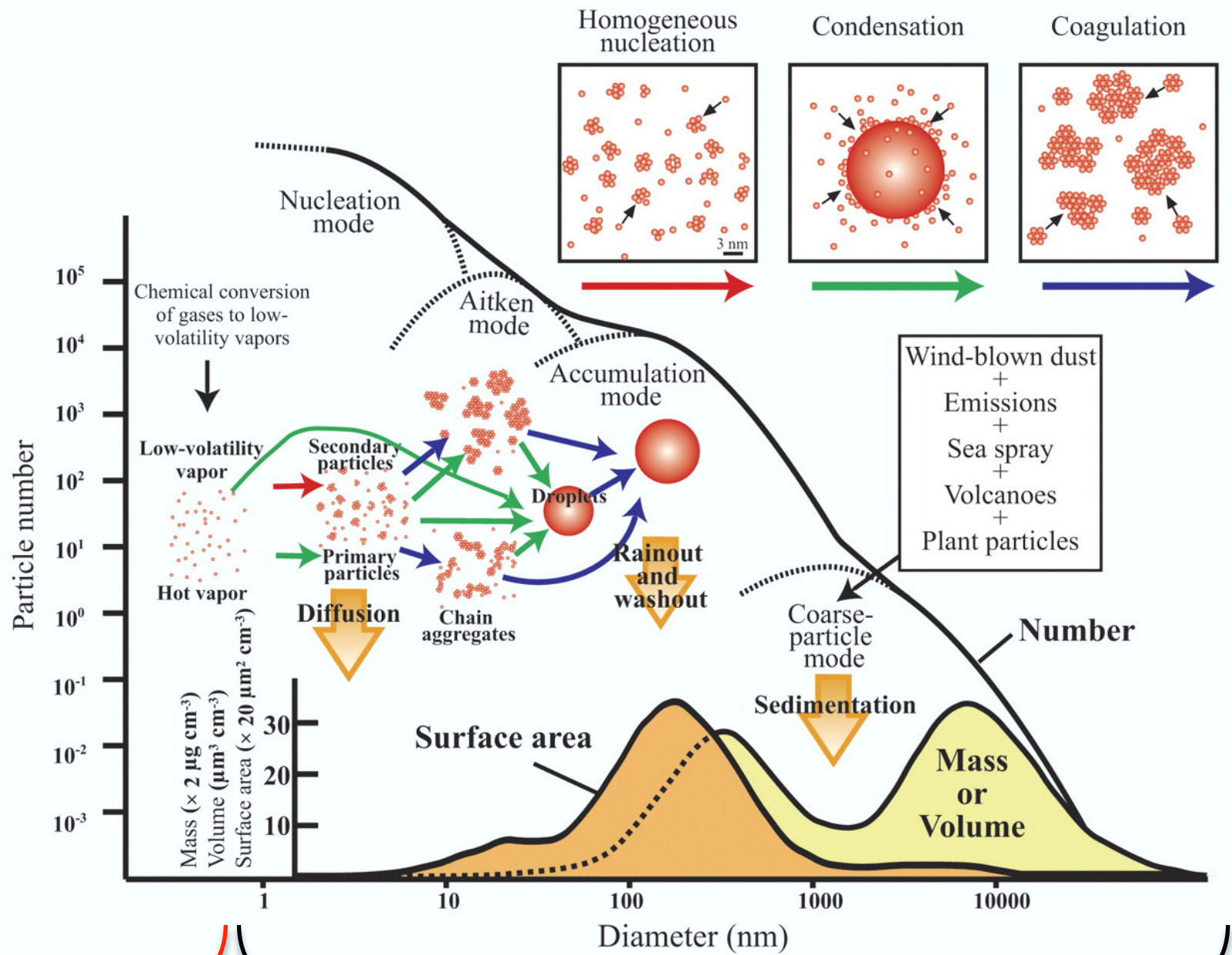
Com slides de HC Hanssen, Peter Tunved, Andre Burger, Marco Franco, etc.

- Sea salt
- Dust
- Biomass burning
- Biogenic
- Volcanic
- Urban



Processos envolvendo aerossóis





gases

aerosols

Basic processes acting on single aerosol particles

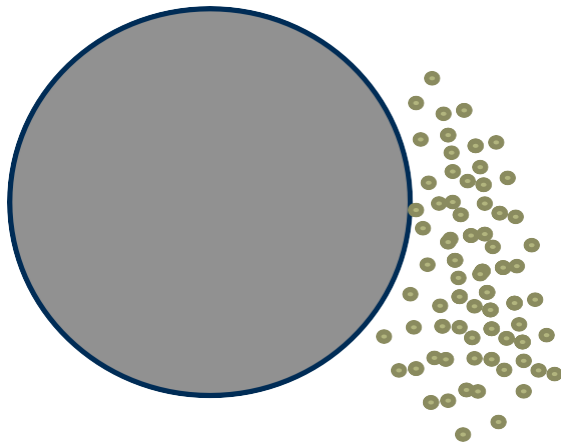
- Gravitational settling
- Drag force
- Brownian motion

Single particle dynamics and Knudsen number

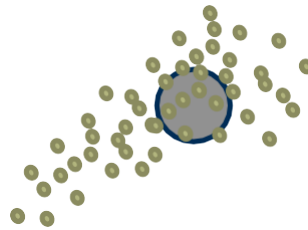
Continuum regime

transition regime

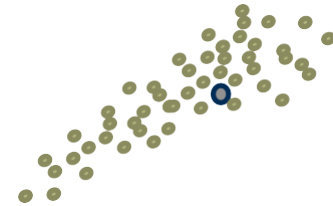
Free molecular regime



$$Kn \rightarrow 0$$



$$Kn \approx 1$$



$$Kn \rightarrow \infty$$

$$K_n = \frac{2\lambda}{D_p}$$

Knudsen number

Where λ is mean free path of air

D_p is diameter of particle

Single particle dynamics and Knudsen number

$$\lambda_a = \frac{2\nu_a}{\bar{v}_a}$$

ν_a Kinematic viscosity of air (m²/s)
 \bar{v}_a Average speed of an air molecule (m/s)

= 65nm

T = 288K

P_a = 1013hPa

$$\bar{v}_a = \sqrt{\frac{2N_a k_B T}{\pi m_a}}$$

T is temperature
 m_a is molar mass

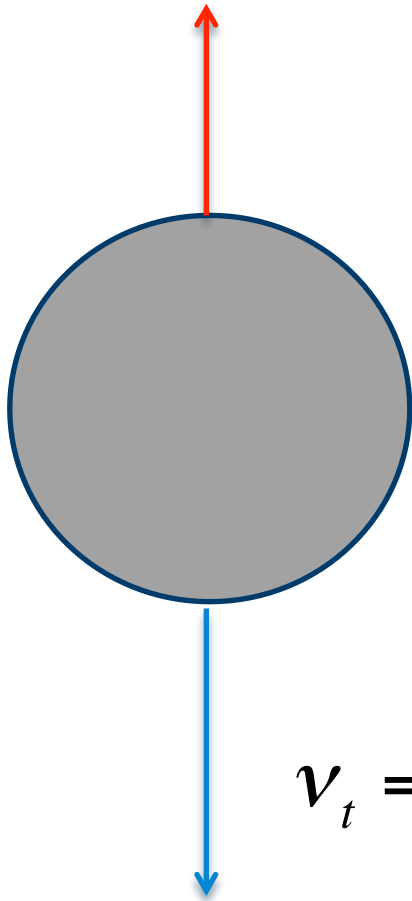
$$\nu_a = \frac{\eta_a}{\rho_a}$$

$$\eta_a = \frac{5}{16A d_a^2} \sqrt{\frac{m_a R^* T}{\pi}}$$

d_a is diameter of air molecule

D _P (μm)	Kn
0.01	12.9
0.1	1.29
1	0.129
10	0.0129

Gravitational settling



$$F_{drag} = F_{grav}$$

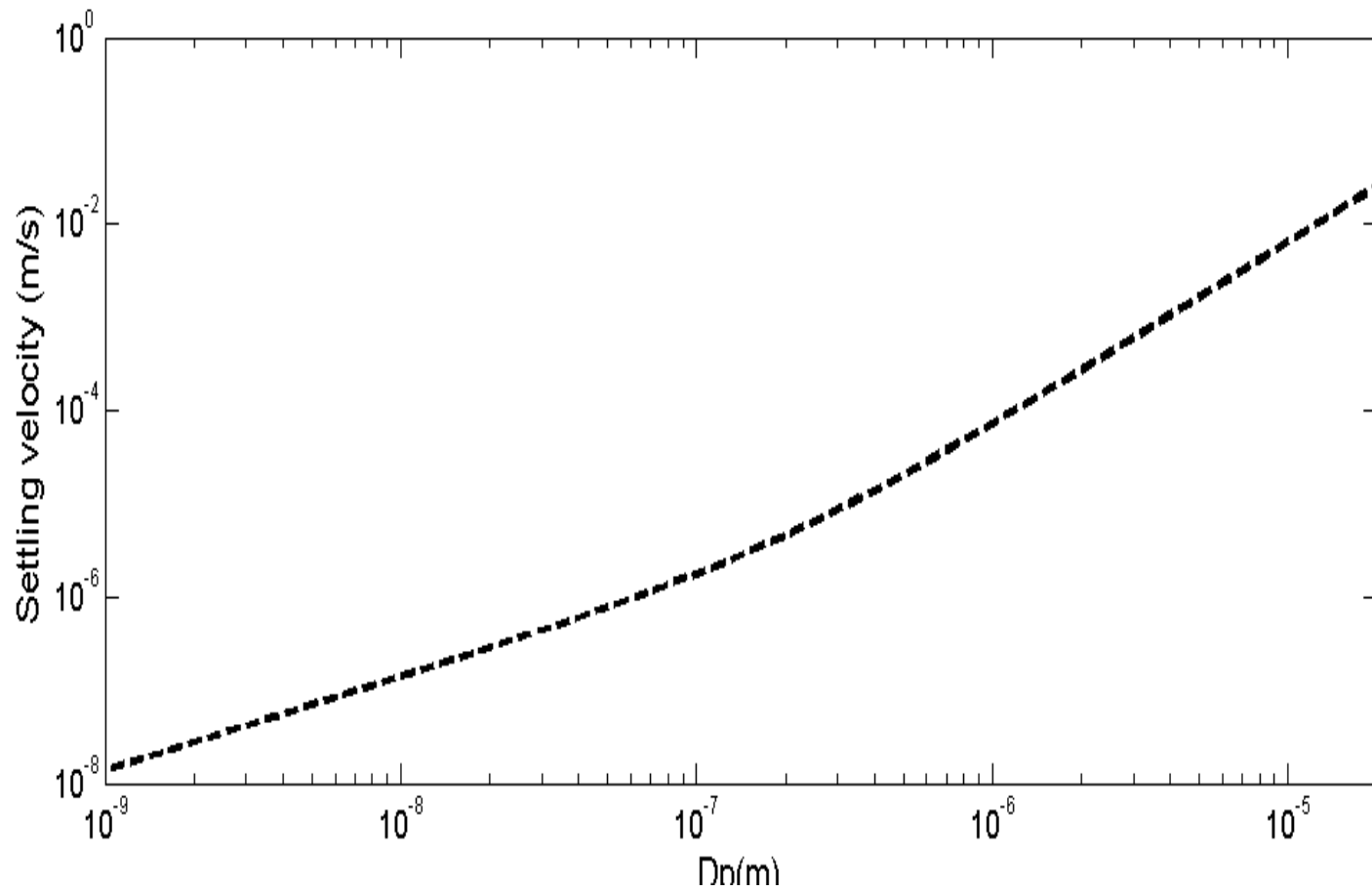
$$F_{grav} = mg = \frac{\pi}{6} \rho_P D_P^3 g$$

$$F_{drag} = \frac{3\pi D_P \eta_a v_t}{C_c}, C_c = 1 + K_n$$

$$v_t = \frac{m_P C_c g}{3\pi \eta_a D_P}$$

D_P (μm)	V_t ($\mu\text{m/s}$)
0.01	0.062
0.1	1
1	51
10	4576

Terminal velocity

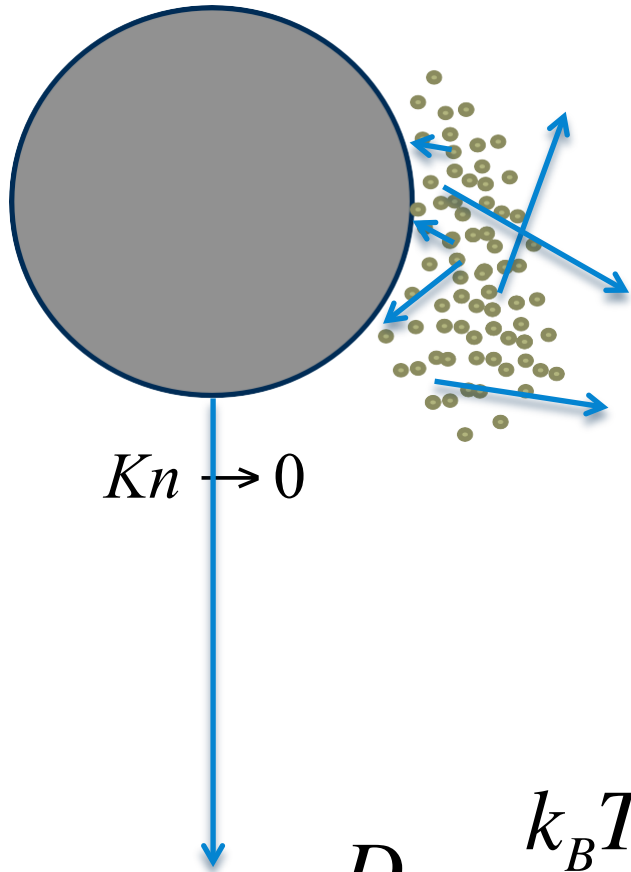


Brownian diffusion

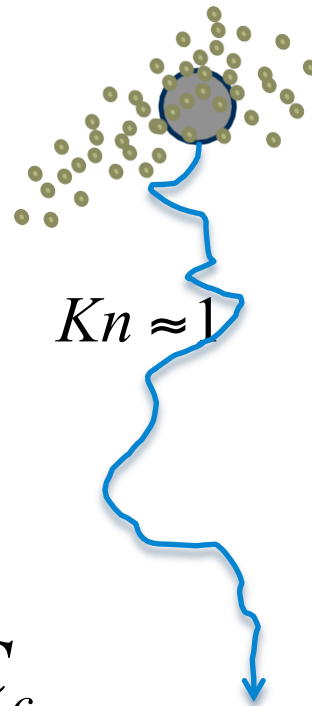
Continuum regime

transition regime

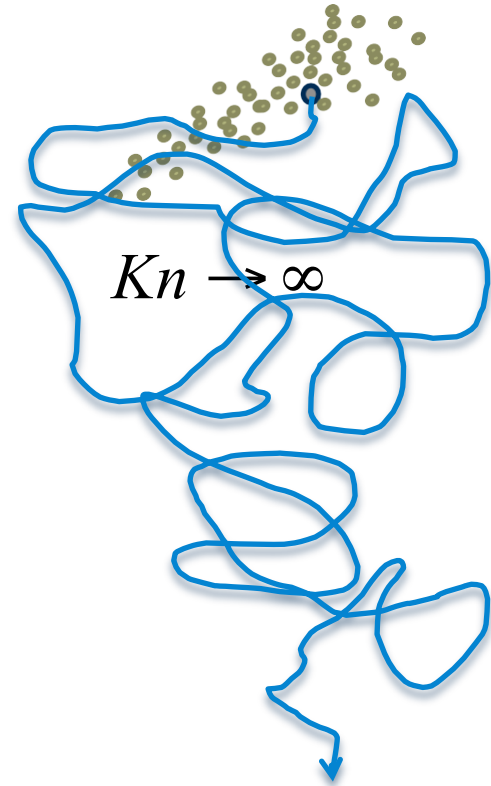
Free molecular regime



$$Kn \rightarrow 0$$



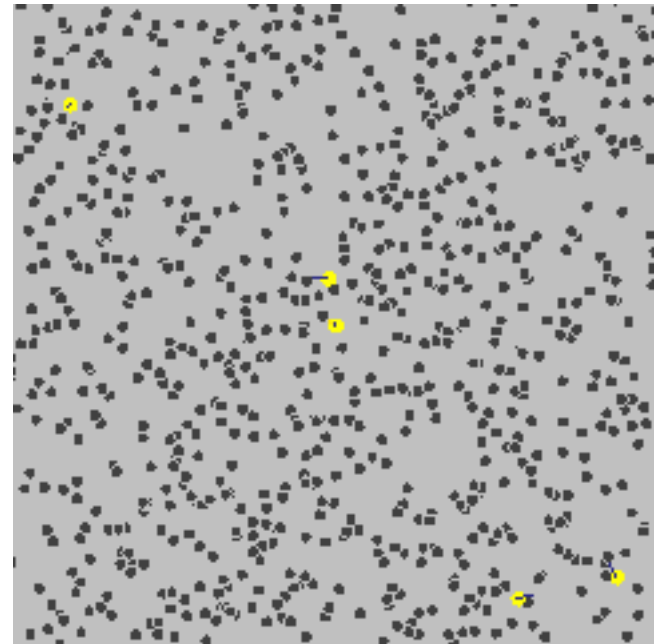
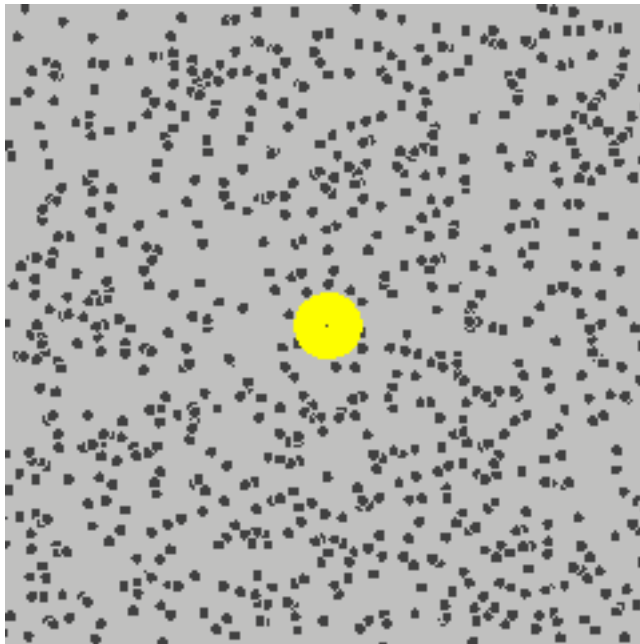
$$Kn \approx 1$$



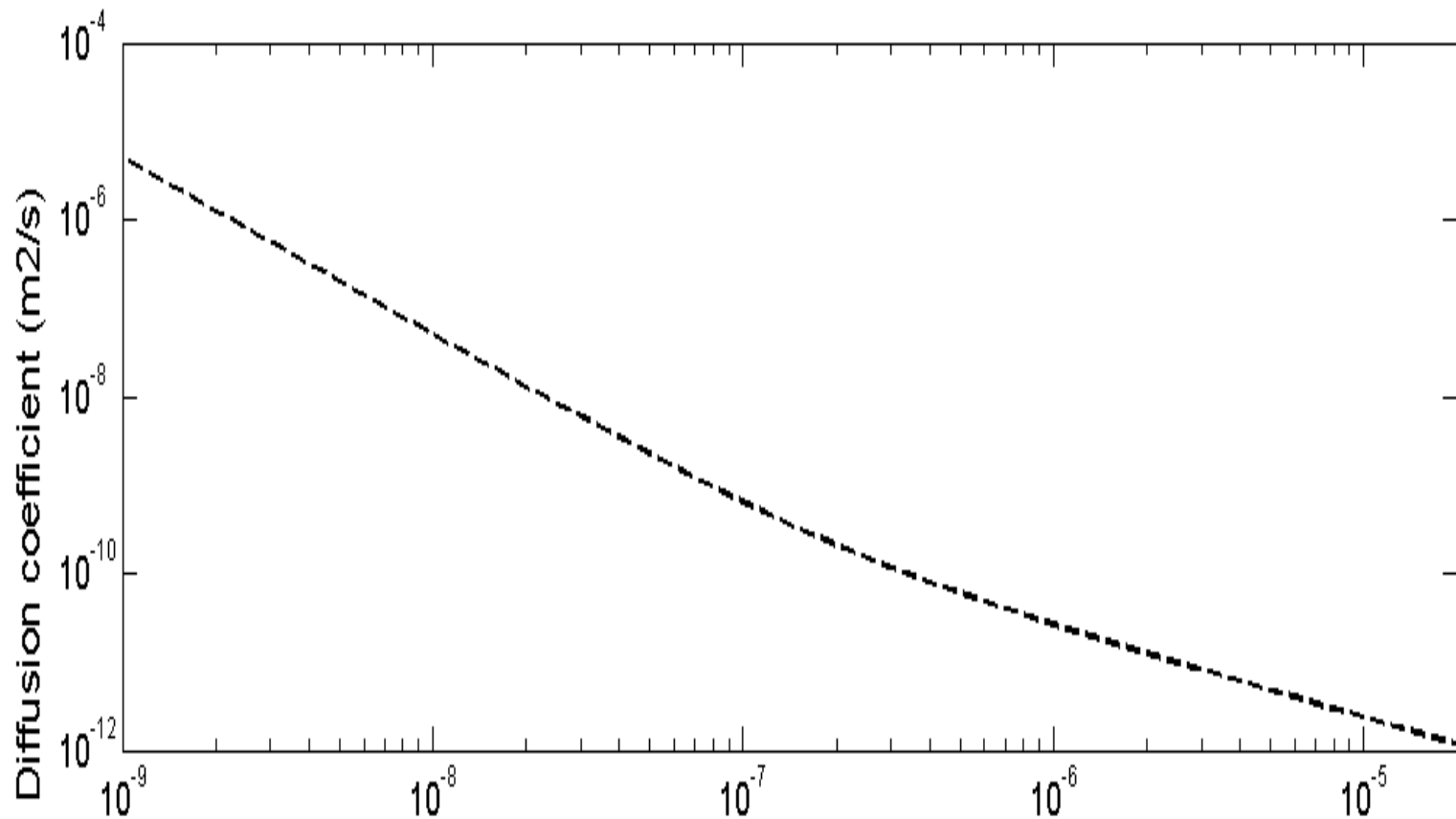
$$Kn \rightarrow \infty$$

$$D = \frac{k_B T C_c}{3\pi\eta_a D_P}$$

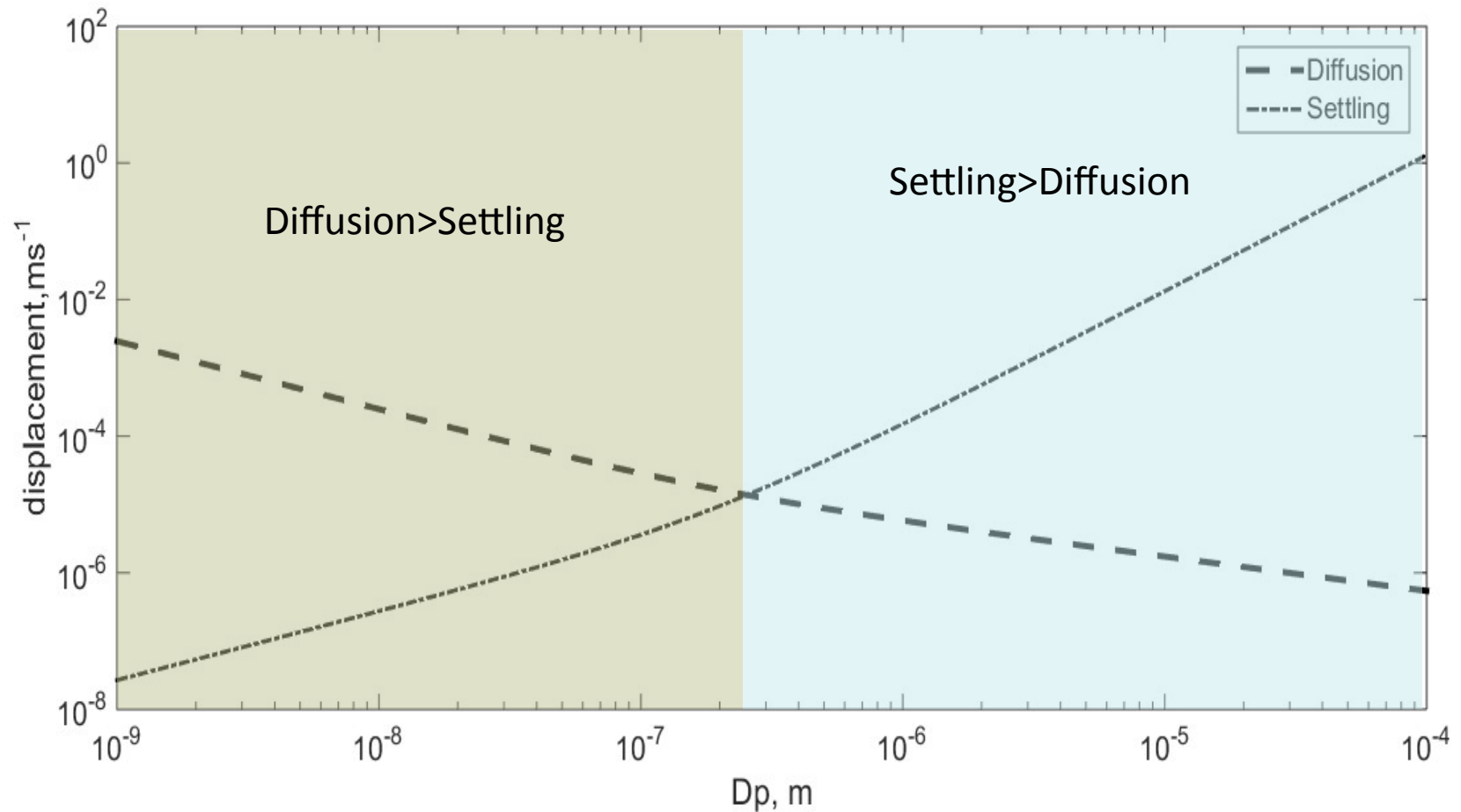
Brownian diffusion



Diffusion coefficient



Estimating displacement as function of size; Brownian diffusion vs gravitational settling



Coagulação

- Resultado, principalmente, de movimento Browniano, mas outras forças também tem um papel (elétrica, gravitacional, etc)
- Duas partículas colidem, agregam e formam uma nova partícula
- Coagulação não afeta a massa, mas reduz o número
- É mais eficiente para partículas pequenas

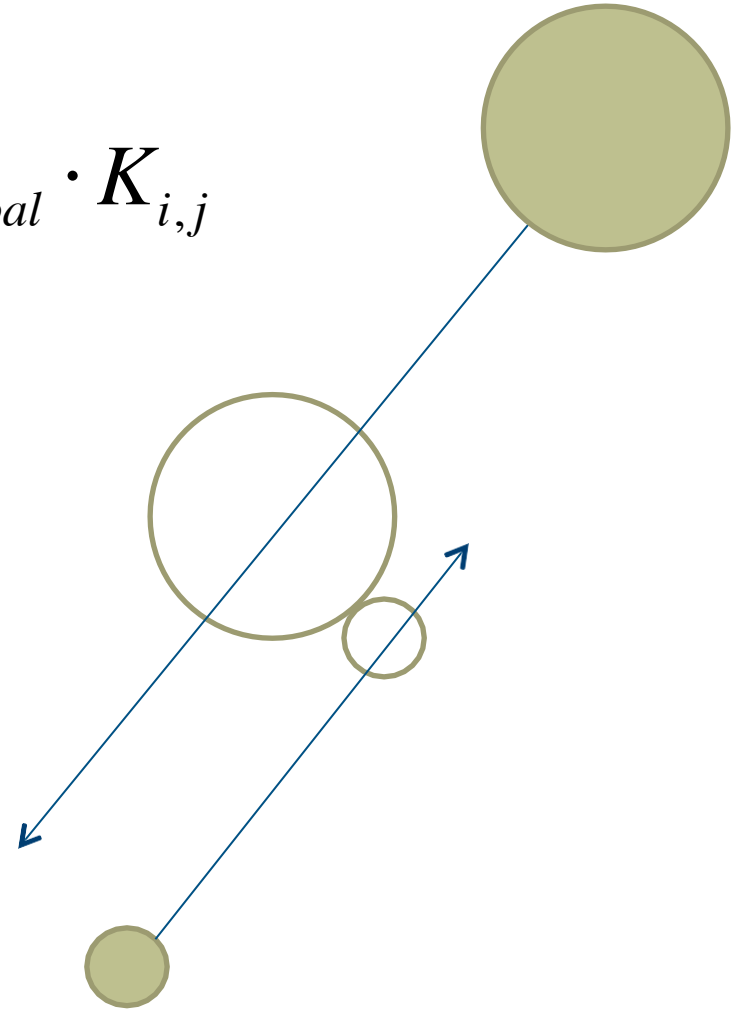
Coagulation

Free molecular regime

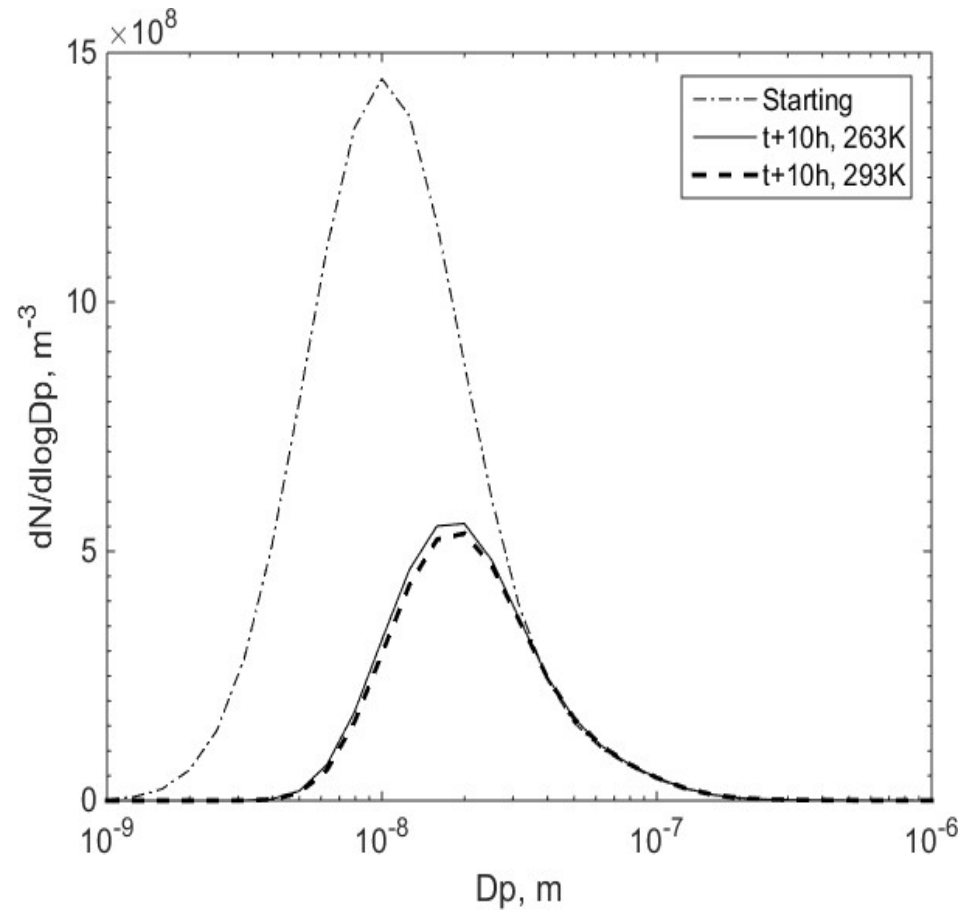
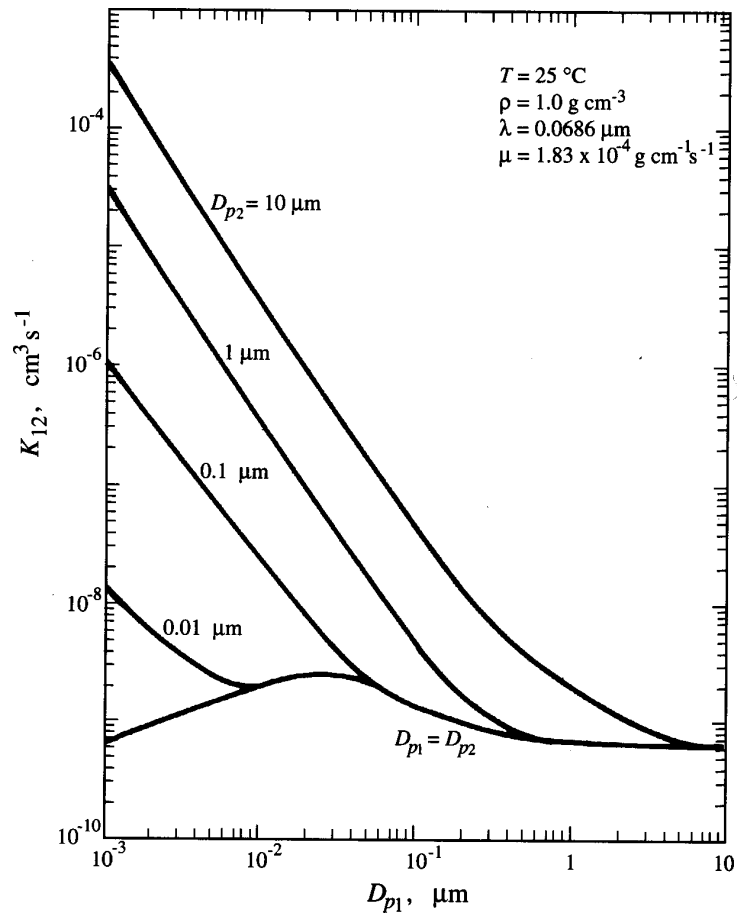
$$\text{Coagulation rate coefficient} = E_{coal} \cdot K_{i,j}$$

When two aerosol particles collide, they may or may not stick together, depending on the efficiency of coalescence. The efficiency of coalescence depends on particle shape, composition, ambient relative humidity, and other factors.

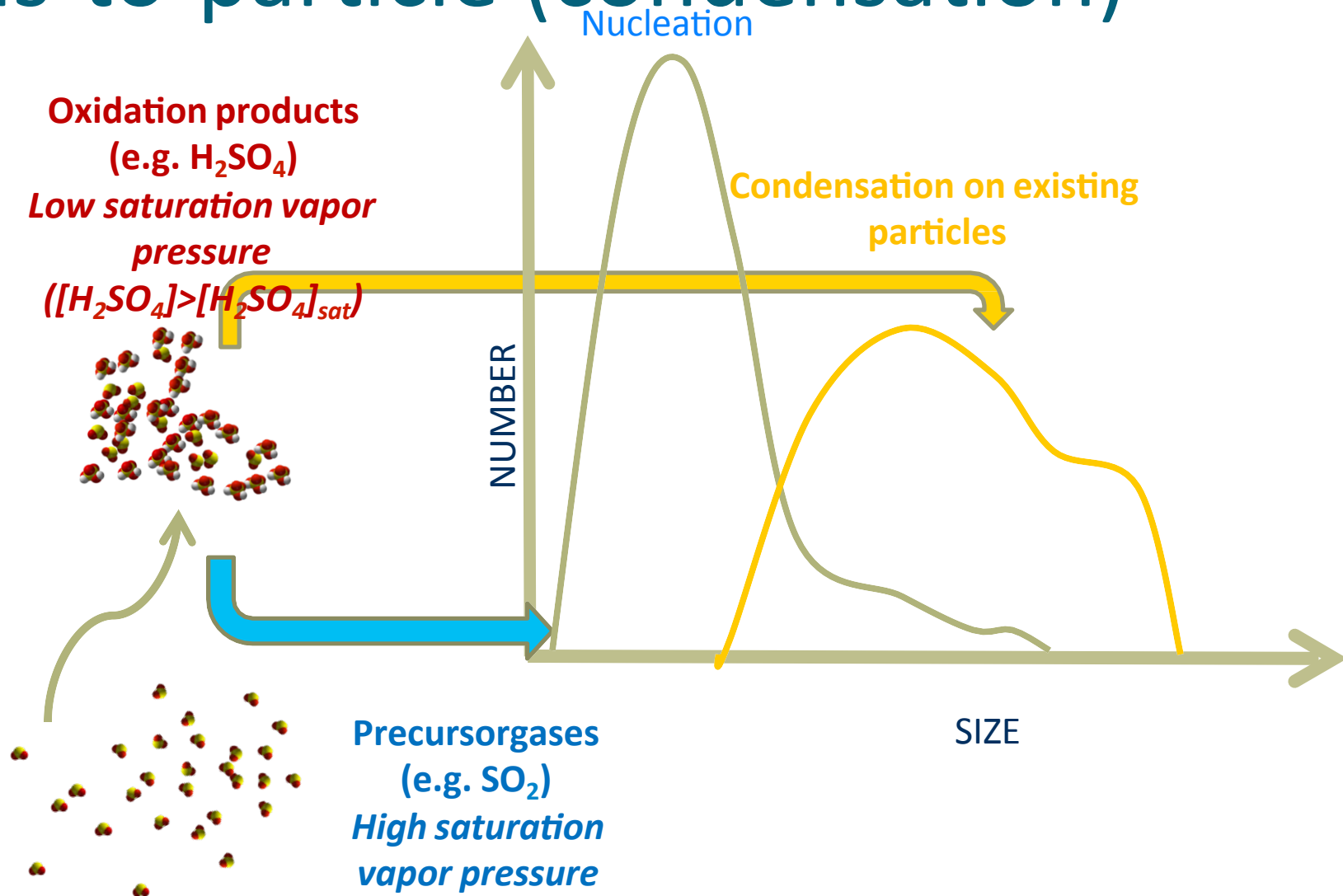
$$K_{i,j}^B = 4\pi(r_i + r_j)^2 \sqrt{v_i^2 + v_j^2}$$



Coagulation



Gas-to-particle (condensation)



Condensation

$$S = \frac{p_a}{p_a^s(T)}$$

$S < 1$, *Subsaturation*

$S > 1$, *Supersaturation*

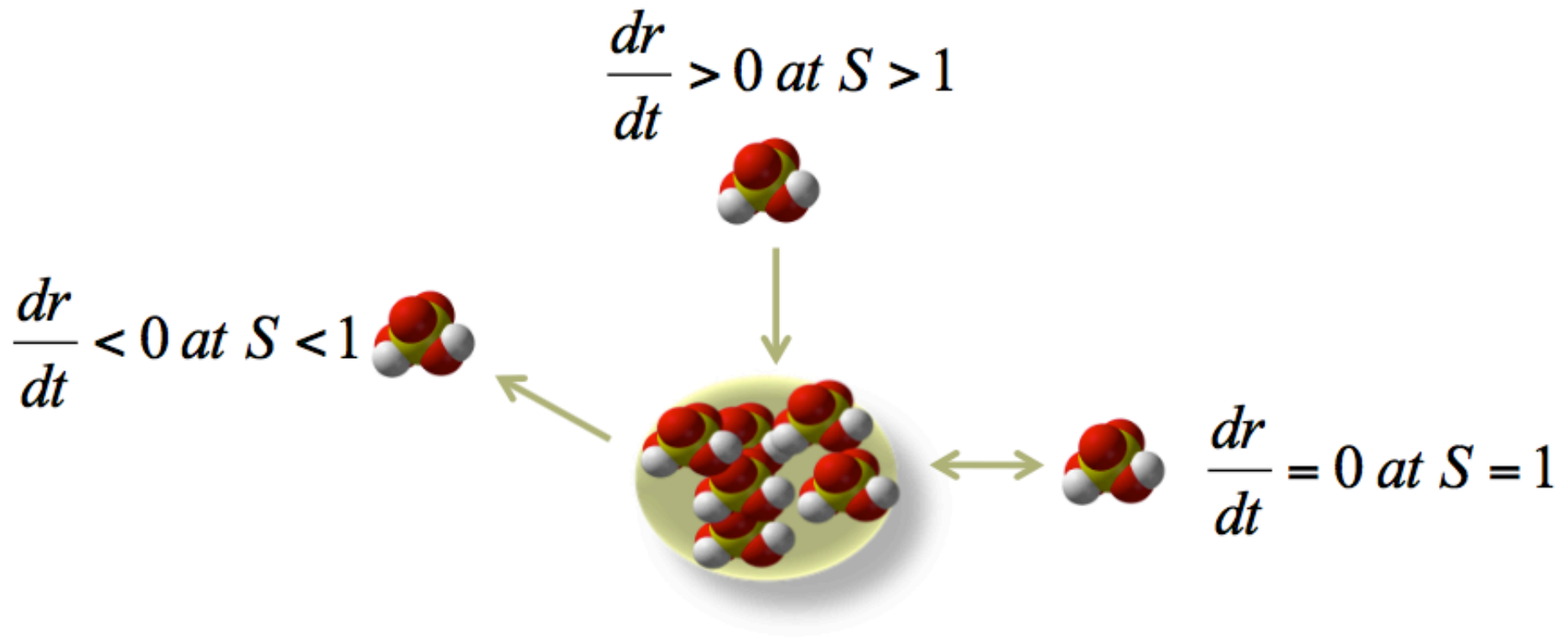
$S = 1$, *Saturation*

S =saturation ratio
 p_a =partial pressure of a
 p_a^s =saturation vapor
pressure of a at
temperature T

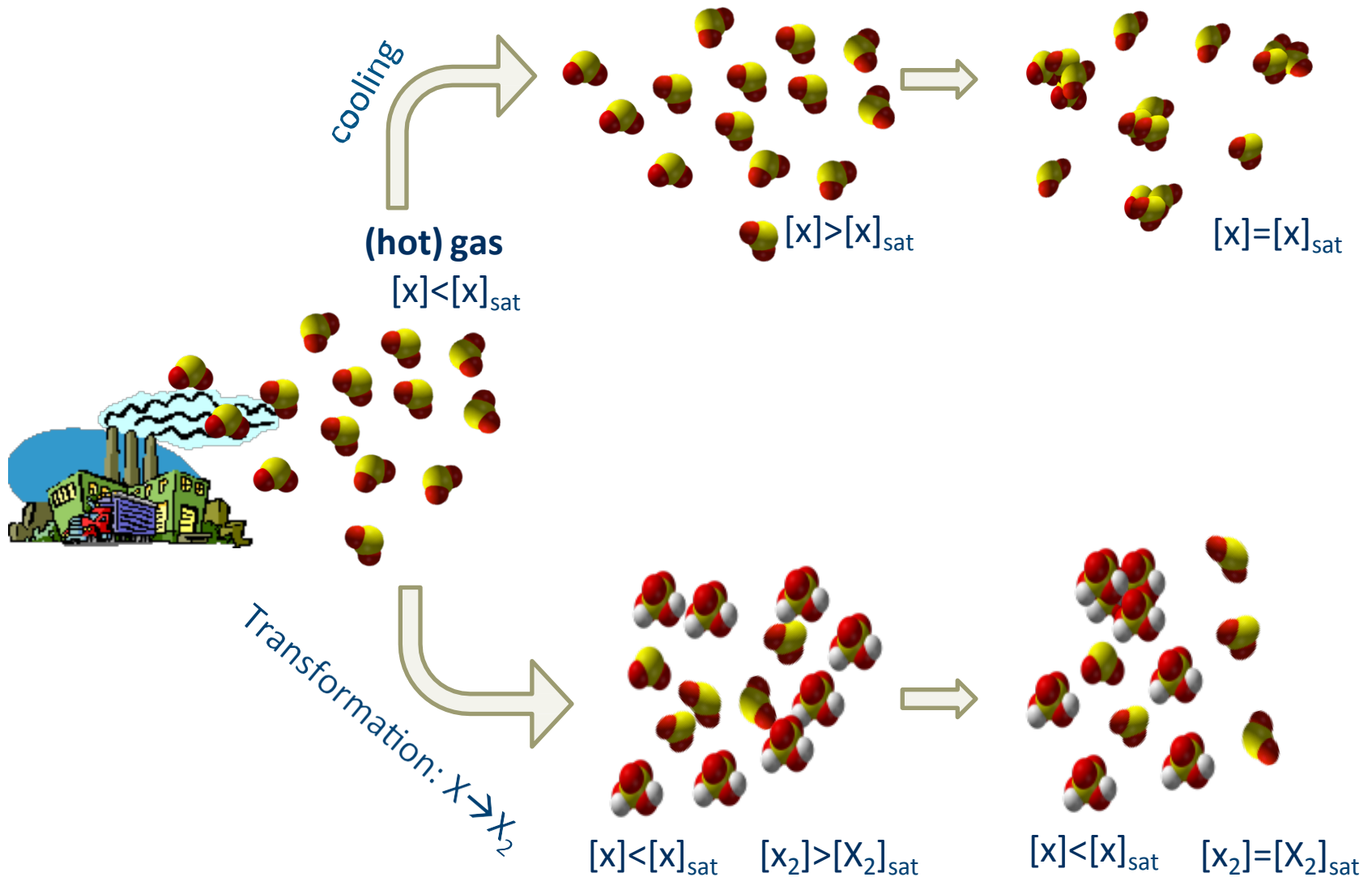
$$\ln\left(\frac{p_{a,1}^s}{p_{a,2}^s}\right) = \frac{\Delta H_{vap}}{R} \left(\frac{1}{T_2} - \frac{1}{T_1}\right)$$

Clausius-Clapeyron relation

Condensation



How is supersaturation reached?



Nucleation x Condensation

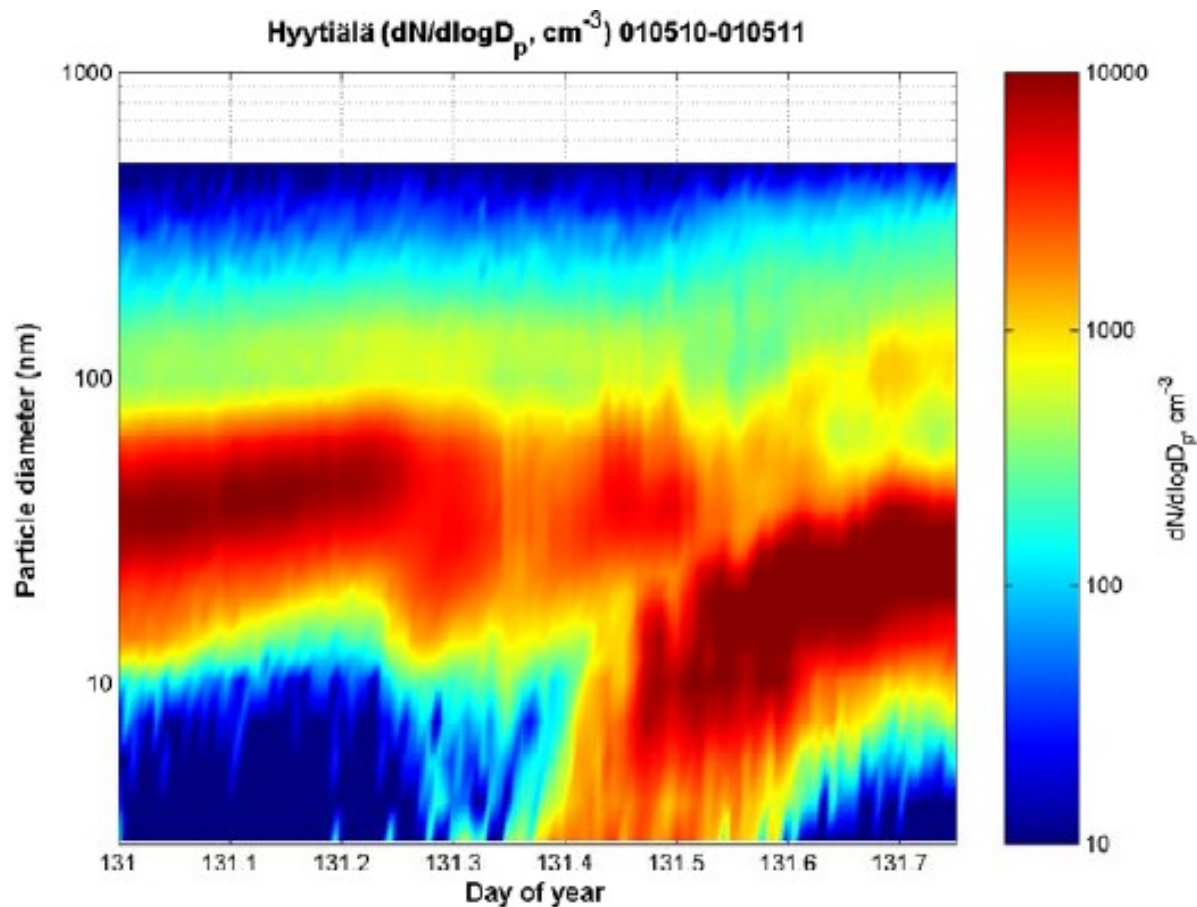
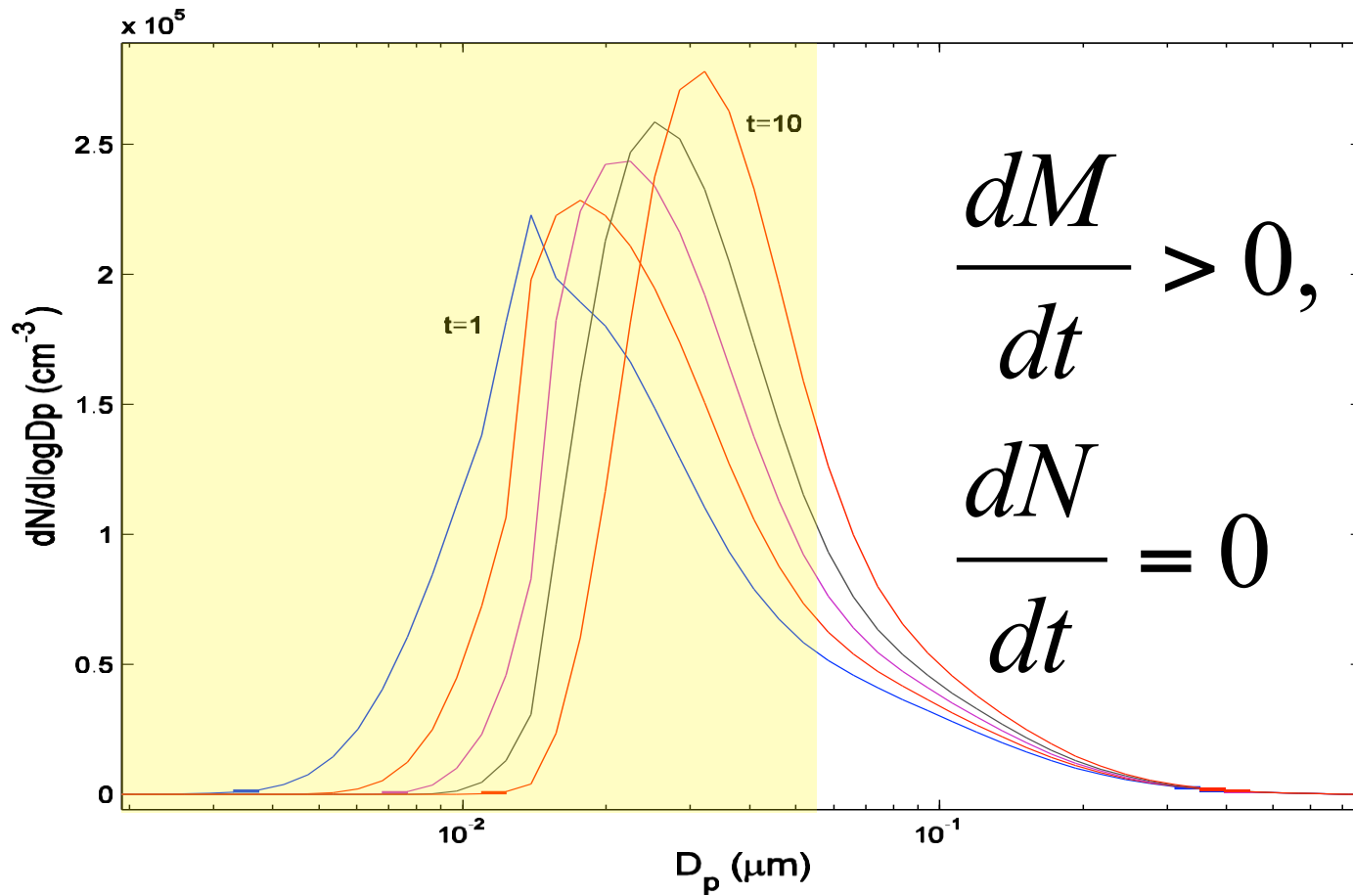


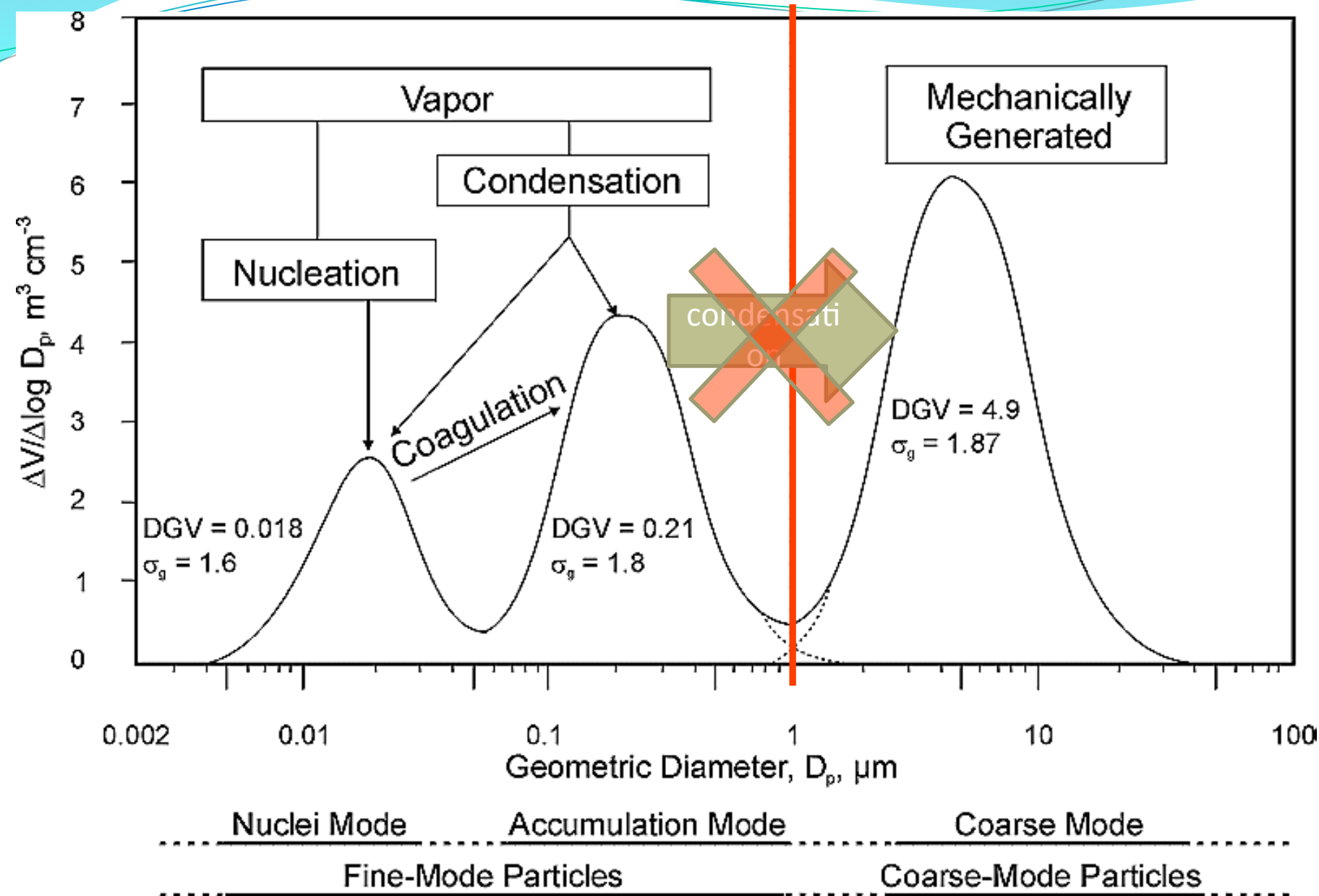
Figure 2: Example of nucleation event observed on 11th May 2001, Hyytiälä (61.51°N, 24.17°E).

Number or Mass?

- Amount of pre-existing aerosol surface crucial
 - Generation of supersaturated conditions + **low surface area** of pre-existing particles favours formation of **particle number via nucleation**
 - Generation of supersaturated conditions + **High concentration** of pre-existing particles favours formation of **particle mass via condensation**

Condensation





Dry deposition

- **Dry deposition** occurs when a gas or particle is removed at an air–surface interface, such as on the surface of a tree, a building, a window, grass, soil, snow, or water.
- Dry-deposition speeds are generally parameterized as the inverse sum of a series of resistances

$$V_{d,part} = \frac{1}{r_t} + v_t$$

$$F = -V_d C$$

F=flux to surface $\text{m}^{-2}\text{s}^{-1}$

v_d =deposition velocity (m/s)

C=concentration of particles

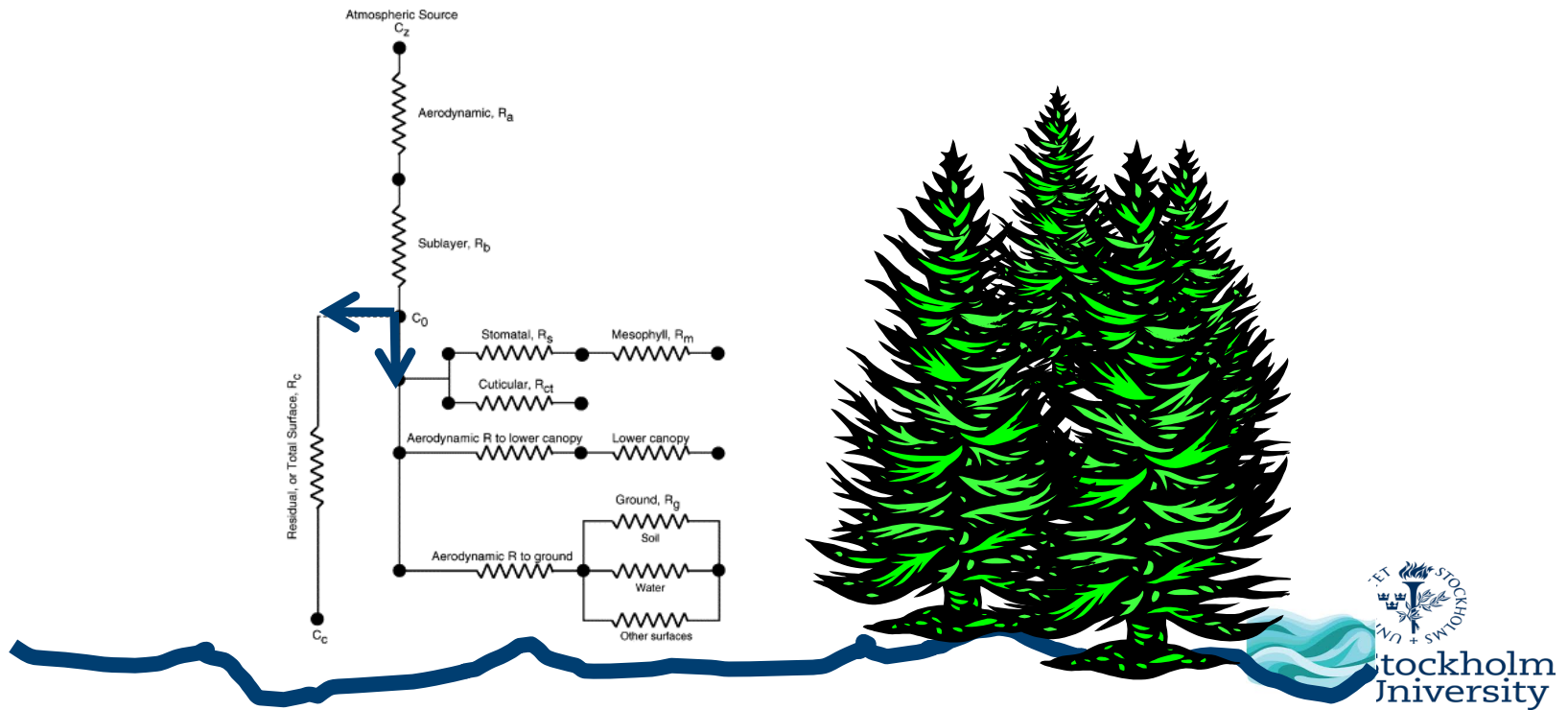
Dry deposition

Only removal path in the dry atmosphere

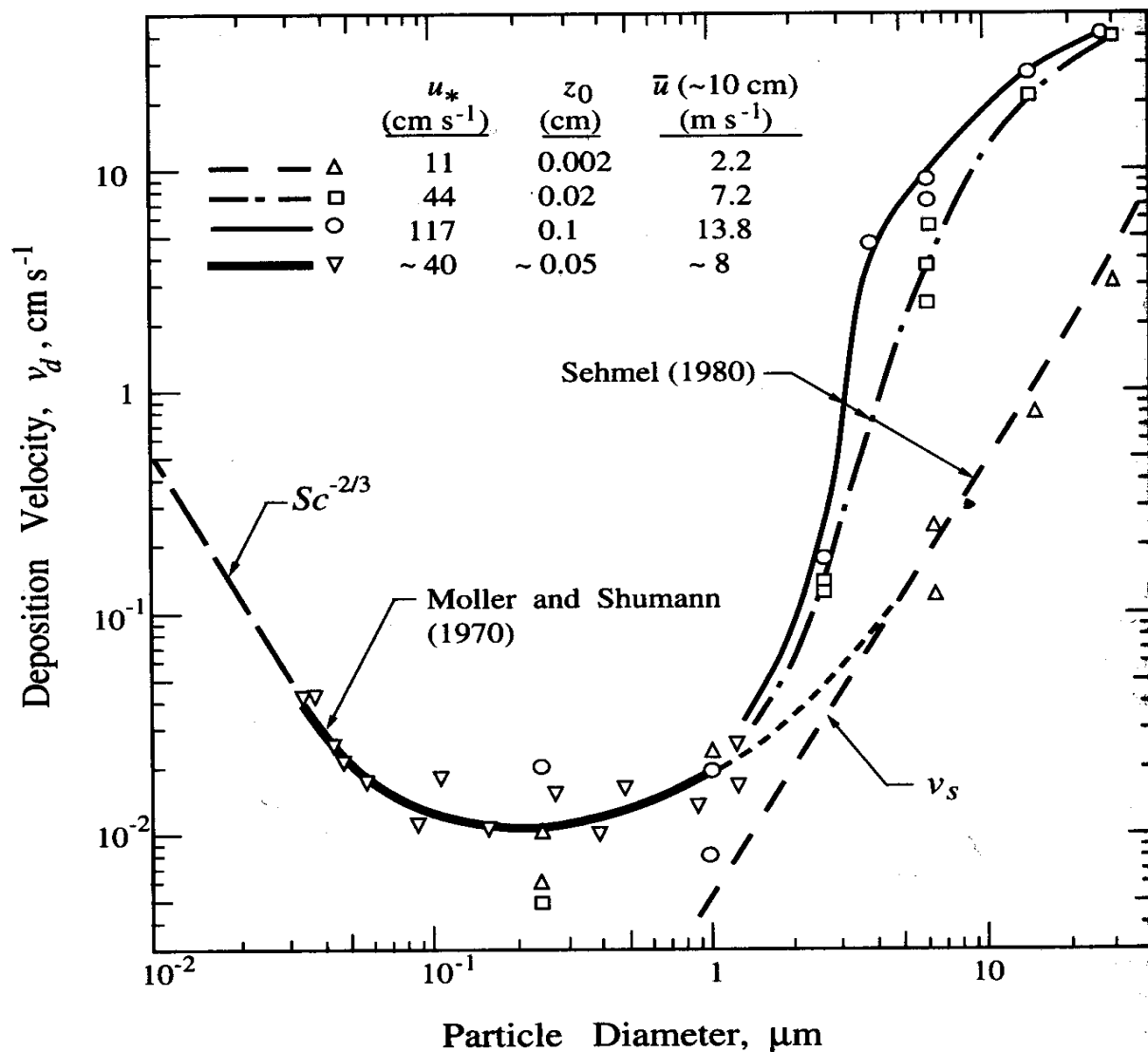
Depends on:

- Atmospheric turbulence
- Phase of species (gas or particle)
- Physio-chemical properties of depositing species
 - Particles: size, density
 - Gases: water solubility, reactivity
- Surface properties: Reactive? Sticky? Irregular? (eg vegetation)

Resistance analogy cont'd



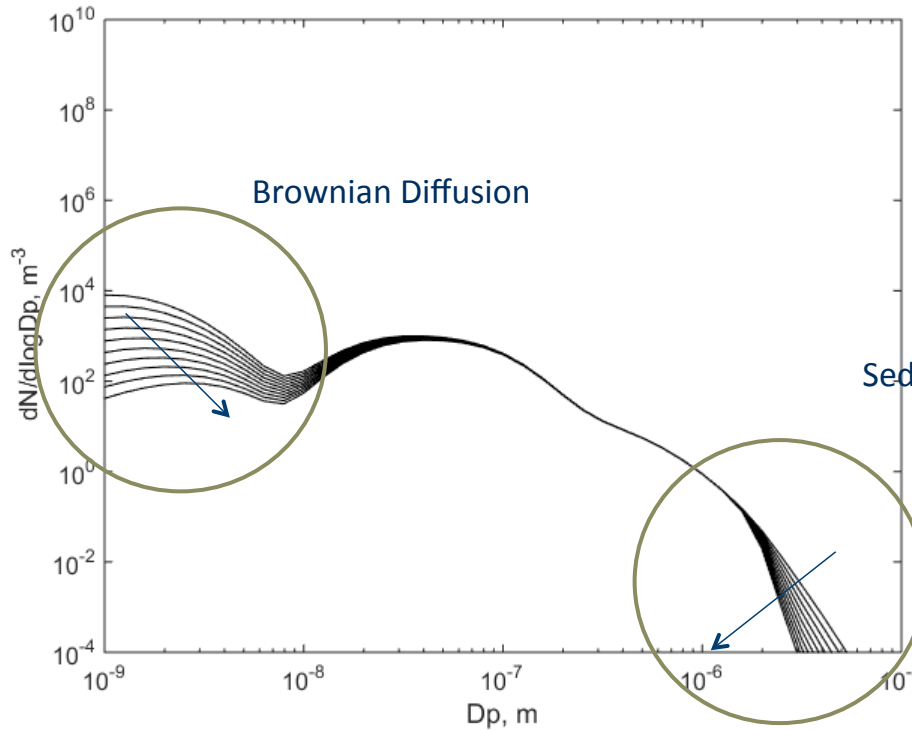
Dry deposition-strongly dependent on size



Ageing due to dry deposition

$$Sc = \mu_{air} / Diff_p$$

$$Sc^{-2/3}$$



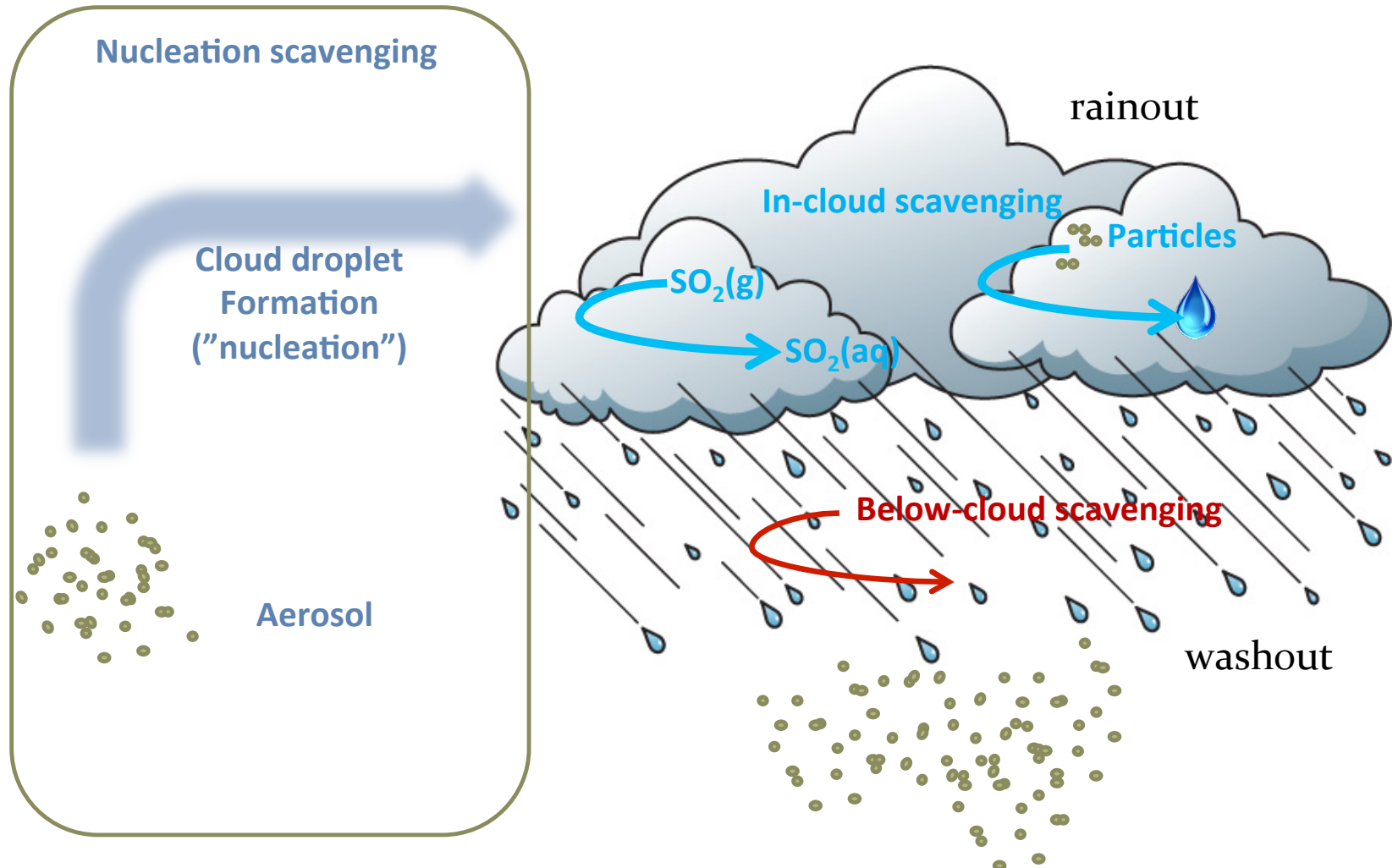
$$\frac{dM}{dt} < 0,$$

$$\frac{dN}{dt} < 0$$

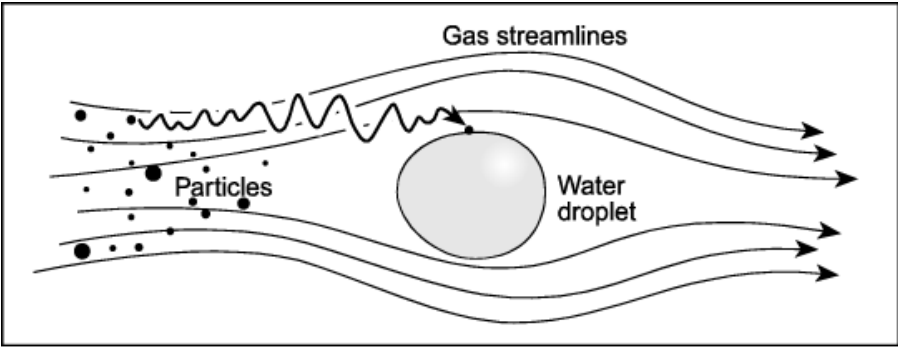
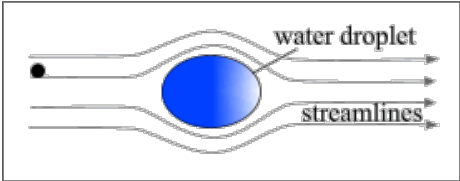
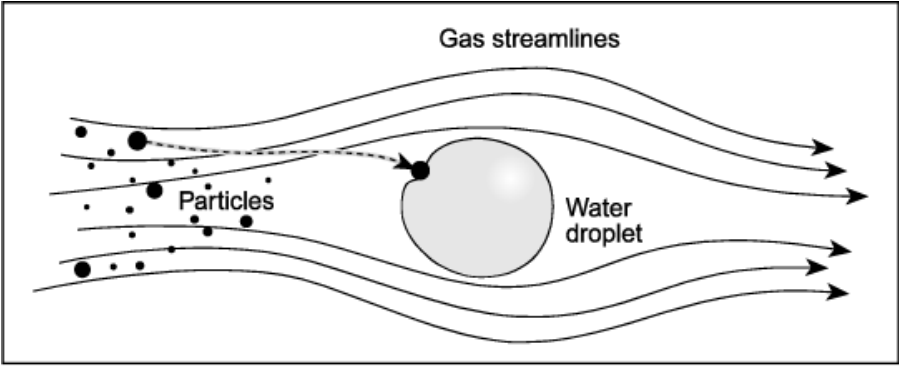
Sedimentation, impaction

$$10^{-3/St}, St = \frac{u^{*2} V_s}{g\mu} \text{ and } V_s$$

Wet deposition



Scavenging of aerosols: Impaction, diffusion and interception



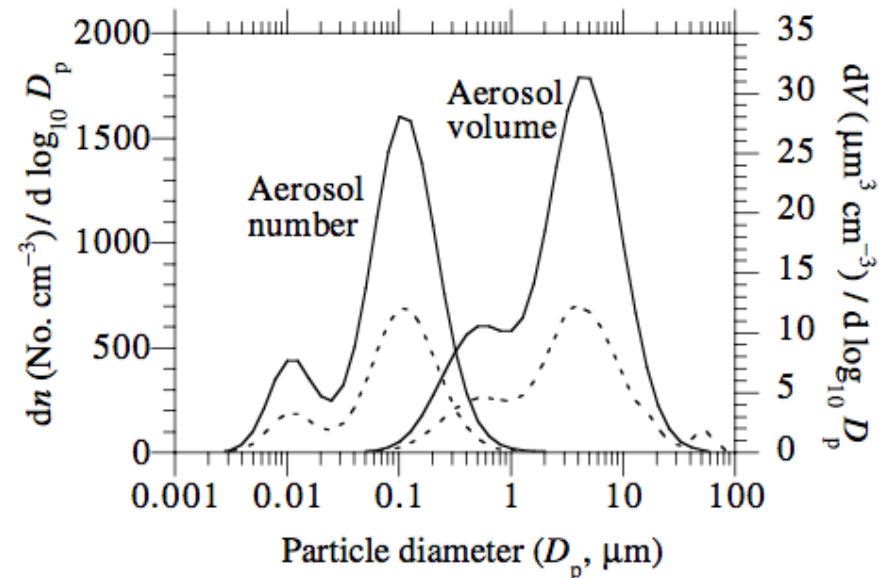
Scavenging

Within a cloud

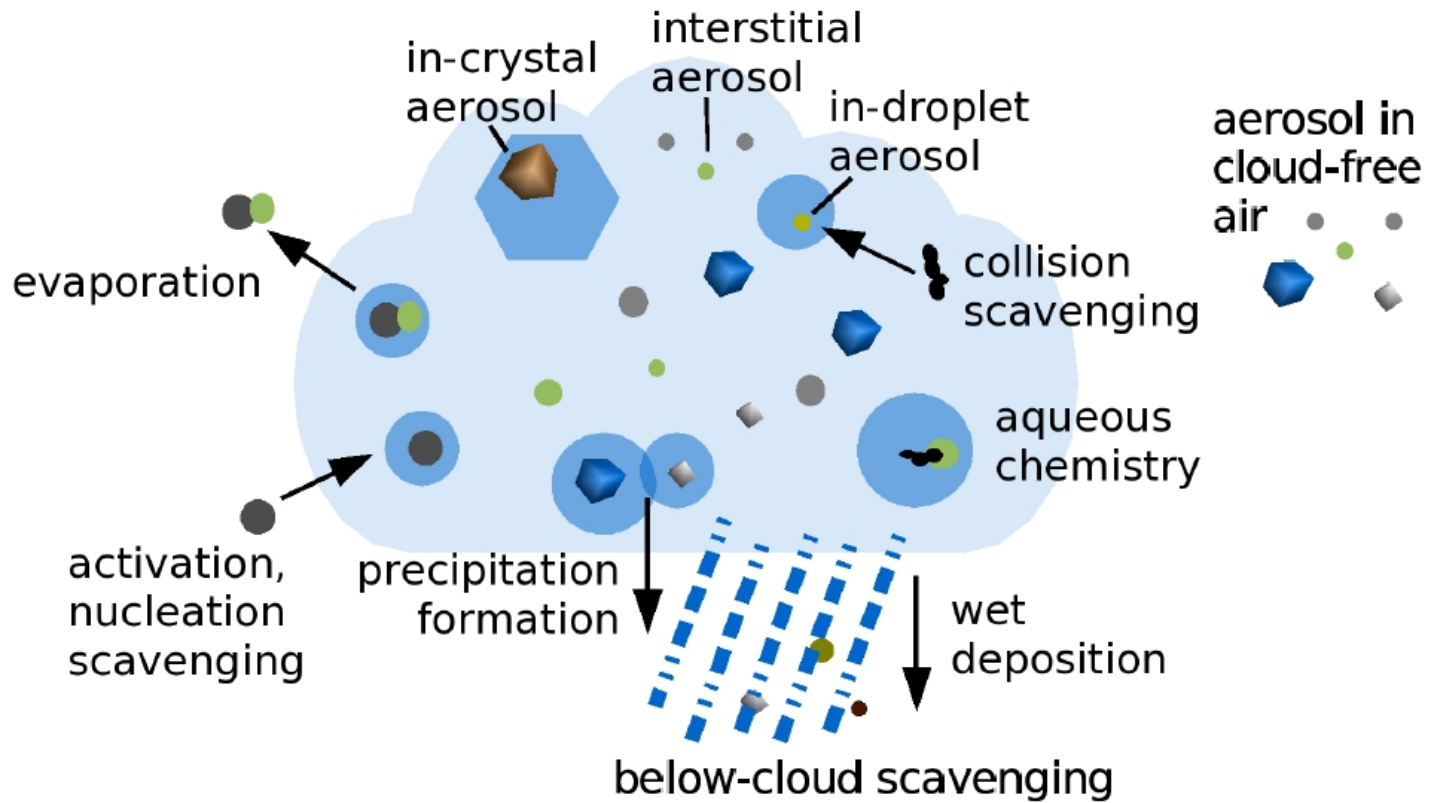
- Rainout scavenges all large and most midsize particles, which have large mass, before washout has a chance to remove particles within a cloud.

Below a cloud (no rainout)

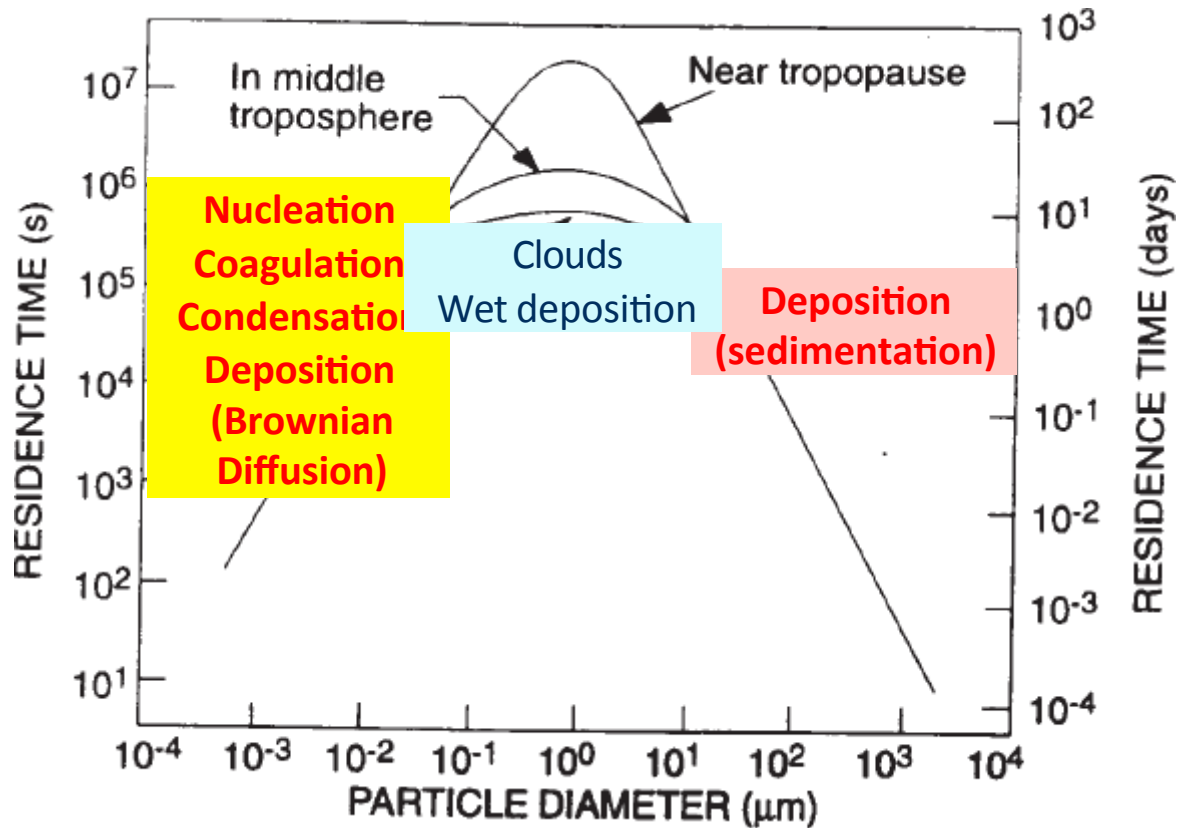
- Washout generally removes more particles by number than does rainout.



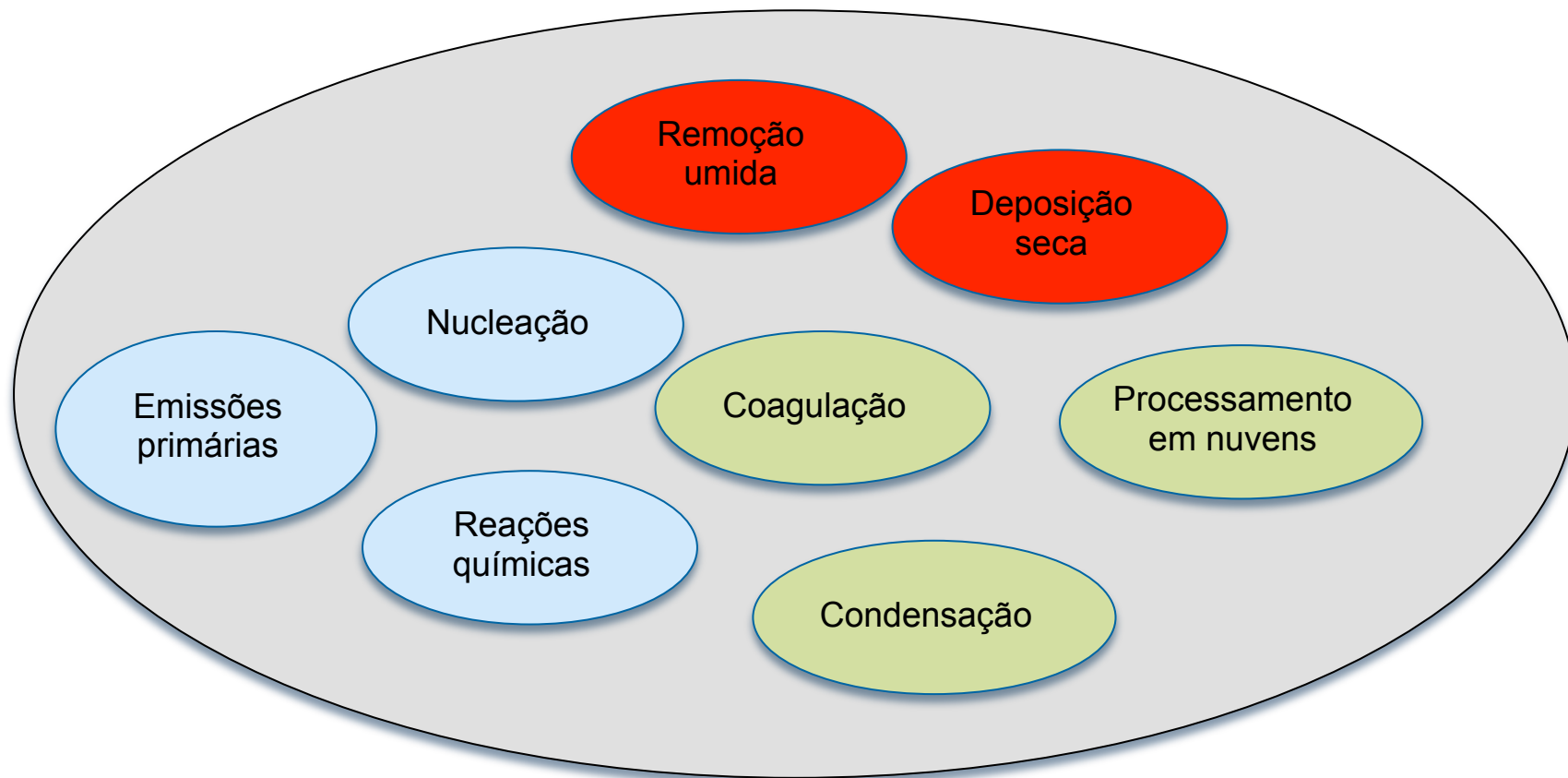
Cloud processing



Explaining the residence time



Processos envolvendo aerossóis



Model treatment of aerosol

Individual processes or sets of processes are treated by “modules” or “schemes”.

