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Organosulfates in aerosols downwind of an urban region in central Amazon⁺

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Organosulfates are formed in the atmosphere from reactions between reactive organic compounds (such as oxidation products of isoprene) and acidic sulfate aerosol. Here we investigated speciated organosulfates in an area typically downwind of the city of Manaus situated in the Amazon forest in Brazil during "GoAmazon2014/5" in both the wet season (February-March) and dry season (August-October). We 15 observe products consistent with the reaction of isoprene photooxidation products and sulfate aerosols, leading to formation of several types of isoprene-derived organosulfates, which contribute 3% up to 42% of total sulfate aerosol measured by aerosol mass spectrometry. During the wet season the average contribution of summed organic sulfate concentrations to total sulfate was $19 \pm 10\%$ and similarly during the dry season the contribution was 19 \pm 8%. This is the highest fraction of speciated organic sulfate to 20 total sulfate observed at any reported site. Organosulfates appeared to be dominantly formed from isoprene epoxydiols (IEPOX), averaging 104 ± 73 ng m⁻³ (range 15–328 ng m⁻³) during the wet season, with much higher abundance 610 ± 400 ng m⁻³ (range 86–1962 ng m⁻³) during the dry season. The concentration of isoprene-derived organic sulfate correlated with total inorganic sulfate ($R^2 = 0.35$ and 0.51 during the wet and dry seasons, respectively), implying the significant influence of inorganic sulfate 25aerosol for the heterogeneous reactive uptake of IEPOX. Organosulfates also contributed to organic matter in aerosols (3.5 \pm 1.9% during the wet season and 5.1 \pm 2.5% during the dry season). The present study shows that an important fraction of sulfate in aerosols in the Amazon downwind of Manaus consists of multifunctional organic chemicals formed in the atmosphere, and that increased SO₂ emissions would substantially increase SOA formation from isoprene.

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Environmental significance

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Aerosols in the atmosphere affect climate and human health, but there are still major gaps in our understanding of fundamental processes including the interaction of natural and man-made emissions. Tropical vegetation emits large amounts of isoprene which undergo photochemical oxidation in the atmosphere. Reactions between oxidation products of isoprene can react with acidic sulfate aerosols, typically of anthropogenic origin, to form organosulfates. Here we find that these organosulfates make up a surprisingly high fraction of total sulfate in aerosols downwind of the city of Manaus situated in the Amazon forest. Increasing anthropogenic emissions of SO2, forming sulfate aerosols, is thus also expected to increase formation of secondary organic aerosols from biogenic isoprene in other tropical regions with high isoprene emissions.

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Introduction 1 1.

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Isoprene (2-methyl-1,3-butadiene), emitted from vegetation, is the most important non-methane hydrocarbon emitted to the atmosphere (about 500 Tg per year),¹ where it undergoes photooxidation, forming a wide range of gaseous and particulate products.²⁻⁵ Tropical forests, such as the Amazonian rainforest, emit particularly high amounts of isoprene and other terpenoids.^{1,6} A substantial fraction of sub-micron organic aerosol mass in tropical regions originates from photochemical oxidation of isoprene7-11 affecting e.g. aerosol hygroscopicity and formation of cloud condensation nuclei.12-14

Heterogeneous acid-catalyzed uptake of reactive isoprene oxidation products such as isoprene epoxydiols (IEPOX) and hydroxyhydroperoxides onto acidic sulfate aerosol particles leads to formation of organosulfates, in addition to other products such as methyl tetrols and their oligomers3,15-19 and thus contributes to the mass concentration of secondary organic aerosol (SOA). This uptake is most efficient for sulfate aerosol particles not coated by other organics,^{20,21} and can contribute to anthropogenic enhancement of biogenic SOA formation,²²⁻²⁴ as previously observed *e.g.* in the southeastern US.^{22,25} Analysis of specific molecular tracers in aerosols can provide information on sources and atmospheric processes.^{26,27}

The region of Manaus was identified as a unique site for studying interactions of anthropogenic and biogenic emissions, since this city of about 2 million inhabitants is situated in the central region of the Amazon basin, which extends for more than 1000 km in every direction.^{28,29} Power generation from combustion of fuel oil and diesel, as well as oil refinery operations, emit about 72 kton SO₂ per year in the region.³⁰ Additional sources of SO₂ and sulfuric acid in the Amazon include photochemical oxidation of dimethylsulfide (DMS) emitted from marine or terrestrial sources.^{8,31,32} Tropical rainforests have been identified as a source of DMS, which may constitute an important regional source, though small on a global scale (about 6%),³³ and affect aerosol formation and the hydrological cvcle.34

Recently de Sá et al.35 reported that urban pollution from Manaus affected the formation of isoprene-derived SOA. In particular, they observed that sulfate enhanced formation of IEPOX-SOA, while NO suppressed it, consistent with expectations based on known isoprene oxidation chemistry.35,36 Using ultra-high-resolution mass spectrometry, Kourtchev et al.37 detected more than 2000 elemental formulae in aerosol samples collected downwind of Manaus. Up to 60% of the organic species were nitrogen- and/or sulfur-containing, and the dominant ion in most samples had the elemental formula of the isoprene-derived IEPOX organosulfate.38

In the present work, we investigate the contribution of organosulfates to the observed PM in the region using analysis of molecular tracers of aerosol sources and processes. Specifically, aerosol samples collected during both the wet season and the dry season were analyzed for organosulfates. Of particular interest is the contribution of organosulfates to total sulfate. In a previous study the average contribution of organic sulfate to

total sulfate was estimated as about 5% (0-16%) for 12 sites distributed across USA, showing higher organic sulfate contributions to total sulfate in regions of lower sulfate concentrations.³⁹ In comparison, airborne measurements of IEPOX OS and acetic acid sulfate (also known as glycolic acid sulfate) in the tropospheric boundary layer in the eastern and western US showed ratios of organic sulfate to total sulfate of 3.5% and 2.4%, respectively.40 Studies in other regions with very low isoprene emissions typically find lower average contributions of organic sulfate to total sulfate, such as a study in Northern 10 Europe during winter $(1.4 \pm 0.9\%)$ where organosulfate precursors were identified as monoterpenes and alkanes.41

Additional information on the biogenic sources of SOA was obtained in the present study from analysis of carboxylic acids 15 from photochemical oxidation of monoterpenes. This methodology has previously been applied to study sources of SOA in Europe.^{42,43} In this study we investigated molecular tracers of monoterpene SOA include pinic acid, pinonic acid and terpenylic acid from photochemical oxidation of α - and β -pinene,⁴⁴⁻⁴⁷ 20 in order to quantify SOA contributions from these potential In addition we measured sources. 3-methyl-1,2,3butanetricarboxylic acid (MBTCA), which is formed through gasphase oxidation of pinonic acid, and is thus a tracer of photochemical processing of α- and β-pinene.48,49 25

2. Experimental

2.1 Sampling

Samples were collected as part of the Green Ocean Amazon 30 (GoAmazon2014/5) field campaign^{28,29} during intensive operating periods in February-March (IOP1 during the wet season) and August-October (IOP2 during the dry season) of 2014. Here we focus on measurements at site "T3" ($3.21^{\circ}S$, $-60.60^{\circ}W$), approximately 70 km downwind of Manaus (Fig. 1), situated on 35 a cleared, pasture site about 500 m from any trees, 2 km north of the nearest paved road and about 2.5 km from the closest surrounding forest.

The prevailing wind direction was from the north-east in the 40wet season and from east during the dry season, transporting air masses over the Amazon forest to be mixed with the outflow from Manaus (to the east) before reaching the T3 site (Fig. 1). Transport time from city to T3 was typically 4-5 h.³⁶ Backward trajectories were simulated using HYSPLIT4 model.⁵⁰ Forcing 45 meteorological conditions were taken from GDAS 3-hourly, global, 0.5° fields provided by NOAA's Air Resources Laboratory. Trajectories started from 10 m over T3 site, at every hour, for the period of our analysis.

Submicron aerosol samples (PM₁) were collected on quartz 50 fiber filters (101.6 mm diameter, QM-A Quartz, Whatman) using a custom-built sequential filter sampler (Aerosol Dynamics, Inc.). Air was sampled 4 m above ground level, passed through tubing kept at temperatures below the dew point for trapping excess water, and then through a greaseless cyclone (with 1 μ m 55 aerodynamic diameter cut-point), before collection of particles. This sampling system is also expected to remove water-soluble reactive gas-phase species and thus reduce potential sampling artefacts.⁵¹ Standard conditions were defined as 21.1 °C and

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Fig. 1 Location of sampling site (yellow star), downwind of Manaus (red circle), in the central Amazon region in Brazil, and trajectory analysis for samples collected during the wet season (left) and dry season (right). Samples were selected to represent periods when airmasses arrived at the site after passing over the urbanized region of Manaus. Percentages denote the fraction of trajectories following each route.

101.3 kPa for the TSI flowmeter used to check sampling flow rate. Filters were pre-treated by baking them at 550 °C for 12 hours.

During IOP1 (wet season), sample collection occurred with 12 hour time resolution, from 06:15–18:00 and 18:30–06:15, local time (UTC-4). During IOP2 (dry season), sample collection occurred with 4 hour time resolution, from 06:00–09:55, 09:55–13:50, 13:50–17:45, 18:15–22:10, 22:10–02:05, and 02:05–06:00, local time. Field blanks were collected each week. Particle filters were kept frozen (-18 °C) and always transported on ice.

⁵⁰ 2.2 Analyses

Filter aliquots (47 mm diameter punches) were extracted in 1:1 mix of methanol and acetonitrile (AcN), aided by sonication for twenty minutes. The extract was filtered using a polytetra-fluoroethylene syringe filter (Q-Max, Frisenette) and evaporated under a flow of dry, high-purity nitrogen gas. The residue was re-dissolved in 200 μ L of 10% AcN with 0.1% acetic acid in MilliQ-water and analyzed using an ultrahigh performance liquid chromatograph coupled through an electrospray

ionization source to a quadrupole time-of-flight mass spectrometer (UHPLC/ESI-qTOF-MS, Bruker Daltonics) operated in negative ionization mode. The extraction and analytical method, as well as identification and quantification of organosulfates, are described in detail elsewhere.^{42,51,52}

Compounds were identified based on comparison of mass spectra and retention times with authentic standards when 45 available or according to mass spectra and comparison with data from previous studies. An authentic standard of an organosulfate with molecular weight 216 (OS 216), derived from IEPOX,⁵³ was kindly obtained from Prof. J. Surratt and Prof. A. 50 Gold (University of North Carolina - Chapel Hill) and used for quantification of all isoprene-derived organosulfates. Standards for carboxylic acids and monoterpene organosulfates have previously been listed.41,54 Six-point calibration curves generally showed $R^2 > 0.99$. All concentrations were corrected according 55 to recovery (59 \pm 15%) except OS 216 which was only corrected according to ionization efficiency. Average blank concentrations, which were <0.3 ng m⁻³ for all organosulfates except OS 200 (0.8 ng m⁻³), 0.5 ng m⁻³ for pinonic acid, and 0.02 ng m⁻³

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Mass concentrations of particle sulfate, organic mass, ammonium and nitrate species were measured using a High-Resolution Time-of-Flight Aerosol Mass Spectrometer from Aerodyne, Inc., USA (HR-ToF-AMS, hereafter denoted AMS).^{55,56} Details of the specific deployment were previously described.³⁵ The AMS measures aerosol composition after vaporization, typically at 600 °C, followed by electron ionization at 70 eV, which is thought to lead to fragmentation of organosulfates into separate sulfate and organic fragments.⁵⁷ In the current version, the instrument can thus detect the total sulfate concentration, i.e. organic plus inorganic, but cannot separate these two components.

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3. Results and discussion

3.1 Overview

Organosulfates detected in atmospheric aerosols are typically 20 reported based on their suggested precursor molecule. In this work we instead use a more stringent nomenclature in which organosulfates are identified by the structure of the molecule containing a sulfate group bonded to one carbon, e.g. acetone 25 sulfate, instead of a suggested precursor molecule, *i.e.* hydroxvacetone sulfate. This nomenclature in accordance with standard IUPAC nomenclature will make it easier to recognize the molecular structures of organosulfates and avoid confusion as there can be several or unknown precursors to a specific orga-30 nosulfate. For simplicity we will here use the abbreviation OS for organosulfate, followed by its molecular weight (i.e. OS 184)

and IEPOX OS for 1,2,4-hydroxy-2-methyl-butyl-3-sulfate.

Table 1 lists average concentrations, as well as ranges, of major detected compounds and sulfate and organic aerosol (OA) measured by AMS during the aerosol sampling periods. Additional information on suggested molecular structures and precursors of organosulfates is provided in ESI Table S1.† The main observed organosulfates were derived from isoprene (namely OS 184, OS 200, OS 212, OS 214 and IEPOX OS15,58,59) and its oxidation products hydroxyacetone and hydroxyacetic acid (acetone sulfate and acetic acid sulfate).59-61 Concentrations of organosulfates were dominated in both the wet and the dry season seasons by IEPOX OS (87 \pm 4% and 61 \pm 15%, respectively), in line with previous results from the same site.³⁷

Summed concentrations of all detected organosulfates were observed at average concentrations of 104 ± 73 ng m⁻³ during the wet season (IOP1), and 612 \pm 397 ng m⁻³ during the dry season (IOP2).

3.2 Results from the wet season

The organosulfate derived from IEPOX, OS 216, constituted about 80% of the organosulfate mass in the samples (Table 1 and Fig. 2). The variation in organosulfate levels between samples showed some relation with air pollution parameters such as ozone and sulfate (Fig. S1[†]). High correlations across different samples ($R^2 > 0.8$) were found between IEPOX OS and acetone sulfate (from hydroxyacetone),60 hydroxy-butanone1

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sulfate (from methyl vinyl ketone),59,62 and OS 214 (also derived from isoprene),⁵⁸ as listed in Table S2.† Acetic acid sulfate, OS 168, OS 170 and OS 182 from the isoprene oxidation products hydroxyacetic acid, methacrolein, 2-hydroxy propanoic acid and 2-methylglyceric acid were not observed at levels above the detection limit, in contrast to other studies in nontropical regions and laboratory studies.59-63 During the wet season, OS 200 was only observed at very low levels and was not correlated with other measured OS. Previous studies found that OS 200 can have several precursors, including the isoprene 10 high-NO_r oxidation products 2-methylglyceric acid and methacrylic acid epoxide (MAE),60,64 and also 2-methyl-3-buten-2-ol (typically emitted from vegetation in temperate climates).65

In some laboratory studies, OS 200 and OS 184 were both formed from ozonolysis of isoprene, rather than photooxidation.^{59,66} A study at a forested site in the south-western US found much lower formation of OS 200 compared to IEPOX-OS, especially at higher temperatures, due to differences in production of their respective precursors (MAE vs. IEPOX).67

20 Of organosulfates derived from monoterpenes, only OS 250 from α-pinene⁶⁸ was observed at concentrations above the detection limit. The highest concentration of OS 250 (3.6 ng m^{-3}) was measured on March 10 (Fig. 2), when other oxidation products of α -pinene and β -pinene also showed slightly elevated 25 concentrations (Fig. S2[†]), including first-generation products pinic acid and terpenylic acid. In general, carboxylic acids from oxidation of α -pinene and β -pinene were observed at low levels in the particle phase (Table 1 and Fig. S2[†]). For some compounds such as pinonic acid this is partly due to their 30 relatively high vapor pressure and the high temperatures at this site, as previously seen.69

Organosulfate species from other sources were not observed above the detection limit or only at very low concentrations, including those formed from green-leaf volatiles, alkanes and 35 polycyclic aromatic hydrocarbons.70-72

3.3 Results from the dry season

40Concentrations of organosulfates were about four times higher on average during the dry season than during the wet season (Table 1). The main organosulfate compounds are IEPOX OS and OS 214 (also from isoprene), while the remaining organosulfates never make up more than 20% of the total organo-45 sulfate concentration.

The largest difference was observed for OS 200, with dry season average concentration (30 ± 13 ng m⁻³) about 40 times the wet season average concentration (0.7 \pm 0.8 ng m⁻³). Additional OS species were detected in the dry season, namely 50 acetic acid sulfate (OS 156) from glycolic acid and OS 170 from methyl glycolic acid or green-leaf volatiles.59,61,70 Other types of organosulfates, either derived from green-leaf volatiles, alkanes and polycyclic aromatic hydrocarbons, were not observed above the detection limit or only at very low concentrations during the 55 dry season.

Furthermore, while concentrations of some first-generation oxidation products of α -pinene and β -pinene (pinonic acid and terpenylic acid) were not observed above the detection

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Table 1 Co	oncentrations of molecular tracers, together	with sulfate and	d organic aerosol concentrations measured by AMS. All concentrations are
ng m ⁻³ . No	ote that n.d. means that the compound was	"not detected"	" above the detection limit (see Experimental section)

	Wet season		Dry season		
Compound	Average \pm 1SD ($n = 28$)	Range	Average \pm 1SD ($n = 70$)	Range	5
Acetone sulfate (OS 154)	4.6 ± 3.2	0.9-14.1	30 ± 15	n.d. to 70	
Acetic acid sulfate (OS 156)	n.d	n.d.	25 ± 11	n.d. to 48	
Propanoic acid 2-sulfate (OS 170)	n.d.	n.d.	8.8 ± 3.5	n.d. to 19.7	
Hydroxy-butanone-sulfate (OS 184)	2.5 ± 1.6	0.7-6.3	19.0 ± 9.0	n.d. to 45.6	
OS 200	0.7 ± 0.8	n.d. to 3.0	30 ± 13	n.d. to 71	10
OS 212	1.3 ± 1.5	n.d. to 5.1	n.d.	n.d.	
OS 214	11.5 ± 9.3	1.5 - 41.3	101 ± 61	n.d. to 283	
IEPOX OS (1,2,4-hydroxy-2-methyl-butyl-3-sulfate, OS 216)	83 ± 59	11-260	399 ± 308	39-1450	
Total OS	104 ± 73	15-328	612 ± 397	86-1962	
Pinonic acid	2.0 ± 1.5	0.3 - 5.4	n.d.	n.d.	
MBTCA	4.4 ± 3.0	0.9-10.8	14.1 ± 6.2	n.d. to 32.0	15
Terpenylic acid	1.8 ± 3.1	0.4-16.8	n.d.	n.d.	
Total sulfate	268 ± 136	59-550	1600 ± 917	595-4616	
Organic aerosol (OA)	1671 ± 579	625-2680	8783 ± 2642	4539-16 064	

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limit, the second-generation oxidation product MBTCA showed an average concentration of 14.1 ± 6.2 ng m⁻³, which was higher than during the wet season. This is in line with a longer photochemical aging time during the dry season, while rain-out during the wet season results in a larger influence of early oxidation products compared to more oxidized compounds. It should be noted that only a small fraction of pinonic acid (about 0–30%) is expected to be in the particle phase at this site, while MBTCA is primarily found in the particle phase.⁶⁹ resolution during the dry season allowed us to study the diurnal variation of molecular tracers. Levels of organosulfates were on some days higher during daytime than nighttime (*e.g.* 28 August, 22 September, 26 September, 27 September), while the picture was less clear on *e.g.* 19 September and 10 October.

Isoprene-derived OS peaked during daytime (10:00–18:00) and had minimum concentrations during night (22:00–06:00), similar to sulfate concentrations (Fig. 4) and the typical diurnal variation of isoprene emissions.^{36,73} For comparison, full timeseries of pollutants such as ozone and sulfate are shown in Fig. S1.† This is also in line with previous results from the same

Fig. 3 shows concentrations of organosulfates in aerosol samples collected during the dry season. The higher time-



Fig. 2 Concentrations of OS in aerosol samples collected during the wet season (plotted according to local time). Note that the timeline is not continuous.

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site during the wet season, where peak concentrations of isoprene oxidation products (2-methyl tetrols and methyl butene triols) were observed during afternoon hours, probably caused by the diurnal variation in isoprene emission and oxidation.⁶⁹ Isoprene-derived OS are formed locally or regionally during daytime, while during night-time there is less production and only loss processes (such as chemical reactions and particle deposition).

Fig. 4 also shows the average diurnal variation in the dry 10 season of environmental parameters (relative humidity, RH, and temperature) and selected gases (ozone and nitric oxide + nitrogen dioxide, NO_x), as well as sulfate and OA measured by AMS. It is clear that sulfate, temperature and ozone all peak during daytime, while NO_x, and OA show lower concentrations 15 in the morning and peak during night, probably due to a combination of boundary layer dynamics and dynamics of pollutant emission and advection from Manaus. In the dry season, OA concentrations are strongly affected by emissions from biomass burning,²⁸ while the effect on organosulfate 20 concentrations is expected to be minimal, as biomass burning aerosols are typically less acidic.74

During the dry season the correlations between IEPOX OS and other OS were not as high as observed in the wet season. Only the coefficient of determination (R^2) between IEPOX OS and OS 214 (also from isoprene) was larger than 0.5, while R^2 for correlations of IEPOX OS with acetone sulfate, acetic acid

sulfate, OS 184 and OS 200 were in the range 0.3–0.45 (see Table S2†). The correlation between IEPOX OS and sulfate is high during both the wet season ($R^2 = 0.50$) and the dry season ($R^2 = 0.69$), while the correlations between sulfate and other OS are lower (R^2 in the range 0.1–0.4). The correlation between OA and all OS species is also lower during the dry season (R^2 below 0.25).

There is a notable change regarding OS 200, which showed no correlation with IEPOX OS in the wet season but had $R^2 = 0.6$ in the dry season. Interestingly, OS 200 actually shows better 10 correlation with all other OS and MBTCA than IEPOX OS does. Since OS 200 can also be formed from isoprene ozonolysis (in addition to being a product of 2-methylglyceric acid), this is in line with a somewhat larger influence of isoprene ozonolysis during the sampling period in the dry season (18 ± 14 ppbv 15 ozone) than in the wet season (10.4 ± 6.6 ppbv ozone).

3.4 Contribution of organic sulfate to total sulfate

The observed contributions of organic sulfate to total sulfate are generally much higher than in most previous studies showing 0-16%.^{39-41,53} Organic sulfate constituted on average $19 \pm 10\%$ of total sulfate concentration measured by AMS during the wet season (range 7–37%) and $19 \pm 8\%$ of total sulfate (3–42%) during the dry season. The average contributions and ranges were thus similar during the two periods.



Fig. 3 Concentrations of OS in samples collected during the dry season (plotted according to local time). Note that the timeline is not continuous.

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Fig. 4 Average diurnal variations during dry season of (A) total sulfate, total organosulfates (total OS), and the ratio of OS sulfate to total sulfate, and (B) relative humidity (RH) and temperature (T), as well as (C) ozone, NO_x, and OA. Note that the data only includes time periods where samples in the current study were collected.

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In the southeastern US, another region with high isoprene emissions, isoprene-derived OS sulfate constituted only about 4.4% of total sulfate during summer.⁵³ During the wet season in the present study sulfate was much lower (0.27 \pm 0.14 µg m⁻³) compared to the southeastern US study $(1.9 \pm 1.2 \,\mu g \,m^{-3})$, while levels of isoprene-derived OS were more similar (average concentration of 104 ng m⁻³ during the wet season compared to 181 ng m^{-3} in the southeastern US). On the other hand, during the dry season the concentration of total sulfate (1.6 \pm 0.9 μg m^{-3}) was similar to the previous study in the US, but higher concentrations of isoprene-derived OS ($610 \pm 400 \text{ ng m}^{-3}$) were observed in the Amazon, indicating a factor of ~ 4 higher contribution of organic sulfate to total sulfate in our samples during the dry season.

55 A study in Northern Europe (during winter) found OS from monoterpenes and alkanes at higher levels than in the present study, while sulfate levels were in the same range $(0.2-6 \ \mu g \ m^{-3})$, average 1.9 \pm 1.6 μ g m⁻³) yielding an overall OS to total sulfate

40ratio of only 1.4 \pm 0.9%.⁴¹ This shows that besides sulfate concentration, other variables such as precursor reactivity, particle pH and coating, as well as RH, influence formation of organosulfates. Furthermore, low NO_r levels means that HO₂ reaction pathways and peroxide chemistry are relatively more 45 important in the Amazon region compared to previously investigated areas. This difference in photochemical oxidation pathways may also have implications for organosulfate formation.

Sulfate from isoprene-derived OS showed a clear association 50 with inorganic sulfate during both the wet and the dry season (Fig. 5). The slope of the least squares regression line (not shown) is 0.17 \pm 0.05 for IOP1 and 0.17 \pm 0.02 for IOP2. The correlation of OS sulfate with inorganic sulfate was notable with 55 $R^2 = 0.35$ during the wet season and $R^2 = 0.51$ in the dry season (both statistically significant with p-values < 0.01). During the wet season high OS concentrations often occur coincident with high OA levels (Fig. 5 and S3[†]), while the correlation between

total OS and OA ($R^2 = 0.30$) is lower than for OS and inorganic 1 sulfate (yet still significant with p < 0.01). The correlations indicate that both sulfate and OA may influence the formation of organosulfates during the wet season, when the atmospheric trace composition is more pristine. An association between OS 5 and OA is not observed during the dry season (Fig. 5 and S3[†]), possibly due to the influence of additional OA from other sources. Correlation analyses did not identify other significant parameters affecting the formation of organosulfates. This 10 included NO_x, ozone, temperature and RH.

The organic moiety of organosulfates contributed on average more than 3% to OA during both seasons. The average contribution was $3.5 \pm 1.9\%$ (range 0.8–7.3%) during the wet season, and $5.1 \pm 2.5\%$ (range 0.8–10.7%) during the dry season. This is also influenced by meteorological factors, as Fig. 1 shows that the sampling site was more often directly in the plume of Manaus during the dry season. The highest contributions of

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Fig. 5 Concentration of sulfate from isoprene-derived OS plotted against inorganic sulfate and colored according to concentration of organic PM₁ (AMS OA). Note the different scales for the two seasons.

organosulfates to OA were observed during daytime, when organosulfate concentrations typically peaked.

Formation of organosulfates through heterogeneous reactions has been shown in laboratory studies to depend on availability of acidic sulfate aerosols and their coating with SOA.^{20,21} At low RH (15%) the reactive uptake of IEPOX onto ammonium bisulfate particles is suppressed by coating with apinene SOA, while the influence of SOA coating is minimal at 50% RH.²¹ Reactive uptake of IEPOX is thus expected to be less sensitive to SOA coating of sulfate aerosols in tropical regions 10 where high relative humidity (>50% RH) is often observed. In the current study, air humidity was 74 \pm 15% RH during the dry season and $81 \pm 5\%$ RH during the wet season.

The acidity of sulfate aerosols is a key factor for reactive uptake of IEPOX.75 Measurements during the dry season showed that ammonium contained in aerosol particles was not sufficient to neutralize the particulate sulfate and nitrate.⁷⁶ A recent modeling study also suggests that the aerosol was highly acidic during the wet season of GoAmazon2014/5.77

Furthermore, uptake of IEPOX increases with increasing particle surface area.75,78 The influence of particle liquid water content (LWC) is complex, since higher LWC can increase aqueous phase processes and surface area, but also dilutes aerosol acidity and nucleophile (sulphate ion) concentration, reducing heterogeneous reactive uptake.79-81 Another yet uninvestigated factor is the possibility of aqueous phase oxidation of HSO_3^{-} by peroxides of isoprene, analogous to the mechanism suggested for peroxides from monoterpenes, which also was observed to form organosulfates.82

The very high organic sulfate to total sulfate ratio observed in the Amazon compared to other investigated sites globally can thus be explained by availability of reactive organic precursors (such as IEPOX) and presence of acidic sulfate aerosols at high relative humidity.

Overall, the results of the present study are relevant for understanding SOA formation in other regions in the tropics where highly urbanized settings coincide with high emissions of biogenic volatile organic compounds (VOC). It is especially important to further investigate the effect of sulfate aerosols, typically from combustion of fossil fuels but also from oxidation of naturally occurring DMS, for SOA formation in these regions, including effects on human health and regional climate. A better understanding of aerosol sources and formation processes will also lead to development of the most efficient mitigation measures.

4. Conclusion

In the present work we have investigated formation of organosulfate aerosol downwind of an urbanized region in the Amazon forest. The results show a high level of organosulfates derived from isoprene, which correlates well with concentration of total sulfate measured by AMS. On average, organic sulfate makes up 19 \pm 10% of total sulfate during the wet season and 19 \pm 8% during the dry season. Organosulfates also contribute to 3.5 \pm 1.9% of organic matter in aerosols during the wet season and 5.1 \pm 2.5% during the dry season. Additional analyses of 1

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molecular tracers and modeling studies are needed to improve our understanding of the formation of isoprene SOA and other sources of organic aerosols in general in the region. The findings of the present work show that acidic sulfate aerosols can critically affect SOA processes in tropical and sub-tropical areas with high emissions of isoprene. Increasing SO₂ emissions may thus lead to a substantial increase in SOA formation from biogenic isoprene in tropical forested regions.

10 Conflicts of interest

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References

- 40 1 A. B. Guenther, X. Jiang, C. L. Heald, T. Sakulyanontvittaya, T. Duhl, L. K. Emmons and X. Wang, The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1): an extended and updated framework for modeling biogenic emissions, Geosci. Model Dev., 2012, 5, 45 1471-1492.
 - 2 S. E. Paulson and J. H. Seinfeld, Development and evaluation of a photooxidation mechanism for isoprene, J. Geophys. Res., 1992, 97, 20703-20715.
- 3 M. Claeys, B. Graham, G. Vas, W. Wang, R. Vermeylen, 50 V. Pashynska, J. Cafmeyer, P. Guyon, M. O. Andreae, P. Artaxo and W. Maenhaut, Formation of secondary organic aerosols through photooxidation of isoprene, Science, 2004, 303, 1173-1176.
- 4 A. G. Carlton, C. Wiedinmyer and J. H. Kroll, A review of 55 Secondary Organic Aerosol (SOA) formation from isoprene, Atmos. Chem. Phys., 2009, 9, 4987-5005.
 - 5 F. Paulot, J. D. Crounse, H. G. Kjaergaard, A. Kurten, J. M. St. Clair, J. H. Seinfeld and P. O. Wennberg,

Epoxide Formation in the Gas-Phase Unexpected Photooxidation of Isoprene, Science, 2009, 325, 730-733.

- 6 D. Gu, A. B. Guenther, J. E. Shilling, H. Yu, M. Huang, C. Zhao, Q. Yang, S. T. Martin, P. Artaxo, S. Kim, R. Seco, T. Stavrakou, K. M. Longo, J. Tota, R. A. F. de Souza, O. Vega, Y. Liu, M. Shrivastava, E. G. Alves, F. C. Santos, G. Leng and Z. Hu, Airborne observations reveal elevational gradient in tropical forest isoprene emissions, Nat. Commun., 2017, 8, 15541.
- 7 Q. Chen, D. K. Farmer, J. Schneider, S. R. Zorn, C. L. Heald, 10 T. G. Karl, A. Guenther, J. D. Allan, N. Robinson, H. Coe, J. R. Kimmel, T. Pauliquevis, S. Borrmann, U. Pöschl, M. O. Andreae, P. Artaxo, J. L. Jimenez and S. T. Martin, Mass spectral characterization of submicron biogenic 15 organic particles in the Amazon Basin, Geophys. Res. Lett., 2009, 36, L20806.
- 8 S. T. Martin, M. O. Andreae, P. Artaxo, D. Baumgardner, Q. Chen, A. H. Goldstein, A. Guenther, C. L. Heald, O. L. Mayol-Bracero, P. H. McMurry, T. Pauliquevis, 20 U. Pöschl, K. A. Prather, G. C. Roberts, S. R. Saleska, M. A. Silva Dias, D. V. Spracklen, E. Swietlicki and I. Trebs, Sources and properties of Amazonian aerosol particles, Rev. Geophys., 2010, 48, RG2002.
- 9 N. H. Robinson, J. F. Hamilton, J. D. Allan, B. Langford, 25 D. E. Oram, Q. Chen, K. Docherty, D. K. Farmer, J. L. Jimenez, M. W. Ward, C. N. Hewitt, M. H. Barley, M. E. Jenkin, A. R. Rickard, S. T. Martin, G. McFiggans and H. Coe, Evidence for a significant proportion of Secondary Organic Aerosol from isoprene above a maritime tropical 30 forest, Atmos. Chem. Phys., 2011, 11, 1039-1050.
- 10 Q. Chen, D. K. Farmer, L. V. Rizzo, T. Pauliquevis, M. Kuwata, T. G. Karl, A. Guenther, J. D. Allan, H. Coe, M. O. Andreae, U. Pöschl, J. L. Jimenez, P. Artaxo and S. T. Martin, Submicron particle mass concentrations and sources in the Amazonian wet season (AMAZE-08), Atmos. Chem. Phys., 2015, 15, 3687-3701.
- 11 W. W. Hu, P. Campuzano-Jost, B. B. Palm, D. A. Day, A. M. Ortega, P. L. Hayes, J. E. Krechmer, Q. Chen, 40M. Kuwata, Y. J. Liu, S. S. de Sá, K. McKinney, S. T. Martin, M. Hu, S. H. Budisulistiorini, M. Riva, J. D. Surratt, J. M. St. Clair, G. Isaacman-Van Wertz, L. D. Yee, A. H. Goldstein, S. Carbone, J. Brito, P. Artaxo, J. A. de Gouw, A. Koss, A. Wisthaler, T. Mikoviny, T. Karl, L. Kaser, 45 W. Jud, A. Hansel, K. S. Docherty, M. L. Alexander, N. H. Robinson, H. Coe, J. D. Allan, M. R. Canagaratna, F. Paulot and J. L. Jimenez, Characterization of a real-time tracer for isoprene epoxydiols-derived secondary organic aerosol (IEPOX-SOA) from aerosol mass spectrometer 50 measurements, Atmos. Chem. Phys., 2015, 15, 11807-11833.
- 12 J. Rissler, E. Swietlicki, J. Zhou, G. Roberts, M. O. Andreae, L. V. Gatti and P. Artaxo, Physical properties of the submicrometer aerosol over the Amazon rain forest during the wet-to-dry season transition - comparison of modeled and 55 measured CCN concentrations, Atmos. Chem. Phys., 2004, 4, 2119-2143.
- 13 M. L. Pöhlker, C. Pöhlker, F. Ditas, T. Klimach, I. Hrabe de Angelis, A. Araújo, J. Brito, S. Carbone, Y. Cheng, X. Chi,

1

5

35

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35

40

45

- R. Ditz, S. S. Gunthe, J. Kesselmeier, T. Könemann,
 J. V. Lavrič, S. T. Martin, E. Mikhailov, D. Moran-Zuloaga,
 D. Rose, J. Saturno, H. Su, R. Thalman, D. Walter, J. Wang,
 S. Wolff, H. M. J. Barbosa, P. Artaxo, M. O. Andreae and
 U. Pöschl, Long-term observations of cloud condensation
 nuclei in the Amazon rain forest Part 1: Aerosol size
 distribution, hygroscopicity, and new model
 parametrizations for CCN prediction, *Atmos. Chem. Phys.*,
 2016, 16, 15709–15740.
- 14 R. Thalman, S. S. de Sá, B. B. Palm, H. M. J. Barbosa, M. L. Pöhlker, M. L. Alexander, J. Brito, S. Carbone, P. Castillo, D. A. Day, C. Kuang, A. Manzi, N. L. Ng, A. J. Sedlacek Iii, R. Souza, S. Springston, T. Watson, C. Pöhlker, U. Pöschl, M. O. Andreae, P. Artaxo,
- ¹⁵ J. L. Jimenez, S. T. Martin and J. Wang, CCN activity and organic hygroscopicity of aerosols downwind of an urban region in central Amazonia: seasonal and diel variations and impact of anthropogenic emissions, *Atmos. Chem. Phys.*, 2017, **17**, 11779–11801.
- ²⁰ 15 J. D. Surratt, A. W. H. Chan, N. C. Eddingsaas, M. Chan, C. L. Loza, A. J. Kwan, S. P. Hersey, R. C. Flagan, P. O. Wennberg and J. H. Seinfeld, Reactive intermediates revealed in secondary organic aerosol formation from isoprene, *Proc. Natl. Acad. Sci. U. S. A.*, 2010, **107**, 6640–6645.
- ²⁵ 16 Y. H. Lin, Z. F. Zhang, K. S. Docherty, H. F. Zhang, S. H. Budisulistiorini, C. L. Rubitschun, S. L. Shaw, E. M. Knipping, E. S. Edgerton, T. E. Kleindienst, A. Gold and J. D. Surratt, Isoprene Epoxydiols as Precursors to Secondary Organic Aerosol Formation: Acid-Catalyzed Reactive Uptake Studies with Authentic Compounds, *Environ. Sci. Technol.*, 2012, 46, 250–258.
 - 17 Y. Gómez-González, J. D. Surratt, F. Cuyckens, R. Szmigielski, R. Vermeylen, M. Jaoui, M. Lewandowski,
 - J. H. Offenberg, T. E. Kleindienst, E. O. Edney, F. Blockhuys, C. Van Alsenoy, W. Maenhaut and M. Claeys, Characterization of organosulfates from the photooxidation of isoprene and unsaturated fatty acids in ambient aerosol using liquid chromatography/(–) electrospray ionization mass spectrometry, *J. Mass Spectrom.*, 2008, **43**, 371–382.
 - 18 M. Riva, S. H. Budisulistiorini, Y. Chen, Z. Zhang, E. L. D'Ambro, X. Zhang, A. Gold, B. J. Turpin, J. A. Thornton, M. R. Canagaratna and J. D. Surratt, Chemical Characterization of Secondary Organic Aerosol from Oxidation of Isoprene Hydroxyhydroperoxides, *Environ. Sci. Technol.*, 2016, **50**, 9889–9899.
 - 19 M. Riva, S. H. Budisulistiorini, Z. Zhang, A. Gold,
- J. A. Thornton, B. J. Turpin and J. D. Surratt, Multiphase reactivity of gaseous hydroperoxide oligomers produced from isoprene ozonolysis in the presence of acidified aerosols, *Atmos. Environ.*, 2017, **152**, 314–322.
 - 20 M. Riva, D. M. Bell, A.-M. K. Hansen, G. T. Drozd, Z. Zhang, A. Gold, D. Imre, J. D. Surratt, M. Glasius and A. Zelenyuk,
- 55 A. Gold, D. Imre, J. D. Surratt, M. Glasius and A. Zelenyuk, Effect of Organic Coatings, Humidity and Aerosol Acidity on Multiphase Chemistry of Isoprene Epoxydiols, *Environ. Sci. Technol.*, 2016, **50**, 5580–5588.

- 21 Y. Zhang, Y. Chen, A. T. Lambe, N. E. Olson, Z. Lei, R. L. Craig, Z. Zhang, A. Gold, T. B. Onasch, J. T. Jayne, D. R. Worsnop, C. J. Gaston, J. A. Thornton, W. Vizuete, A. P. Ault and J. D. Surratt, Effect of the Aerosol-Phase State on Secondary Organic Aerosol Formation from the Reactive Uptake of Isoprene-Derived Epoxydiols (IEPOX), *Environ. Sci. Technol. Lett.*, 2018, 5, 167–174.
- 22 C. R. Hoyle, M. Boy, N. M. Donahue, J. L. Fry, M. Glasius,
 A. Guenther, A. G. Hallar, K. H. Hartz, M. D. Petters,
 T. Petaja, T. Rosenoern and A. P. Sullivan, A review of the anthropogenic influence on biogenic secondary organic aerosol, *Atmos. Chem. Phys.*, 2011, 11, 321–343.
- 23 A. G. Carlton, R. W. Pinder, P. V. Bhave and G. A. Pouliot, To What Extent Can Biogenic SOA be Controlled?, *Environ. Sci. Technol.*, 2010, **44**, 3376–3380.
- 24 D. V. Spracklen, J. L. Jimenez, K. S. Carslaw, D. R. Worsnop, M. J. Evans, G. W. Mann, Q. Zhang, M. R. Canagaratna, J. Allan, H. Coe, G. McFiggans, A. Rap and P. Forster, Aerosol mass spectrometer constraint on the global secondary organic aerosol budget, *Atmos. Chem. Phys.*, 2011, 11, 12109–12136.
- 25 A. H. Goldstein, C. D. Koven, C. L. Heald and I. Y. Fung, Biogenic carbon and anthropogenic pollutants combine to form a cooling haze over the southeastern United States, 2009, **106**, 8835–8840.
- 26 B. Nozière, M. Kalberer, M. Claeys, J. Allan, B. D'Anna, S. Decesari, E. Finessi, M. Glasius, I. Grgić, J. F. Hamilton, T. Hoffmann, Y. Iinuma, M. Jaoui, A. Kahnt, C. J. Kampf, I. Kourtchev, W. Maenhaut, N. Marsden, S. Saarikoski, J. Schnelle-Kreis, J. D. Surratt, S. Szidat, R. Szmigielski and A. Wisthaler, The Molecular Identification of Organic Compounds in the Atmosphere: State of the Art and Challenges, *Chem. Rev.*, 2015, 115, 3919–3983.
- 27 M. Glasius and A. H. Goldstein, Recent Discoveries and 35
 Future Challenges in Atmospheric Organic Chemistry, *Environ. Sci. Technol.*, 2016, 50, 2754–2764.
- 28 S. T. Martin, P. Artaxo, L. A. T. Machado, A. O. Manzi, R. A. F. Souza, C. Schumacher, J. Wang, M. O. Andreae, 40H. M. J. Barbosa, J. Fan, G. Fisch, A. H. Goldstein, A. Guenther, J. L. Jimenez, U. Pöschl, M. A. Silva Dias, J. N. Smith and M. Wendisch, Introduction: Observations and Modeling of the Green Ocean Amazon (GoAmazon2014/5), Atmos. Chem. Phys., 2016, 16, 4785-45 4797.
- 29 S. T. Martin, P. Artaxo, L. Machado, A. O. Manzi, R. A. F. Souza, C. Schumacher, J. Wang, T. Biscaro, J. Brito, A. Calheiros, K. Jardine, A. Medeiros, B. Portela, S. S. d. Sá, K. Adachi, A. C. Aiken, R. Albrecht, L. Alexander, M. O. Andreae, H. M. J. Barbosa, P. Buseck, D. Chand, J. M. Comstock, D. A. Day, M. Dubey, J. Fan, J. Fast, G. Fisch, E. Fortner, S. Giangrande, M. Gilles, A. H. Goldstein, A. Guenther, J. Hubbe, M. Jensen, J. L. Jimenez, F. N. Keutsch, S. Kim, C. Kuang, A. Laskin, K. McKinney, F. Mei, M. Miller, R. Nascimento, T. Pauliquevis, M. Pekour, J. Peres, T. Petäjä, C. Pöhlker, U. Pöschl, L. Rizzo, B. Schmid, J. E. Shilling, M. A. S. Dias, I. N. Smith, J. M. Tomlinson, J. Tóta and M. Wendisch,

1

5

15

20

The Green Ocean Amazon Experiment (GoAmazon2014/5) Observes Pollution Affecting Gases, Aerosols, Clouds, and Rainfall over the Rain Forest, *Bull. Am. Meteorol. Soc.*, 2017, **98**, 981–997.

- 30 S. A. Abou Rafee, L. D. Martins, A. B. Kawashima, D. S. Almeida, M. V. B. Morais, R. V. A. Souza, M. B. L. Oliveira, R. A. F. Souza, A. S. S. Medeiros, V. Urbina, E. D. Freitas, S. T. Martin and J. A. Martins, Contributions of mobile, stationary and biogenic sources to air pollution in the Amazon rainforest: a numerical study with the WRF-Chem model, *Atmos. Chem. Phys.*, 2017, 17, 7977–7995.
 - 31 M. O. Andreae and T. W. Andreae, The cycle of biogenic sulfur compounds over the Amazon Basin: 1. Dry season, J. Geophys. Res.: Atmos., 1988, 93, 1487–1497.
 - 32 M. O. Andreae, H. Berresheim, H. Bingemer, D. J. Jacob, B. L. Lewis, S. M. Li and R. W. Talbot, The atmospheric sulfur cycle over the Amazon Basin: 2. Wet season, *J. Geophys. Res.: Atmos.*, 1990, **95**, 16813–16824.
- 33 S. F. Watts, The mass budgets of carbonyl sulfide, dimethyl sulfide, carbon disulfide and hydrogen sulfide, *Atmos. Environ.*, 2000, 34, 761–779.
 - 34 K. Jardine, A. M. Yañez-Serrano, J. Williams, N. Kunert,
 - A. Jardine, T. Taylor, L. Abrell, P. Artaxo, A. Guenther,
 C. N. Hewitt, E. House, A. P. Florentino, A. Manzi,
 N. Higuchi, J. Kesselmeier, T. Behrendt, P. R. Veres,
 B. Derstroff, J. D. Fuentes, S. T. Martin and M. O. Andreae,
 Dimethyl sulfide in the Amazon rain forest, *Global* Biogeochem. Cycles, 2015, 29, 19–32.
 - 35 S. S. de Sá, B. B. Palm, P. Campuzano-Jost, D. A. Day,
 M. K. Newburn, W. Hu, G. Isaacman-VanWertz, L. D. Yee,
 R. Thalman, J. Brito, S. Carbone, P. Artaxo,
 A. H. Goldstein, A. O. Manzi, R. A. F. Souza, F. Mei,
- J. E. Shilling, S. R. Springston, J. Wang, J. D. Surratt, M. L. Alexander, J. L. Jimenez and S. T. Martin, Influence of urban pollution on the production of organic particulate matter from isoprene epoxydiols in central Amazonia, *Atmos. Chem. Phys.*, 2017, 17, 6611–6629.
- ⁴⁰ 36 Y. Liu, J. Brito, M. R. Dorris, J. C. Rivera-Rios, R. Seco, K. H. Bates, P. Artaxo, S. Duvoisin, F. N. Keutsch, S. Kim, A. H. Goldstein, A. B. Guenther, A. O. Manzi, R. A. F. Souza, S. R. Springston, T. B. Watson, K. A. McKinney and S. T. Martin, Isoprene photochemistry over the Amazon rainforest, *Proc. Natl. Acad. Sci. U. S. A.*, 2016, **113**, 6125–6130.
 - 37 I. Kourtchev, R. H. M. Godoi, S. Connors, J. G. Levine,
 A. T. Archibald, A. F. L. Godoi, S. L. Paralovo,
 C. G. G. Barbosa, R. A. F. Souza, A. O. Manzi, R. Seco,
 S. Sjostedt, J. H. Park, A. Guenther, S. Kim, J. Smith,
 S. T. Martin and M. Kalberer, Molecular composition of
 organic aerosols in central Amazonia: an ultra-highresolution mass spectrometry study, *Atmos. Chem. Phys.*,
 - 2016, 16, 11899–11913.
 38 D. F. Zhao, A. Buchholz, B. Kortner, P. Schlag, F. Rubach, H. Fuchs, A. Kiendler-Scharr, R. Tillmann, A. Wahner,
 - A. K. Watne, M. Hallquist, J. M. Flores, Y. Rudich, K. Kristensen, A. M. K. Hansen, M. Glasius, I. Kourtchev,

M. Kalberer and T. F. Mentel, Cloud condensation nuclei activity, droplet growth kinetics, and hygroscopicity of biogenic and anthropogenic secondary organic aerosol (SOA), *Atmos. Chem. Phys.*, 2016, **16**, 1105–1121.

- 39 M. P. Tolocka and B. Turpin, Contribution of Organosulfur Compounds to Organic Aerosol Mass, *Environ. Sci. Technol.*, 2012, 46, 7978–7983.
- 40 J. Liao, K. D. Froyd, D. M. Murphy, F. N. Keutsch, G. Yu, P. O. Wennberg, J. M. St. Clair, J. D. Crounse, A. Wisthaler, T. Mikoviny, J. L. Jimenez, P. Campuzano-Jost, D. A. Day,
 W. Hu, T. B. Ryerson, I. B. Pollack, J. Peischl, B. E. Anderson, L. D. Ziemba, D. R. Blake, S. Meinardi and G. Diskin, Airborne measurements of organosulfates over the continental U.S, *J. Geophys. Res.: Atmos.*, 2015, 120, 2990–3005.
- 41 M. Glasius, A. M. K. Hansen, M. Claeys, J. S. Henzing, A. D. Jedynska, A. Kasper-Giebl, M. Kistler, K. Kristensen, J. Martinsson, W. Maenhaut, J. K. Nøjgaard, G. Spindler, K. E. Stenström, E. Swietlicki, S. Szidat, D. Simpson and K. E. Yttri, Composition and sources of carbonaceous aerosols in Northern Europe during winter, *Atmos. Environ.*, 2018, **173**, 127–141.
- 42 K. Kristensen and M. Glasius, Organosulfates and oxidation products from biogenic hydrocarbons in fine aerosols from a forest in North West Europe during spring, *Atmos. Environ.*, 2011, **45**, 4546–4556.
- 43 Q. T. Nguyen, M. K. Christensen, F. Cozzi, A. Zare, A. M. K. Hansen, K. Kristensen, T. E. Tulinius, H. H. Madsen, J. H. Christensen, J. Brandt, A. Massling, J. K. Nøjgaard and M. Glasius, Understanding the anthropogenic influence on formation of biogenic secondary organic aerosols in Denmark *via* analysis of organosulfates and related oxidation products, *Atmos. Chem. Phys.*, 2014, 14, 8961–8981.
- 44 T. S. Christoffersen, J. Hjorth, O. Horie, N. R. Jensen, D. Kotzias, L. L. Molander, P. Neeb, L. Ruppert, R. Winterhalter, A. Virkkula, K. Wirtz and B. R. Larsen, *cis*-Pinic acid, a possible precursor for organic aerosol formation from ozonolysis of α -pinene, *Atmos. Environ.*, 1998, **32**, 1657–1661.
- 45 J. Z. Yu, D. R. Cocker, R. J. Griffin, R. C. Flagan and J. H. Seinfeld, Gas-phase ozone oxidation of monoterpenes: Gaseous and particulate products, *J. Atmos. Chem.*, 1999, 34, 207–258.
- 46 M. Glasius, M. Lahaniati, A. Calogirou, D. Di Bella, N. R. Jensen, J. Hjorth, D. Kotzias and B. R. Larsen, Carboxylic acids in secondary aerosols from oxidation of cyclic monoterpenes by ozone, *Environ. Sci. Technol.*, 2000, 34, 1001–1010.
- 47 M. Claeys, Y. Iinuma, R. Szmigielski, J. D. Surratt, F. Blockhuys, C. Van Alsenoy, O. Böge, B. Sierau, Y. Gómez-González, R. Vermeylen, P. Van der Veken, M. Shahgholi, A. W. H. Chan, H. Herrmann, J. H. Seinfeld and W. Maenhaut, Terpenylic acid and related compounds from the oxidation of alpha-pinene: Implications for new particle formation and growth above forests, *Environ. Sci. Technol.*, 2009, 43, 6976–6982.

Paper

1

15

25

30

50

55

45

50

40

25

1

- 48 R. Szmigielski, J. D. Surratt, Y. Gomez-Gonzalez, P. Van der Veken, I. Kourtchev, R. Vermeylen, F. Blockhuys, M. Jaoui, T. E. Kleindienst, M. Lewandowski, J. H. Offenberg, E. O. Edney, J. H. Seinfeld, W. Maenhaut and M. Claeys, 3-Methyl-1,2,3-butanetricarboxylic acid: An atmospheric tracer for terpene secondary organic aerosol, *Geophys. Res.*
- *Lett.*, 2007, **34**.

5

10

15

2.0

25

40

45

50

55

- 49 L. Müller, M. C. Reinnig, K. H. Naumann, H. Saathoff, T. F. Mentel, N. M. Donahue and T. Hoffmann, Formation of 3-methyl-1,2,3-butanetricarboxylic acid *via* gas phase oxidation of pinonic acid – a mass spectrometric study of SOA aging, *Atmos. Chem. Phys.*, 2012, **12**, 1483–1496.
- 50 A. F. Stein, R. R. Draxler, G. D. Rolph, B. J. B. Stunder, M. D. Cohen and F. Ngan, NOAA's HYSPLIT Atmospheric Transport and Dispersion Modeling System, *Bull. Am. Meteorol. Soc.*, 2015, **96**, 2059–2077.
- 51 K. Kristensen, M. Bilde, P. P. Aalto, T. Petaja and M. Glasius, Denuder/filter sampling of organic acids and organosulfates at urban and boreal forest sites: Gas/particle distribution and possible sampling artifacts, *Atmos. Environ.*, 2016, **130**, 36–53.
- 52 A. M. K. Hansen, K. Kristensen, Q. T. Nguyen, A. Zare, F. Cozzi, J. K. Nøjgaard, H. Skov, J. Brandt, J. H. Christensen, J. Ström, P. Tunved, R. Krejci and M. Glasius, Organosulfates and organic acids in Arctic aerosols: speciation, annual variation and concentration levels, *Atmos. Chem. Phys.*, 2014, 14, 7807–7823.
- 53 S. H. Budisulistiorini, X. Li, S. T. Bairai, J. Renfro, Y. Liu,
- Y. J. Liu, K. A. McKinney, S. T. Martin, V. F. McNeill, H. O. T. Pye, A. Nenes, M. E. Neff, E. A. Stone, S. Mueller, C. Knote, S. L. Shaw, Z. Zhang, A. Gold and J. D. Surratt, Examining the effects of anthropogenic emissions on isoprene-derived secondary organic aerosol formation during the 2013 Southern Oxidant and Aerosol Study (SOAS) at the Look Rock, Tennessee ground site, *Atmos. Chem. Phys.*, 2015, 15, 8871–8888.
 - 54 Q. T. Nguyen, M. K. Christensen, F. Cozzi, A. Zare,
 A. M. K. Hansen, K. Kristensen, T. E. Tulinius,
 H. H. Madsen, J. H. Christensen, J. Brandt, A. Massling,
 J. K. Nojgaard and M. Glasius, Understanding the anthropogenic influence on formation of biogenic secondary organic aerosols in Denmark *via* analysis of
 - organosulfates and related oxidation products, *Atmos. Chem. Phys.*, 2014, **14**, 8961–8981.
 - 55 P. F. DeCarlo, J. R. Kimmel, A. Trimborn, M. J. Northway, J. T. Jayne, A. C. Aiken, M. Gonin, K. Fuhrer, T. Horvath, K. S. Docherty, D. R. Worsnop and J. L. Jimenez, Fielddeployable, high-resolution, time-of-flight aerosol mass spectrometer, *Anal. Chem.*, 2006, **78**, 8281–8289.
 - 56 M. R. Canagaratna, J. T. Jayne, J. L. Jimenez, J. D. Allan,
 M. R. Alfarra, Q. Zhang, T. B. Onasch, F. Drewnick, H. Coe,
 A. Middlebrook, A. Delia, L. R. Williams, A. M. Trimborn,
 M. J. Northway, P. F. DeCarlo, C. E. Kolb, P. Davidovits and
 D. R. Worsnop, Chemical and microphysical
 - characterization of ambient aerosols with the aerodyne aerosol mass spectrometer, *Mass Spectrom. Rev.*, 2007, **26**, 185–222.

- 57 D. K. Farmer, A. Matsunaga, K. S. Docherty, J. D. Surratt, J. H. Seinfeld, P. J. Ziemann and J. L. Jimenez, Response of an aerosol mass spectrometer to organonitrates and organosulfates and implications for atmospheric chemistry, *Proc. Natl. Acad. Sci. U. S. A.*, 2010, **107**, 6670– 6675.
- 58 J. D. Surratt, Y. Gomez-Gonzalez, A. W. H. Chan, R. Vermeylen, M. Shahgholi, T. E. Kleindienst, E. O. Edney, J. H. Offenberg, M. Lewandowski, M. Jaoui, W. Maenhaut, M. Claeys, R. C. Flagan and J. H. Seinfeld, Organosulfate 10 formation in biogenic secondary organic aerosol, *J. Phys. Chem. A*, 2008, **112**, 8345–8378.
- 59 M. S. Shalamzari, O. Ryabtsova, A. Kahnt, R. Vermeylen, M.-F. Hérent, J. Quetin-Leclercq, P. Van der Veken, W. Maenhaut and M. Claeys, Mass spectrometric ¹⁵ characterization of organosulfates related to secondary organic aerosol from isoprene, *Rapid Commun. Mass Spectrom.*, 2013, 27, 784–794.
- 60 J. D. Surratt, J. H. Kroll, T. E. Kleindienst, E. O. Edney, M. Claeys, A. Sorooshian, N. L. Ng, J. H. Offenberg, M. Lewandowski, M. Jaoui, R. C. Flagan and J. H. Seinfeld, Evidence for organosulfates in secondary organic aerosol, *Environ. Sci. Technol.*, 2007, 41, 517–527.
- 61 C. N. Olson, M. M. Galloway, G. Yu, C. J. Hedman, M. R. Lockett, T. Yoon, E. A. Stone, L. M. Smith and F. N. Keutsch, Hydroxycarboxylic Acid-Derived Organosulfates: Synthesis, Stability, and Quantification in Ambient Aerosol, *Environ. Sci. Technol.*, 2011, 45, 6468–6474.
- 62 J. Schindelka, Y. Iinuma, D. Hoffmann and H. Herrmann, 30
 Sulfate radical-initiated formation of isoprene-derived organosulfates in atmospheric aerosols, *Faraday Discuss.*, 2013, 165, 237–259.
- 63 A. P. S. Hettiyadura, T. Jayarathne, K. Baumann,
 A. H. Goldstein, J. A. de Gouw, A. Koss, F. N. Keutsch, 35
 K. Skog and E. A. Stone, Qualitative and quantitative analysis of atmospheric organosulfates in Centreville, Alabama, *Atmos. Chem. Phys.*, 2017, 17, 1343–1359.
- 64 Y.-H. Lin, H. Zhang, H. O. T. Pye, Z. Zhang, W. J. Marth, S. Park, M. Arashiro, T. Cui, S. H. Budisulistiorini, K. G. Sexton, W. Vizuete, Y. Xie, D. J. Luecken, I. R. Piletic, E. O. Edney, L. J. Bartolotti, A. Gold and J. D. Surratt, Epoxide as a precursor to secondary organic aerosol formation from isoprene photooxidation in the presence of nitrogen oxides, *Proc. Natl. Acad. Sci. U. S. A.*, 2013, DOI: 10.1073/pnas.1221150110.
- 65 H. Zhang, D. R. Worton, M. Lewandowski, J. Ortega, C. L. Rubitschun, J.-H. Park, K. Kristensen, P. Campuzano-Jost, D. A. Day, J. L. Jimenez, M. Jaoui, J. H. Offenberg, T. E. Kleindienst, J. Gilman, W. C. Kuster, J. de Gouw, C. Park, G. W. Schade, A. A. Frossard, L. Russell, L. Kaser, W. Jud, A. Hansel, L. Cappellin, T. Karl, M. Glasius, A. Guenther, A. H. Goldstein, J. H. Seinfeld, A. Gold, R. M. Kamens and J. D. Surratt, Organosulfates as tracers for secondary organic aerosol (SOA) formation from 2-methyl-3-buten-2-ol (MBO) in the atmosphere, *Environ. Sci. Technol.*, 2012, 46, 9437–9446.

1

66 M. Riva, S. H. Budisulistiorini, Z. Zhang, A. Gold and J. D. Surratt, Chemical characterization of secondary organic aerosol constituents from isoprene ozonolysis in the presence of acidic aerosol, *Atmos. Environ.*, 2016, **130**, 5–13.

Paper

1

5

20

25

40

45

50

55

- 67 D. R. Worton, J. D. Surratt, B. W. LaFranchi, A. W. H. Chan,Y. Zhao, R. J. Weber, J.-H. Park, J. B. Gilman, J. de Gouw,C. Park, G. Schade, M. Beaver, J. M. S. Clair, J. Crounse,P. Wennberg, G. M. Wolfe, S. Harrold, J. A. Thornton,
- D. K. Farmer, K. S. Docherty, M. J. Cubison, J.-L. Jimenez, A. A. Frossard, L. M. Russell, K. Kristensen, M. Glasius, J. Mao, X. Ren, W. Brune, E. C. Browne, S. E. Pusede, R. C. Cohen, J. H. Seinfeld and A. H. Goldstein, Observational Insights into Aerosol Formation from Isoprene, *Environ. Sci. Technol.*, 2013, 47, 11403–11413.
 - 68 J. D. Surratt, S. M. Murphy, J. H. Kroll, N. L. Ng, L. Hildebrandt, A. Sorooshian, R. Szmigielski, R. Vermeylen, W. Maenhaut, M. Claeys, R. C. Flagan and J. H. Seinfeld, Chemical composition of secondary organic aerosol formed from the photooxidation of isoprene, *J. Phys. Chem. A*, 2006, **110**, 9665–9690.
 - 69 G. Isaacman-VanWertz, L. D. Yee, N. M. Kreisberg, R. Wernis, J. A. Moss, S. V. Hering, S. S. de Sá, S. T. Martin, M. L. Alexander, B. B. Palm, W. Hu, P. Campuzano-Jost, D. A. Day, J. L. Jimenez, M. Riva, J. D. Surratt, J. Viegas, A. Manzi, E. Edgerton, K. Baumann, R. Souza, P. Artaxo and A. H. Goldstein, Ambient Gas-Particle Partitioning of Tracers for Biogenic Oxidation, *Environ. Sci. Technol.*, 2016, **50**, 9952–9962.
- 30 Environ. Sci. Technol., 2016, 50, 9952–9962.
 70 M. S. Shalamzari, A. Kahnt, R. Vermeylen, T. E. Kleindienst, M. Lewandowski, F. Cuyckens, W. Maenhaut and M. Claeys, Characterization of polar organosulfates in secondary organic aerosol from the green leaf volatile 3-Z-hexenal, Environ. Sci. Technol., 2014, 48, 12671–12678.
 - 71 M. Riva, T. Da Silva Barbosa, Y. H. Lin, E. A. Stone, A. Gold and J. D. Surratt, Chemical characterization of organosulfates in secondary organic aerosol derived from the photooxidation of alkanes, *Atmos. Chem. Phys.*, 2016, 16, 11001–11018.
 - 72 M. Riva, S. Tomaz, T. Cui, Y.-H. Lin, E. Perraudin, A. Gold, E. A. Stone, E. Villenave and J. D. Surratt, Evidence for an unrecognized secondary anthropogenic source of organosulfates and sulfonates: Gas-phase oxidation of polycyclic aromatic hydrocarbons in the presence of sulfate aerosol, *Environ. Sci. Technol.*, 2015, **49**, 6654–6664.
 - 73 A. Guenther, C. N. Hewitt, D. Erickson, *et al.*, A global model of natural volatile organic compound emissions, *J. Geophys. Res.: Atmos.*, 1995, **100**, 8873–8892.
 - 74 A. Bougiatioti, P. Nikolaou, I. Stavroulas, G. Kouvarakis,
 R. Weber, A. Nenes, M. Kanakidou and N. Mihalopoulos,
 Particle water and pH in the eastern Mediterranean: source variability and implications for nutrient availability, *Atmos. Chem. Phys.*, 2016, 16, 4579–4591.

- 75 C. J. Gaston, T. P. Riedel, Z. F. Zhang, A. Gold, J. D. Surratt and J. A. Thornton, Reactive Uptake of an Isoprene-Derived Epoxydiol to Submicron Aerosol Particles, *Environ. Sci. Technol.*, 2014, **48**, 11178–11186.
- 76 C. Schulz, J. Schneider, B. Amorim Holanda, O. Appel, A. Costa, S. S. de Sá, V. Dreiling, D. Fütterer, T. Jurkat-Witschas, T. Klimach, M. Krämer, S. T. Martin, S. Mertes, M. L. Pöhlker, D. Sauer, C. Voigt, B. Weinzierl, H. Ziereis, M. Zöger, M. O. Andreae, P. Artaxo, L. A. T. Machado, U. Pöschl, M. Wendisch and S. Borrmann, Aircraft-based observations of isoprene epoxydiol-derived secondary organic aerosol (IEPOX-SOA) in the tropical upper troposphere over the Amazon region, *Atmos. Chem. Phys. Discuss.*, 2018, 18, 1–32.
- 77 M. Shrivastava, M. O. Andreae, P. Artaxo, H. Barbosa, ¹⁵ L. K. Berg, J. Brito, J. Ching, R. C. Easter, J. Fan, J. D. Fast, Z. Feng, J. D. Fuentes, M. Glasius, A. H. Goldstein, D. G. H. Gomes, A. Guenther, S. H. Jathar, S. Kim, Y. Liu, S. Lou, S. T. Martin, V. F. McNeill, A. Medeiros, S. S. de Sá, J. E. Shilling, S. R. Springston, R. F. Souza, J. A. Thornton, G. Isaacman-VanWertz, R. Y. L. Yee, R. A. Zaveri, A. Zelenyuk and C. Zhao, Urban pollution greatly enhances formation of natural aerosols over the Amazon rainforest, *Nat. Commun.*, 2018, submitted.
- 78 Y. Zhang, L. Tang, Y. Sun, O. Favez, F. Canonaco, A. Albinet,
 F. Couvidat, D. Liu, J. T. Jayne, Z. Wang, P. L. Croteau,
 M. R. Canagaratna, H. c. Zhou, A. S. H. Prévôt and
 D. R. Worsnop, Limited formation of isoprene
 epoxydiols-derived secondary organic aerosol under 30
 NO_x-rich environments in Eastern China, *Geophys. Res. Lett.*, 2017, 44, 2035–2043.
- 79 A. G. Carlton and B. J. Turpin, Particle partitioning potential of organic compounds is highest in the Eastern US and driven by anthropogenic water, *Atmos. Chem. Phys.*, 2013, 35 13, 10203–10214.
- 80 H. Guo, L. Xu, A. Bougiatioti, K. M. Cerully, S. L. Capps, J. R. Hite Jr, A. G. Carlton, S. H. Lee, M. H. Bergin, N. L. Ng, A. Nenes and R. J. Weber, Fine-particle water and pH in the southeastern United States, *Atmos. Chem. Phys.*, 2015, 15, 5211–5228.
- 81 L. Xu, H. Y. Guo, C. M. Boyd, M. Klein, A. Bougiatioti, K. M. Cerully, J. R. Hite, G. Isaacman-VanWertz, N. M. Kreisberg, C. Knote, K. Olson, A. Koss, A. H. Goldstein, S. V. Hering, J. de Gouw, K. Baumann, S. H. Lee, A. Nenes, R. J. Weber and N. L. Ng, Effects of anthropogenic emissions on aerosol formation from isoprene and monoterpenes in the southeastern United States, *Proc. Natl. Acad. Sci. U. S. A.*, 2015, **112**, 37–42.
- 82 J. Ye, J. P. D. Abbatt and A. W. H. Chan, Novel pathway of SO_2 oxidation in the atmosphere: reactions with monoterpene ozonolysis intermediates and secondary organic aerosol, *Atmos. Chem. Phys.*, 2018, **18**, 5549–5565.

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